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# **LBC perturbations for a convection-permitting COSMO ensemble system**

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# Introduction

- 3 major uncertainty sources in LAM:
  - lateral boundary conditions (LBCs)
  - initial conditions
  - model physics
- Investigate 3 different methods to perturb the LBC
  - Downscale global EPS
  - Downscale global EPS perturbations
  - Downscale global climatological perturbations (Torn et al., 2006)
- Results shall give a hint which method(s) to use in a future COSMO-E (EPS and EDA) system at MeteoSwiss



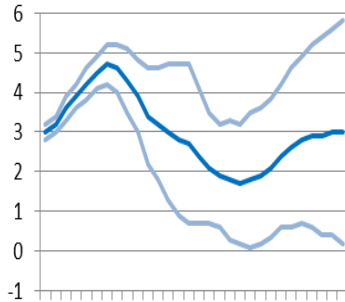
# COSMO-E Configuration

- Operational MeteoSwiss COSMO-2 setup (2.2km, no deep convection scheme)
- Nested directly into ECMWF EPS (32km/20km) and ECMWF DET (16km/10km) model
- 21 ensemble members, integrated out to +120h
- All members started from same operational COSMO-2 analysis
- No physics perturbations
- No obs assimilation



# LBC Perturbation Methods

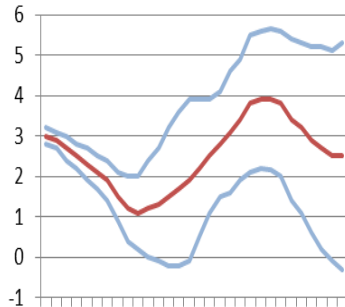
**Method 1**  
«Downscale EPS»



EPS  
EPS  
EPS

$$X_{ijk}^m = \overline{EPS}_{ijk} + \Delta EPS_{ijk}^m$$

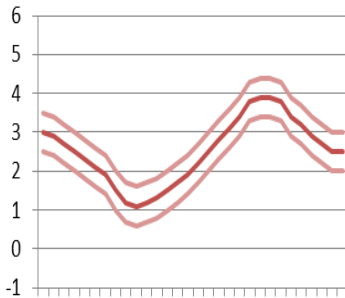
**Method 2**  
«Downscale  
EPS Perturbations»



EPS  
DET  
EPS

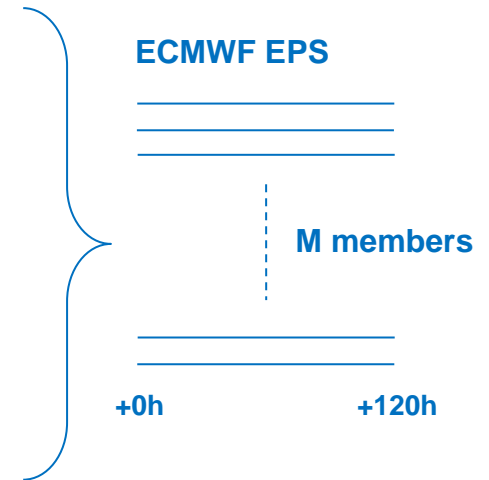
$$X_{ijk}^m = DET_{ijk} + \Delta EPS_{ijk}^m$$

**Method 3**  
«Downscale Clim.  
Perturbations»



CLIM  
DET  
CLIM

$$X_{ijk}^m = DET_{ijk} + \alpha \cdot \Delta CLIM_{ijk}^m$$



Climatological ensemble:  
ECMWF analysis time series

02.01.2007 00UTC  
22.05.2010 00UTC  
14.07.2008 00UTC

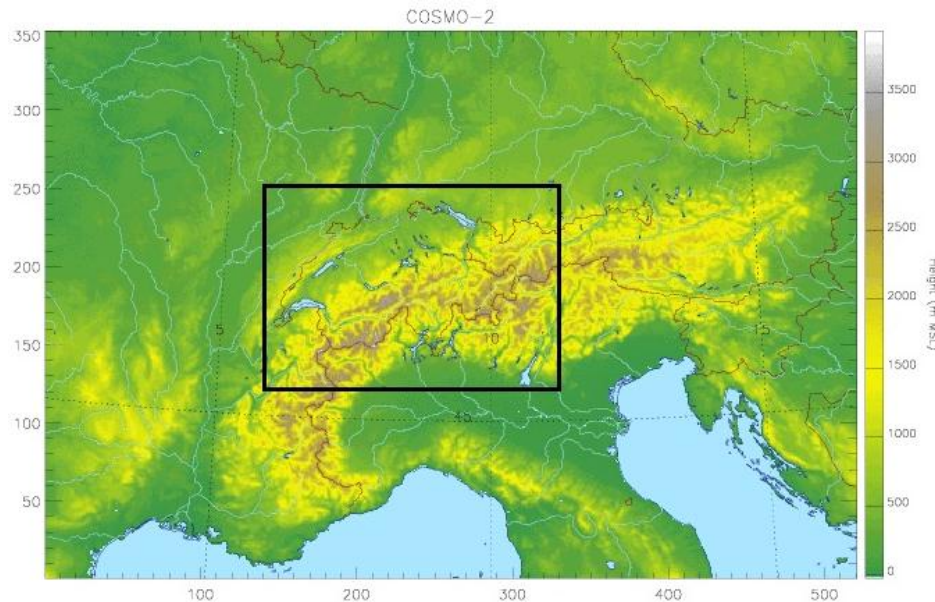
Random dates M members

30.11.2011 00UTC  
05.09.2006 00UTC  
+0h +120h



# Ensemble Assessment

- Results are based on 4 case studies:
  - 2 summer cases (low advection, convection)
  - 2 autumn cases (high advection, large scale forcing)
- Compare temporal and spatial evolution of ensemble mean and spread of different methods

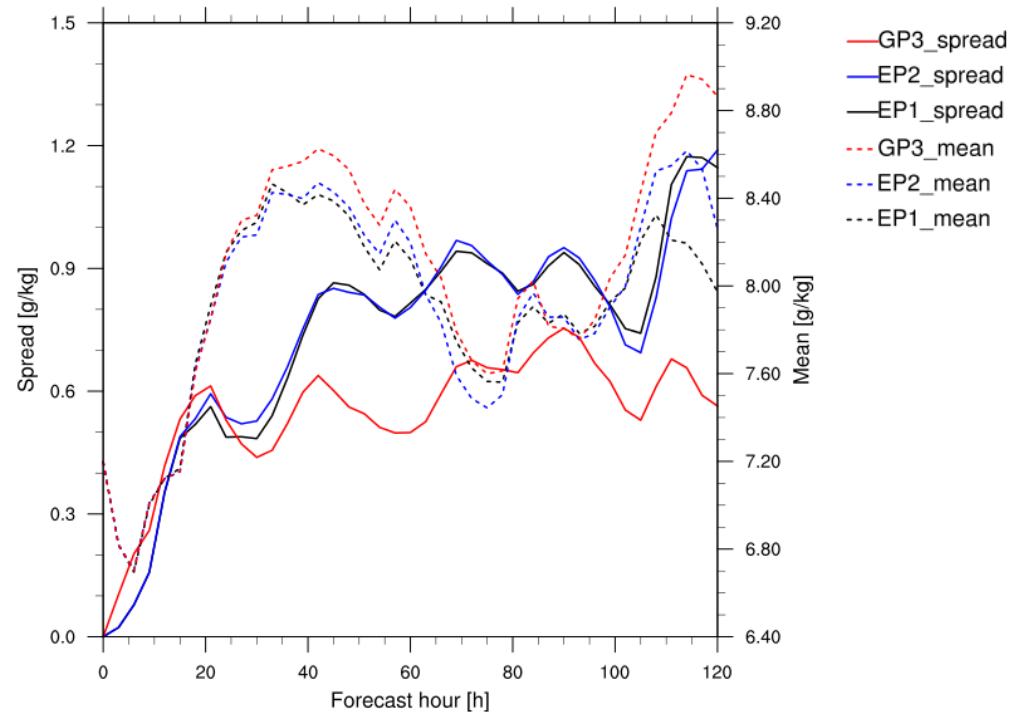
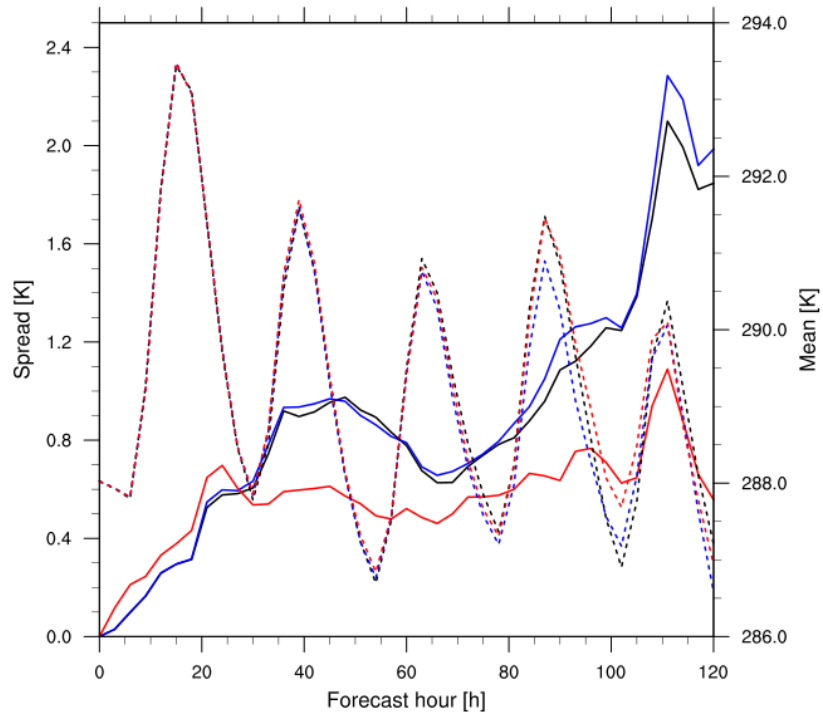




# Temporal Evolution

## Temperature

## Specific humidity



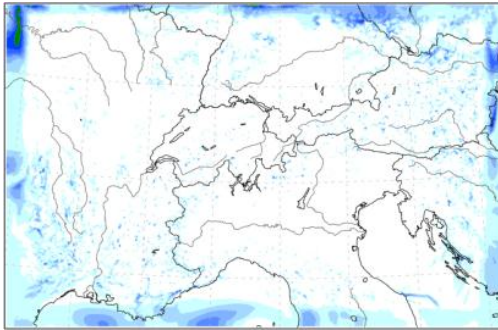
Mean and spread, averaged over SWISS domain,  
 $z = 0\text{m}-1500\text{m}$  (case 1)



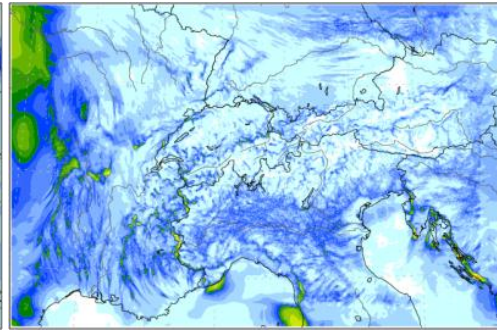
# Spatial Evolution

Method 1

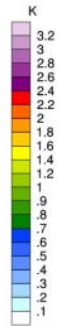
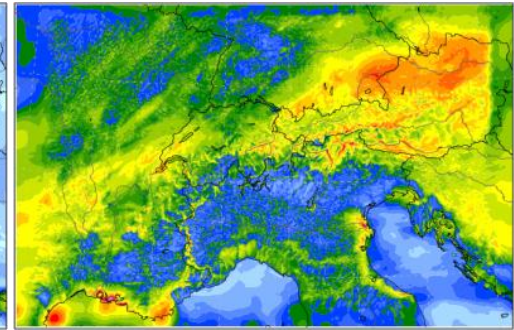
FC +03h



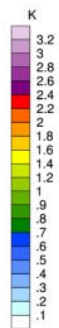
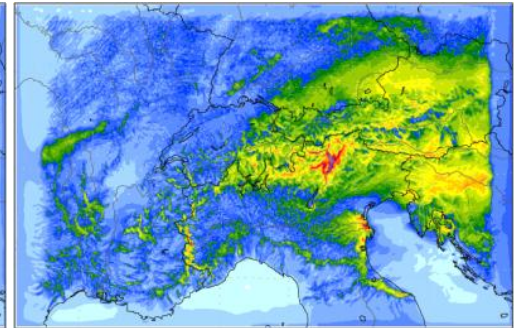
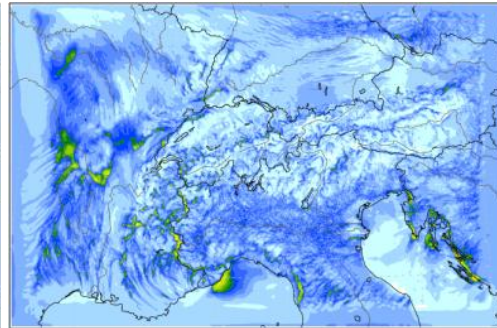
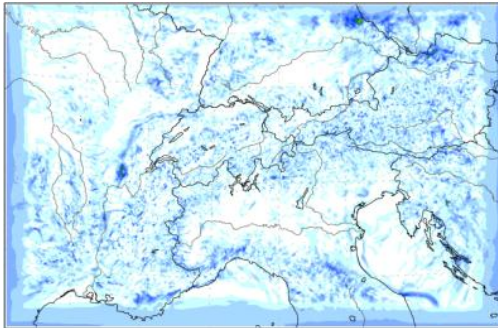
FC +12h



FC +84h



Method 3



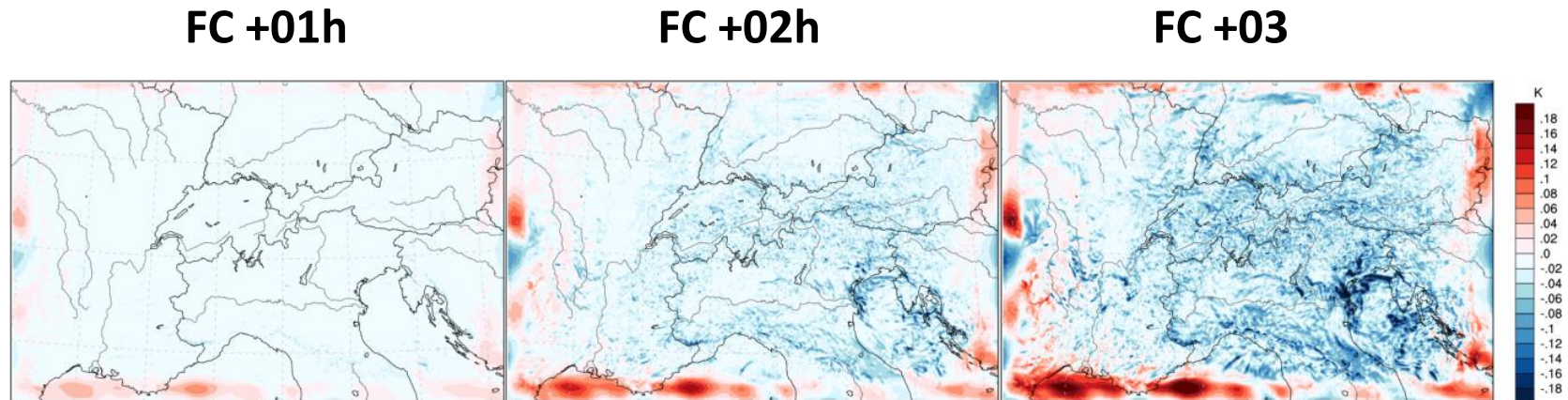
Temperature spread at 500m (case 1)





# Can Method 3 be used for EDA?

- Torn et al. (2006) used method 3 for EDA in an idealized, regional ( $\Delta x \approx 100\text{km}$ ) model environment
- How does it compare to method 1 in a high-resolution model in real case studies?



- Differences in spread, at +03h:  $< 0.07$  K





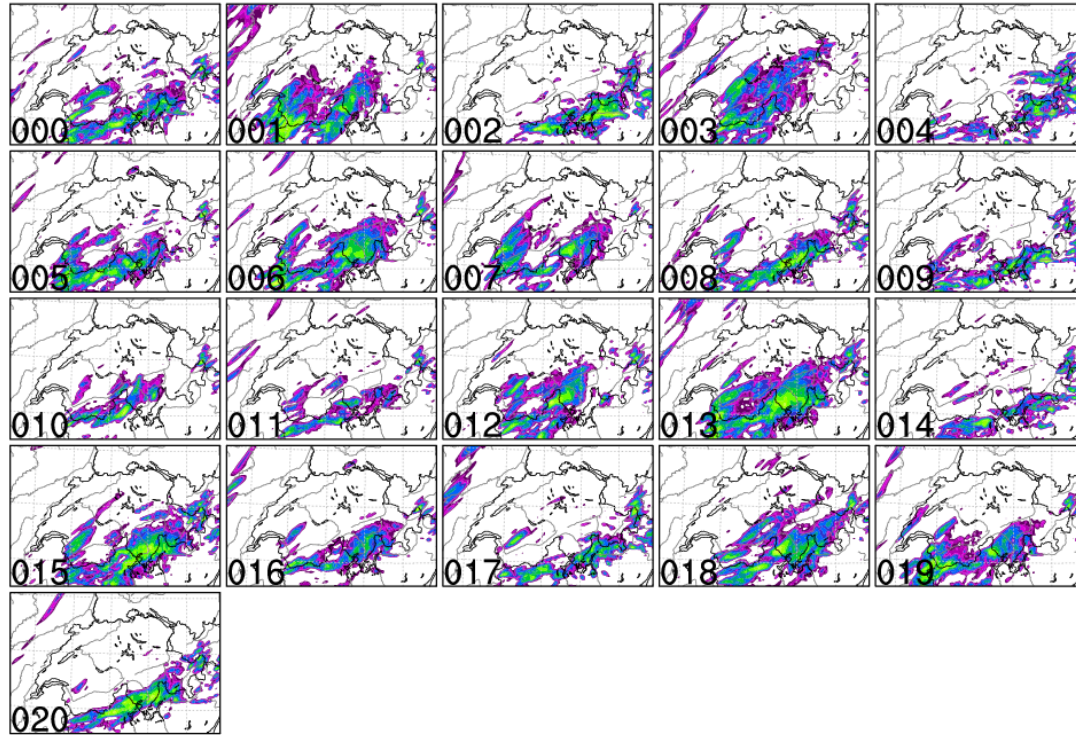
# Further results

- Influence of COSMO-E strongest near surface
  - Upper level mainly downscaling of global model
- No significant model imbalances found (analysis of surface pressure tendencies)
- Sufficient to perturb T, U, V, QV (no large differences with additional perturbation of QI, QC, PP)
- Ensemble results are realistic in comparison with obs.

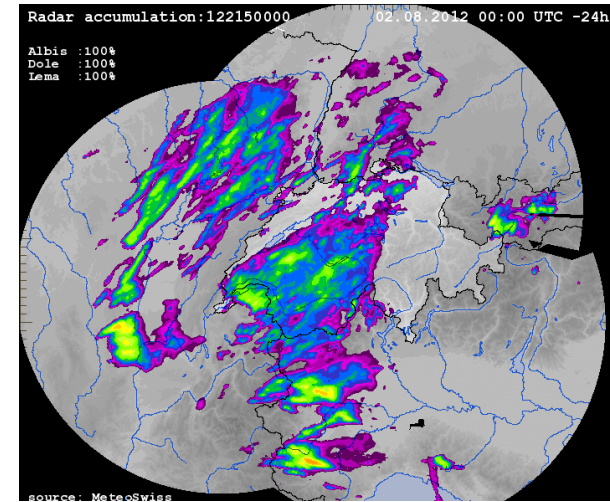
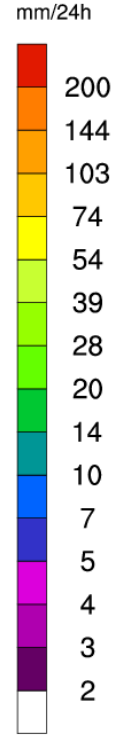


24h Sum of Total Precipitation

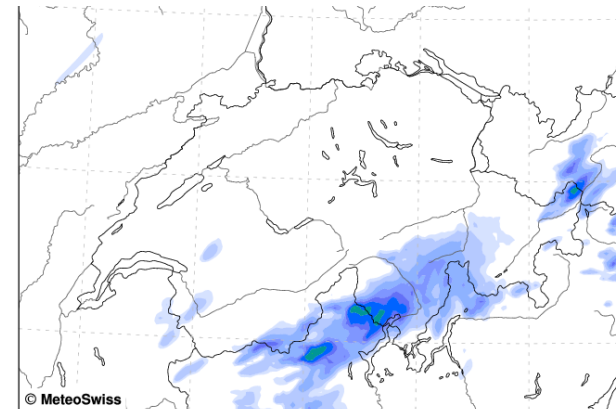
01.08.2012 00UTC +24h



COSMO-E



Observation



COSMO-2



# Conclusions

- **Generally small differences between EP1 and EP2 up to 40h**
- **Method 3 seems suitable for data assimilation, but**
  - **not flow-dependent**
  - **initially faster growth of spread**
    - > **stronger gravity & sound waves?**
- *Only four case studies -> to make general conclusions more cases have to be calculated*



# Outlook

- Use method 2 for EPS and EDA:
  - Flow-dependent pert.
  - High-res. and more frequently updated ensemble mean than using only global EPS (4x vs. 2x per day)
  - If underdispersive: Use scaled perturbations to enhance ensemble spread
- Potential use of method 3:
  - Only for EDA
  - Advantageous if more members are required than provided by global EPS