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Investigating effects of latent heating on midlatitude cyclone intensification with potential vorticity diagnostics **Dominik Büeler & Stephan Pfahl**

Introduction Data & Method Simulation of the historical European winter storm *Klaus* in 2009 [10] Features of midlatitude cyclones Midlatitude cyclones are often associated with weather extremes like COSMO model setup: North Atlantic / European domain wind, precipitation, and floodings Spatial resolution: 0.125° x 0.125° (horizontal), 40 levels (vertical) Adiabatic energy source: baroclinic instability [1] Initial and boundary conditions from ECMWF analysis Diabatic energy source: latent heating (LH) during cloud formation \rightarrow Sensitivity simulations: scale specific latent heat constants in the can have a significant contribution to cyclone intensification [2,3,4,5,6] model (lh_v, lh_f, lh_s) with a constant α *Motivation:* Global warming \rightarrow higher atmospheric moisture content Reference (**REF**): $\alpha = 1.0$ \rightarrow more LH \rightarrow effect on midlatitude cyclone intensity and tracks? Reduced / enhanced LH (L05, L15): $\alpha = 0.5$, $\alpha = 1.5$ Potential vorticity (PV) • No LH (**L00**): α = 0.0 Key variable for investigation of cyclone dynamics [7]: Analyze 6h-heating-rates due to microphysics (TMPHYS) and convection (TCONV) $Q = \frac{1}{2} \cdot \boldsymbol{\eta} \cdot \boldsymbol{\nabla} \boldsymbol{\theta}$ Cyclone area definition:



Conserved under adiabatic conditions but can be changed through frictional and diabatic processes:

 $\frac{DQ}{Dt} = \frac{1}{\rho} \left(\nabla \theta \cdot \nabla \times F \right) + \frac{1}{\rho} \left(\eta \cdot \nabla \dot{\theta} \right)$

PV distribution in midlatitude cyclones [3,8,9]:



- Radius of 200 km around SLP minimum (for PV budget)
- Rectangular box including SLP minima at times t_0 and t_{-6} (for heating) rate budgets)

Klaus, 2009



• One of the most damaging storms for Northern Iberia and Southern France with extreme winds and heavy precipitation [10] Explosive deepening

after crossing a region of strong upper-level divergence (left exit region of jet streak) [10] • LH strongly contributed to the

intensification [11]

Results



Science

Climate

0

Cyclone area examples (radius and box)

Conclusions & Outlook

Very weak cyclone without LH

maximum \rightarrow generated PV is lifted upwards

with strongest LH (structure of warm conveyor belt outflow)

onclusions & Outlook	References
The dynamic contribution of LH was significant during the explosive intensification of <i>Klaus</i> Both the diabatically produced positive low-tropospheric and negative upper-tropospheric PV anomalies of <i>Klaus</i> scale approximately linearly with intensified LH With more intense LH, the track of <i>Klaus</i> shifts slightly southward during the intensification phase The vertical PV distribution is a good diagnostic metric to investigate the effect of different LH conditions on the intensity and dynamics of midlatitude cyclones The result of this study motivates to use the PV budget to further investigate the LH sensitivity of midlatitude cyclones both on a case-to- case and climatological basis	 Eady, E. T., 1949: Long waves and cyclone waves. Tellus, 1 (3), 33-52. Kuo, YH., Shapiro, M. A., and Donall, E. G., 1991: The interaction between baroclinic and diabatic processes in a numerical simulation of a rapidly intensifying extratropical marine cyclone. Mon. Wea. Rev., 119, 368-384. Davis, C. A. and Emanuel, K. A., 1991: Potential vorticity diagnostics of cyclogenesis. Mon. Wea. Rev., 119, 1929-1953. Davis, C. A., 1992: A potential-vorticity diagnosis of the importance of initial structure and condensational heating in observed extratropical cyclogenesis. Mon. Wea. Rev., 120, 2409-2428. Stoelinga, M. T., 1996: A potential vorticity-based study of the role of diabatic heating and friction in a numerically simulated baroclinic cyclone. Mon. Wea. Rev., 124, 849-873. Ahmadi-Givi, F., Graig, G. C., and Plant, R. S., 2004: The dynamics of a midlatitude cyclone with very strong latentheat release. Quart. J. Roy. Meteor. Soc., 130, 295-323. Hoskins, B. J., McIntyre, M.E., and Robertson, A. W., 1985: On the use and signicance of isentropic potential vorticity maps. Quart. J. Roy. Meteor. Soc., 111, 877-946. Reed, R. J., Stoelinga, M. T., and Kuo, YH., 1992: A model-aided study of the origin and evolution of the anomalously high potential vorticity in the inner region of a rapidly deepening marine cyclone. Mon. Wea. Rev., 120, 893-913. Campa, J. and Wernli, H., 2012: A PV perspective on the vertical structure of mature midlatitude cyclones in the Northern Hemisphere. J. Atmos. Sci., 69, 725-740. Liberato, M. L. R., Pinto, J. G., Trigo I. F., and Trigo, R. M., 2011: Klaus an exceptional winter storm over northerm lberia and southern France. Weather, 66, 330-334. Fink, A. H., Pohle, S., Pinto, J. G., and Knippertz, P., 2012: Diagnosing the influence of diabatic processes on the explorie dearpartine of extratropical cyclones. Res. Lett. 30.