

Impact of Higher Order Quadratic Conserving Advection Scheme in the COSMO model on Regional climate Simulation

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Outline

- ▶ Introduction
 - ▶ Quadratic conserving scheme in the COSMO model
- ▶ Regional Climate Simulation
 - ▶ Results
- ▶ Conclusions

INTRODUCTION:

Higher order quadratic conserving scheme in COSMO model

Horizontal Advection of u : $u \frac{\partial u}{\partial \lambda} + v \frac{\partial u}{\partial \phi}$

divergence form: conservative a priori

$$S4^* := \frac{9}{8} \overline{u^{O4,\lambda}} \delta_\lambda u^\lambda - \frac{1}{8} \overline{u^{O4,\lambda}} \delta_{3\lambda} u^{3\lambda} + \frac{9}{8} \overline{v^{O4,\phi}} \delta_\phi u^\lambda - \frac{1}{8} \overline{v^{O4,\phi}} \delta_{3\phi} u^{3\lambda}$$

$$= \left[\frac{9}{8} \delta_\lambda (\overline{u^{O4,\lambda}} \overline{u^\lambda}) - \frac{1}{8} \delta_{3\lambda} (\overline{u^{O4,\lambda}} \overline{u^{3\lambda}}) \right] - u \delta_\lambda^{O4} (\overline{u^{O4}})$$

$$+ \underbrace{\left[\frac{9}{8} \delta_\phi (\overline{v^{O4,\phi}} \overline{u^\lambda}) - \frac{1}{8} \delta_{3\phi} (\overline{v^{O4,\phi}} \overline{u^{3\lambda}}) \right]}_{\equiv \frac{F_{i+n/2} - F_{i-n/2}}{n \delta \phi}} - \overline{v^{O4,\phi}} \delta_\phi^{O4,\lambda} (\overline{u^{O4,\phi}})$$

*: Morinishi, 1998 (JCP), 'Higher order conservative schemes'

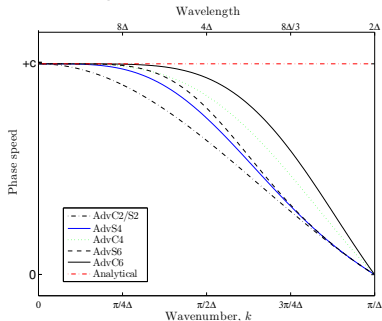
Advantages:

- Higher Order accurate
- Improved model stability due to conservation of the integrated Kinetic energy for non-divergent flows

Motivation: Higher Order Quadratic Conserving Spatial Schemes

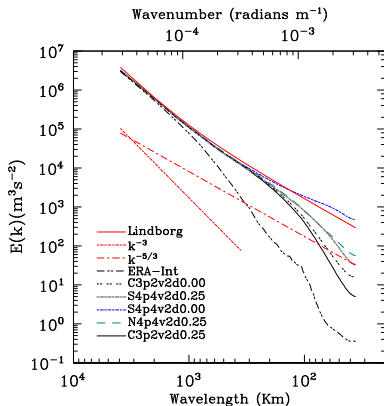
- Conserves 1st and 2nd moments of the advected quantity i.e. kinetic energy

- Reduced phase speed error compared to second order scheme.



- **AdvS4/S4p4v2d0.00: Quadratic conserving scheme**

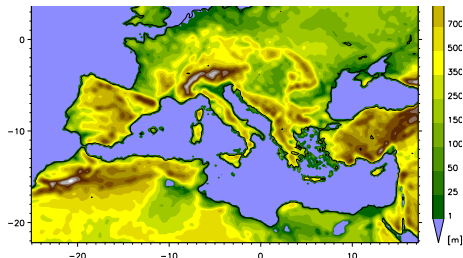
- Effective resolution of the model is increased by atleast a factor of 2



REGIONAL CLIMATE SIMULATION:

Configuration

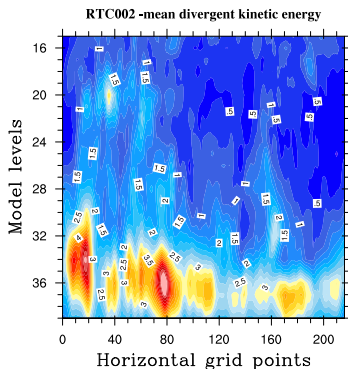
- ▶ COSMO-CLM evaluation domain for EU/Euro CORDEX
- ▶ ERA-interim 1979-1998
- ▶ $\Delta\lambda = 0.165^\circ$
- ▶ $ke_tot(\text{model levels}) = 40$
- ▶ Parameterization schemes
 - Prognostic TKE turbulence closure scheme -level 2.5 (Mellor and Yamada (1982))
 - Tiedtke mass-flux convection scheme (Tiedtke (1989))



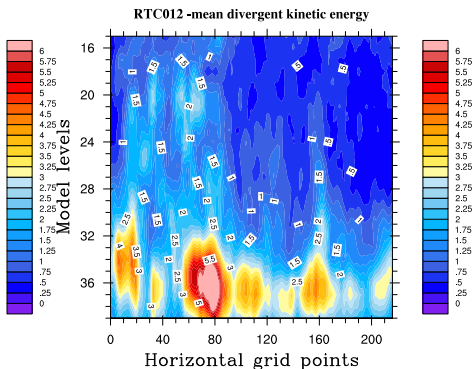
Simulation-ID:	Configuration-ID:	hor.dif. []	dt [s]	simulated time [y]
RTC002	<i>C3p2v2d0.25</i>	0.25	150	79-98
RTC004	<i>N4p4v2d0.25</i>			
RTC016	<i>S4p4v2d0.25</i>			
RTC012	<i>S4p4v2d0.00</i>	0.00		

Summer mean Kinetic Energy from Divergent wind (m^2s^{-2})

C3p2v2d0.25 meridional average

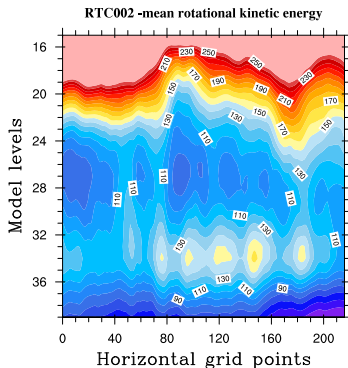


S4p4v2d0.00 meridional average

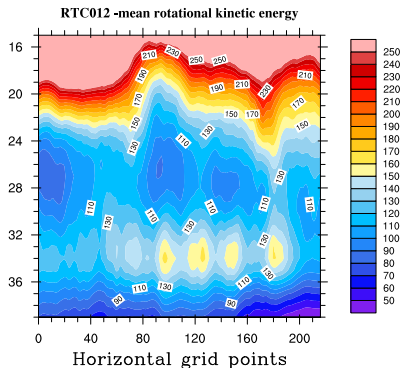


Summer mean Kinetic Energy from Rotational wind (m^2s^{-2})

C3p2v2d0.25 meridional average



S4p4v2d0.00 meridional average

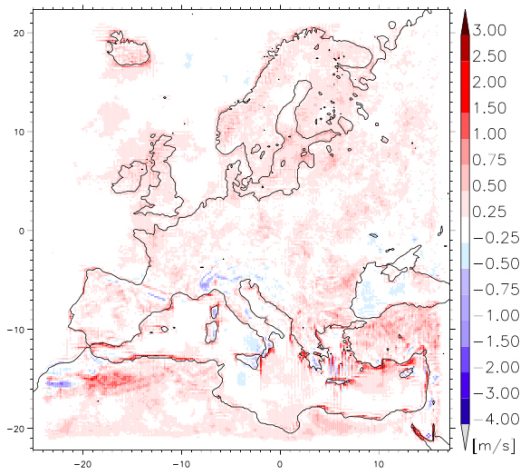


Precipitation, 2m Air temperature and uv_10m maximum

- ▶ Differences between observations and simulations
- ▶ Relative differences between simulations using different spatial schemes

The difference in maximum 10m wind speed (*S4p4v2d0.00-C3p2v2d0.25*)

DIFF: Maximum 10m Windspeed RTC012-RTC002, 1988-198807



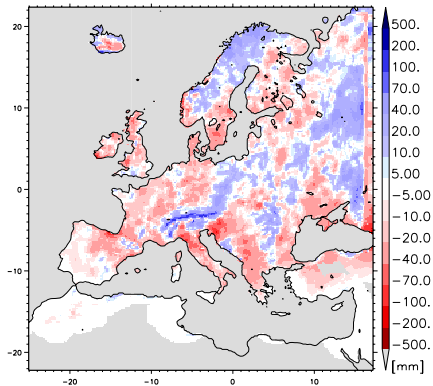
- Significant increase in 10m wind speed

TOT_PREC July 1979-1988 differences

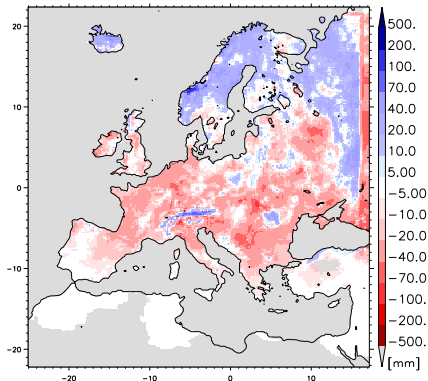
C3p2v2d0.25 – ECAD

S4p4v2d0.00 – ECAD

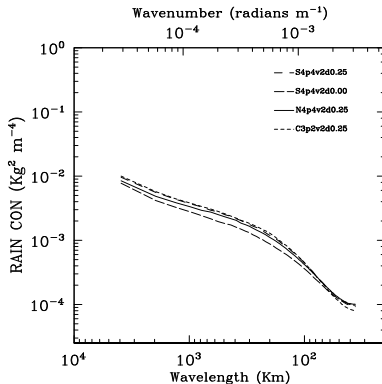
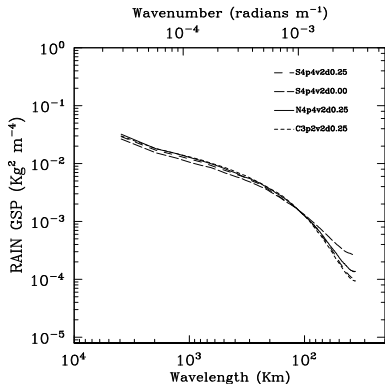
DIFF: Precipitation RTC002–ECAD48, 1983–1983, 08, 00_24



DIFF: Precipitation RTC012–ECAD53, 1988–1988, 07, 00_24



Power spectra of precipitation(SUMMER)



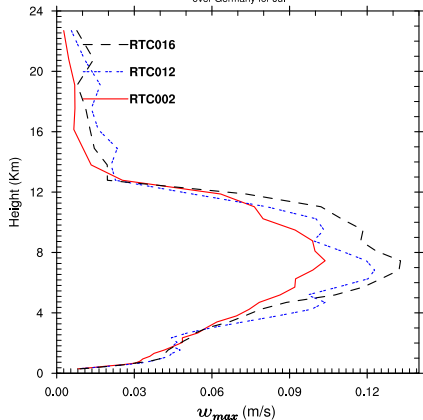
Convective precipitation(RAIN_CON) contributes majorly for the differences between schemes in summer

Summer Maximum updraft (m/s)

summer maximum w , Germany

17-year Maximum vertical velocity profiles

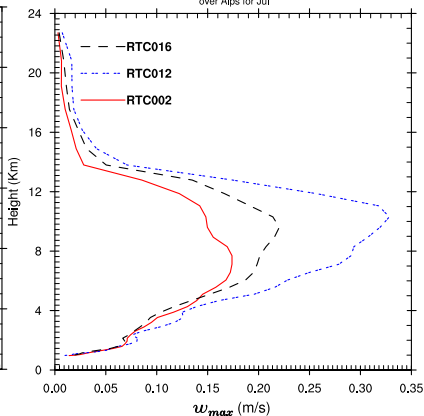
over Germany for Jul



summer maximum w , Alps

17-year Maximum vertical velocity profiles

over Alps for Jul

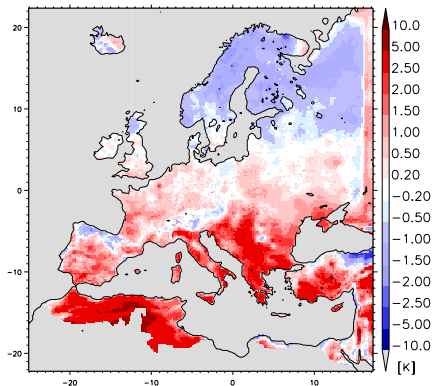


T_2M evaluation: July 1979-1988 differences

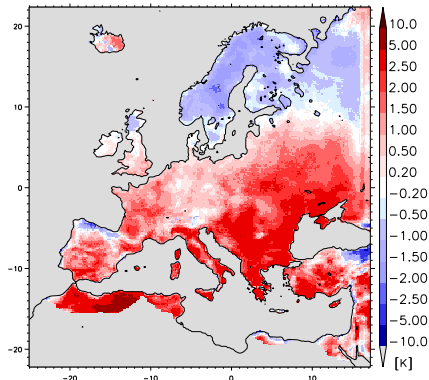
C3p2v2d0.25 – ECAD

S4p4v2d0.00 – ECAD

DIFF: 2m Temperature RTC002–ECAD48, 1983–1983, 08, 00_24



DIFF: 2m Temperature RTC012–ECAD53, 1988–1988, 07, 00_24



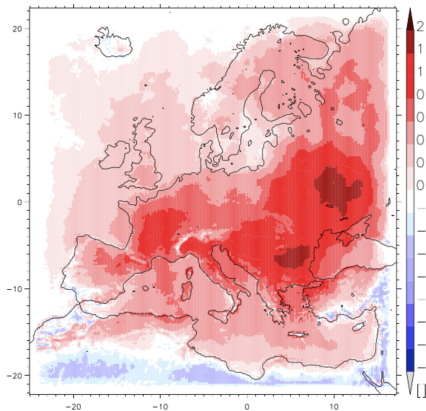
Absolute air temperature: **Relative differences**

S4p4v2d0.00-C3p2v2d0.25 lev:36

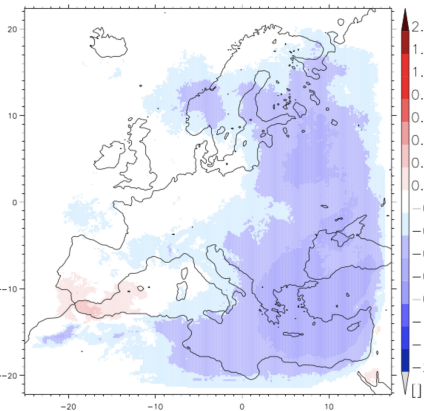
S4p4v2d0.00-C3p2v2d0.25 lev:20

DIFF: 1, RIC012-RIC002, 1998-19980/
LEVEL: 036

DIFF: 1, RIC012-RIC002, 1998-19980/
LEVEL: 020



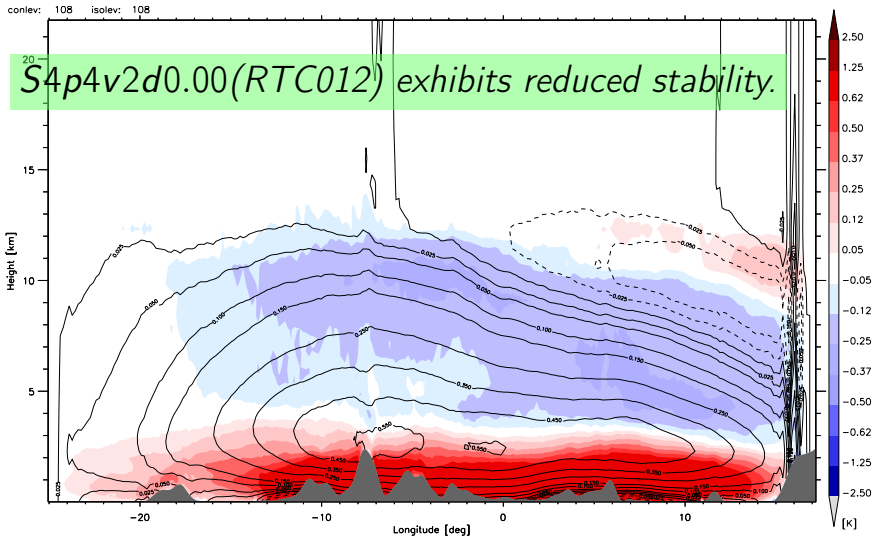
(a) RTC012-RTC002



(a) RTC012-RTC002

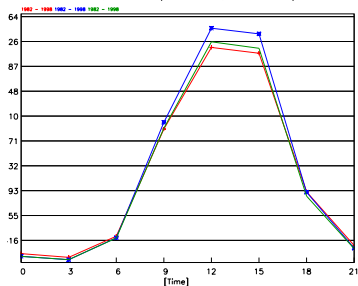
Height-lon cross section: T(Temperature) and PP(Pressure deviation)

Relative differences $S4p4v2d0.00(RTC012)$ - $C3p2v2d0.25(RTC002)$

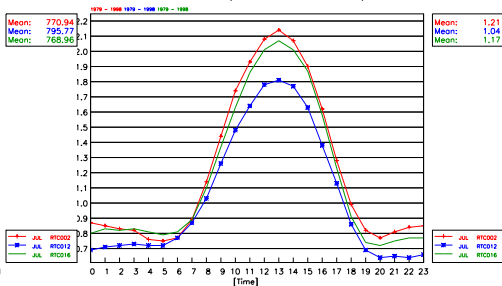


Diurnal cycle for Conv. Precipitation and PBL Height

PBL Height(July), Germany



Convective precipitation(July), Germany



Conclusions

- ▶ Quadratic conserving Fourth order advection scheme `S4p4v2d0.00(RTC012)` is as stable as the third order upwind scheme `C3p2v2d0.25` (with potential for longer time steps) and with higher effective resolution ($\approx 5\Delta\lambda$) compared to the third order upwind scheme ($\approx 8\Delta\lambda$).
- ▶ The increased effective resolution changes the regional climate significantly, particular in summer.
 - ▶ intensifies small scale turbulence, increases the height of the planetary boundary layer and reduces the convective precipitation in summer months.
 - ▶ The reduced precipitation in summer results to reduced soil moisture content and an increase in the summer 2m air temperature(`T_2M`) positive bias.