Modeling the Spatial and Temporal Variability of Fog in the Namib Desert with COSMO

Maike Hacker, Andreas Bott



ICCARUS DWD, Offenbach, 26th February 2018





Outline





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Relevance



Fog precipitation (mm). 7/14 - 6/16. R. Vogt

- Namib is one of the driest deserts on earth
- only a few millimeters precipitation per year
- up to 200 days with fog
- fog water deposition is important water source







NaFoLiCA – Namib Fog Life Cycle Analysis









Model Setup





Model Setup

| Namelist Parameter | COSMO-DE Setup | Tropical Setup (2.8 km) |
|---|----------------|----------------------------|
| domain height | 22 000 | 30 000 |
| number of vertical layers (ke) | 50 | 57 |
| reference temperature on sea level (t0sl) | 288.15 | 300 |
| temperature difference sea level ↔ stratosphere (delta_t) | 75 | 90 |
| scale height (h_scal) | 10 000 | 12 000 |
| coordinate value to change to z-system (vcflat) | 11 357 | 15 000 |
| bottom height of Rayleigh sponge layer (rdheight) | 15 000 | 18 000 |





Model Setup



- COSMO-Version 5.01
- initialization at 12 UTC
- forecast time 30 hours
- twofold nesting
 - COSMO 7 km driven by ICON
 - COSMO 2.8 km driven by COSMO 7 km
- three different vertical grids, Δz_{min} =20m
- 1D TKE-based closure on level 2.5 (Mellor and Yamada, 1982)
- one-moment bulk scheme (Reinhardt and Seifert, 2006)









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Preliminary Results

Case Study 18./19. September 2017





Fog and Stratus Defined by Liquid Water Path

12 UTC





Ke = 57

Ke = 92



Ke = 75



Ke = 75, tkvmin = 0.05





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Fog and Stratus Defined by Liquid Water Path





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Fog and Stratus Defined by Liquid Water Path

06 UTC + 1d





Ke = 57

Ke = 92



Ke = 75



Ke = 75, tkvmin = 0.05





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Fog and Stratus Defined by Liquid Water Path



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Comparison of Model Configurations



- well-mixed layer of 300 to 500 m thickness
- fog layer beneath the inversion
- slight improvement by increased number of vertical layers
- tuning of turbulence scheme is more effective than change of vertical layers



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- unphysical behavior of temperature profiles for decreased minimal turbulent exchange coefficients
- oscillations of turbulent exchange coefficients
 - → numerical instabilities









- filtering based on Buzzi et al. (2011): $K_{k}^{new} = 0.5K_{k} + 0.2 [K_{k+1} + K_{k-1}] + 0.05 [K_{k+2} + K_{k-2}]$
- oscillations of turbulent exchange coefficients are eliminated





Changes in Boundary Layer Dynamics due to Filtering





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Changes in Boundary Layer Dynamics due to Filtering



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Conclusion

Setup of COSMO Simulations

- ✓ tropical setup
- ✓ increase of vertical resolution in the planetary boundary layer

Comparison of Model Configurations

- ✓ initialization from ICON reproduces spatial distribution of stratus and fog
- COSMO strongly underestimates fog and stratus
- × decrease of model layer thickness has no significant influence on results
- ✓ reduction of TKVH_{min} and TKVM_{min} is very effective
- numerical instabilities occur for configurations with decreased mixing and increased number of vertical layers
- x filtering of exchange coefficients changes boundary layer dynamics





Outlook

Treatment of Numerical Instabilities

- objective detection of numerical instabilities
- investigate behavior of dimensionless gradients and stability functions
- improve filtering
- other solutions to instability problem?

Tuning of Model Parameters

- further decrease minimal exchange coefficients
- ..

Dissolution of Fog and Stratus

• sensitivity to initial time

Simulations with COSMO-PAFOG







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Fog and Stratus Defined by Vertically Integrated Liquid Water Content





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Fog and Stratus Defined by Vertically Integrated Liquid Water Content

06 UTC





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Fog and Stratus Defined by Vertically Integrated Liquid Water Content

17 UTC





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Fog and Stratus Defined by Vertically Integrated Liquid Water Content

02 UTC + 1d





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Dissolution of the Stratus Over the Atlantic



- temperature inversion between 300 and 500 m
- saturated layer beneath the inversion
- slight improvement by increased number of vertical layers, but inversion still too weak





Dissolution of the Stratus Over the Atlantic



- moist layer close to inversion layer is still present, but not saturated
- dissolution of the stratus







- unphysical behavior of temperature profiles for changed vertical coordinates
- unexpected oscillations of turbulent exchange coefficients
 - → numerical instabilities



