



## Assimilating Cloud-affected Radiances in Idealized Simulations of Deep Convection

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## I) Motivation

## **II) Numerical Configuration**

- Idealized Simulations
- Variability in Ensemble
- Data Assimilation Experiments

## III) Nature Run

- Synthetic Satellite Fields of Ice Clouds
- Lower Boundary Layer Induced Water Clouds

## **IV) Assimilating Cloud-affected Radiances**

- Brightness Temperature Fields
- Impact on Ice, Water, Wind & Temperature

## V) Conclusions & Outlook





#### I) Motivation: Why do we assimilate Cloud-affected radiances?

- Early coherent and strong signal
- > 6000 measurements every hour in case study
- Satellite measures radiation at suitable spatial (~ km) and temporal (15 min) scales



#### Goal: improve the forecast of

•Precipitation, Temperature, Wind over the whole domain and especially inside convective cells at the convective time scale, for short term weather forecasts





#### **II) Numerical Configuration**

Lange & Craig (2014): The Impact of Data Assimilation Length Scales on Analysis and Prediction of Convective Storms, MWR

**Nature Run** for Observation System Simulation Experiments (OSSE), *5* min output Initial U(z), V(z),  $\theta(z)$ , ... profile from Radiosonde (**Payerne**, Switzerland at 12 UTC, July 20<sup>th</sup> 2007) **Cyclic** boundary conditions with (*n*, *m*, *l*) = (200, 200, 50) The increments *dx* and *dy* are  $\approx$  2 km. The model levels *dz* vary from **100 m** at the surface to **800 m** at domain top. Overall, (*Lx*, *Ly*, *Lz*) corresponds to (394 km, 394 km, 22 km) Timestep *dt* = 6 s, can be increased to 12 s, 24 s **120** members are computationally affordable

new:

**RTTOV** 12, for infrared satellite images **MFASIS** for solar reflectance





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#### II) Variability in Idealized Ensemble nature •Operational local area models comprise large scale variability in initial conditions due to boundary conditions. In idealized setup so far only spatial variability, i.e., position of cells due to white noise in initial conditions, e.g., after 8 h at 14 UTC (right) •We add **temporal variability** by superimposing random large scale perturbations on initial conditions as T'(z), u'(z), v'(z), rh'(z) with standard deviation 0.25 K, 0.25 m/s and 2 % relative humidity variation at a vertical wavelength of approximately 8 km, up to z = 30 km.







nature

#### II) Variability in Idealized Ensemble around Nature Run



#### x / km onset of deep convection varies by more than three hours

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## II) Data Assimilation Experiments (with LETKF in COSMO-KENDA)



Brightness temperature is simulated with an error of 3 K for WV 6.2  $\mu$ m, 7.3  $\mu$ m.



We assimilate every 15 min for 8 cycles, after 8 h lead time with 2 h free forecast. Currently, 40 members run in the cycle. We do not localize in the vertical.

Horizontal localization is  $L_h$ = 32 km.

Vertical localization is  $L_v = \infty$ .

Error model by Harnisch et al. (2016) ranging from  $\approx$  2 K in clear sky to  $\approx$  10 K in cloud regions for WV 6.2 µm

### What is the **potential of assimilating cloud-affected radiances**?



km

x / km



#### **III)** Nature Run: Time Series of Satellite Fields & Radar Reflectivity





Radar Reflectivity (Column Maximum) [dBz]

In the infrared, ice clouds are present at later times.

Deep convection sets as cirrus **anvils** develop.

The ice clouds block the view on lower clouds in the infrared.

Lower clouds cause rain as visible in more intermittent radar reflectivity fields.





#### **III) Nature Run: Clouds in Reflectance of Solar Radiation**



Scheck et al. (2016): A fast radiative transfer method for the simulation of visible satellite imager, Journal of Quantitative Spectrospcopy and Radiative Transfer, 175, 54-67.

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#### **III)** Nature Run: Clouds in Reflectance of Solar Radiation

In observation system simulations high, medium and low clouds occur as: •Ice clouds in the water vapor bands •Water clouds in the infrared window channels & visible •Lower boundary layer induced clouds in the visible reflectances

(all in half hour intervals)





#### **III) Assimilating Cloud-affected Radiances**

1430

1400

Member #1 No data assimilation

**Member #1** Assimilating *Brightness Temperature* of Water Vapor 6.2 µm band

015 Water Vapor 6.2 μm band Brighntess Temperature / K

## **Nature Run**

BT(x, y) for 6.2 µm

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#### **III) Assimilating Cloud-affected Radiances**

LETKF creates conditions for clouds to: •Dissolve, where no clouds exist in nature •Form, where clouds exist in nature run

What is the impact on other variables?





#### IV) Impact on Wind, Temperature, Ice, Water Vapor

17-18 UTC (Second hour of free forecast,  $4 \times 15$  min first guesses): Horizontal and ensemble **mean** of ...







#### IV) Impact on Wind, Temperature, Ice, Water Vapor

14-15 UTC (First hour of data assimilation,  $4 \times 15$  min first guess): Horizontal and ensemble mean absolute **error** of ...







#### **IV) Impact on Wind, Temperature, Ice & Cloud Water**

15-16 UTC (Second hour of data assimilation, 4 x 15 min first guess): Horizontal and ensemble **mean** absolute error of ...







#### IV) Impact on Wind, Temperature, Ice & Radar Ref ectivity

17-18 UTC (Second hour of free forecast, 4 y 15 min first guess): Horizontal and ensemble mean absolute **error** of ...







# V) Conclusions: What is the potential of assimilating brightness temperature?

MIM



- •The assimilated brightness temperature fields improve.
- •Radar reflectivity improves significantly over the height of the ice clouds.
- •Cloud related quantities as cloud ice and water vapor error decrease.
- •Temperature and wind fields improve significantly.
- •Improvement depends on cloud impact, i.e., the improvement is largest at times when and at locations where clouds are present (in OSSE).

#### V) Outlook: What is the most efficient way to assimilate Clouds?

- •Longer Forecasts of up to 5 hours
- •Update Variables: select cloud variables or only wind & temperature
- •Direct assimilation of another instrument, e.g. visible 0.6  $\mu m$
- •Cloud structure, e.g., average cloud cover  $\langle C(x,y) \rangle_{Area}$ , cloud displacement & amplitude score  $DAS(x_0, y_0)$ , smoothness  $\triangle BT(x_0, y_0)$ , fractal dimension  $D_H$
- •**Temporal difference** of brightness temperature  $BT_{diff}(x_0, y_0)$

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## Thank you for your attention!