

# Representation of model error for data assimilation on convective scale

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## Background information:

- Kilometre-scale ENsemble Data Assimilation (**KENDA**) system at DWD since May 2016 (Schraff et al. (2016))
- Data assim. scheme: Local Ensemble Transform Kalman Filter (**LETKF**)
- “Sufficient” background (analysis) spread  $\sigma^b$  ( $\sigma^a$ ) to represent **sampling error** (due to limited size of ensemble) and **model error**
- **Adaptive multi. inflation** (Anderson (2008)):  $\mathbf{P}^b = \frac{1}{N-1} \mathbf{X}^b \mathbf{X}^{bT} \leftarrow \alpha \mathbf{P}^b$
- **Relaxation method:**

1. Relaxation to prior perturbations (**RTPP**, Zhang et al. (2004))

$$\mathbf{X}^a \leftarrow (1 - \alpha_p) \mathbf{X}^a + \alpha_p \mathbf{X}^b \quad \text{operational } \alpha_p = 0.75$$

2. Relaxation to prior spread (**RTPS**, Whitaker and Hamill (2012))

$$\sigma^a \leftarrow (1 - \alpha_s) \sigma^a + \alpha_s \sigma^b \Leftrightarrow \mathbf{X}^a \leftarrow \left( \alpha_s \frac{\sigma^b - \sigma^a}{\sigma^a} + 1 \right) \mathbf{X}^a$$

e.g.,  $\alpha_s = 0.95$  (Bick et al. (2016))

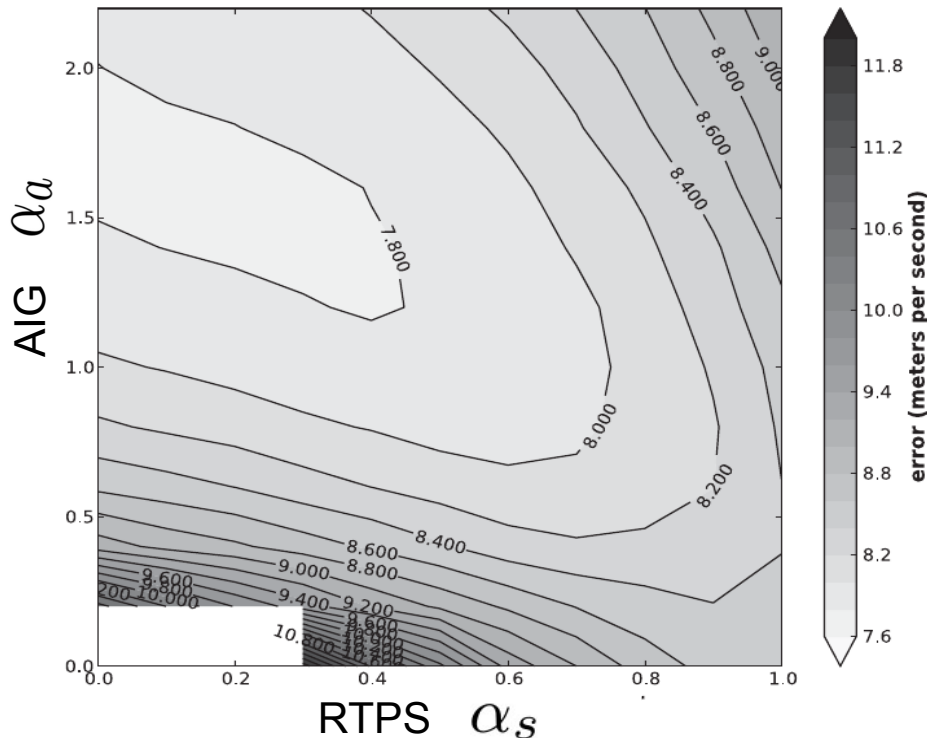
- **Additive inflation:**  $\mathbf{x}^{a(i)} \leftarrow \mathbf{x}^{a(i)} + \alpha_a \boldsymbol{\eta}^{(i)}$

Currently in KENDA: random samples of climatological background error covariances from global EnVar data assimilation system for ICON.

We call it “**large-scale**” additive inflation, denoted by “**AIG**”,  
operational  $\alpha_a = 0.1$

# Motivation

Whitaker and Hamill (2012) compare AIG with RTPS in combination using two-level primitive equation global model. Ensemble size is 200, so sampling error is very small



**Fig.:** Contours of the ensemble mean background error using a combination of AIG and RTPS

“ when model error is the dominant source of unrepresented background errors, additive inflation alone outperforms any combination of RTPS and additive inflation.”

**Model error is prevailing at convective-scale**

**Question: AIG, RTPP/RTPS or else?**

# Outline

1. Comparison of AIG, RTPP and combination
2. Comparison of AIG, RTPS and combination
3. Introduction of additive inflation based on model truncation error
4. Conclusion and outlook

## Experimental design:

**Period:** 00:00 UTC 27 May 2016 – 00:00 UTC 03 June 2016

**Weather situation:** atmospheric blocking, stationary thunderstorms

**Observations:** conventional data (AIREP, TEMP, PILOT, SYNOP) +  
radar reflectivity

**Data assim. scheme:** LETKF (also for radar reflectivity, using forward operator  
EMVORADO (Zeng et al. (2016)))

**Assimilation window:** one hour

**Size of ensemble:** 40 members, and 20 members are used for 6-h ensemble  
forecasts, initiated at 10, 11, ..., 18:00 UTC

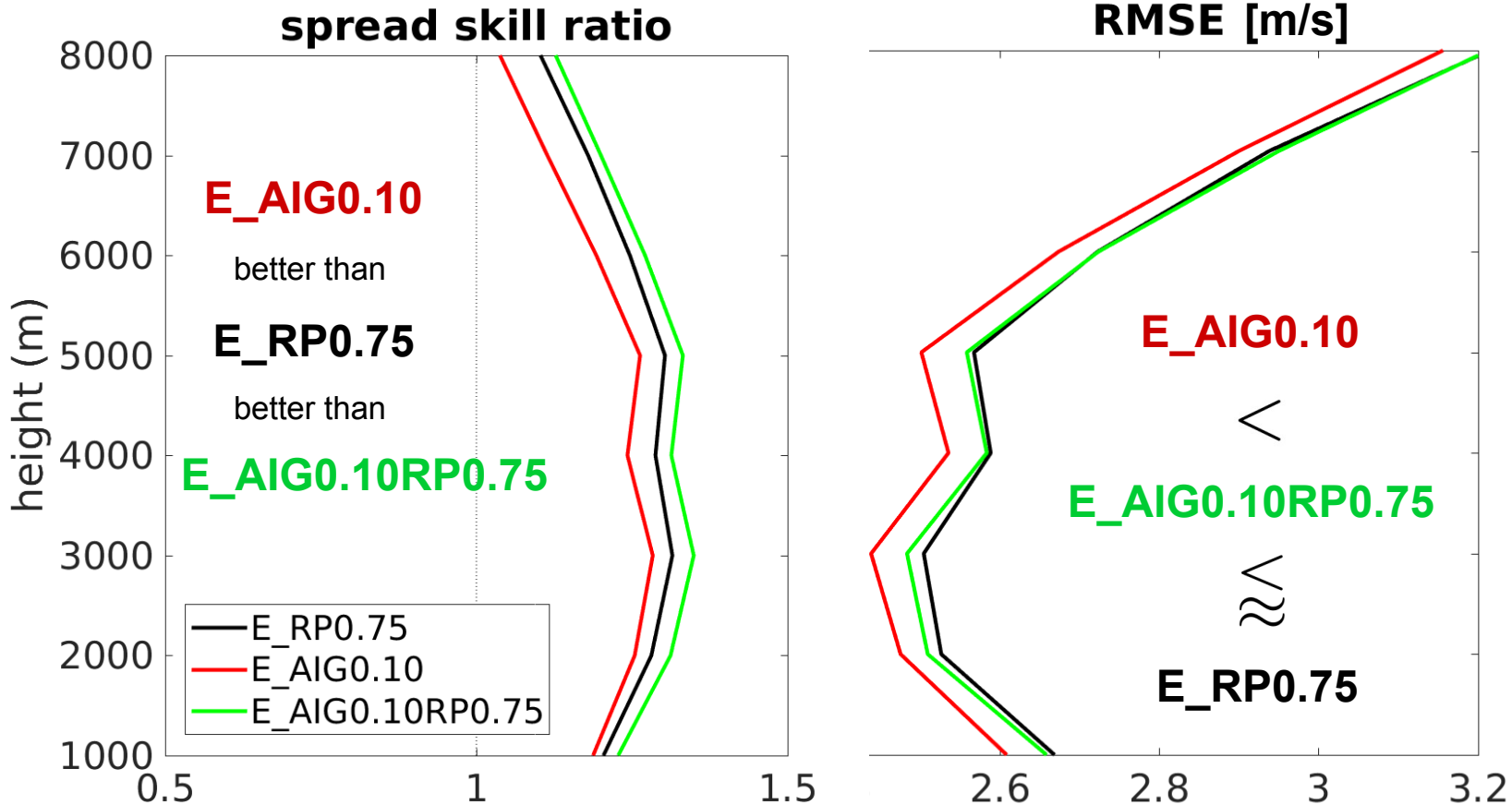
**Localization:** adaptive horizontal localization for conventional data, constant  
horizontal localization (16 km) for reflectivity

**Observation error:** adaptive for conventional data, constant (10 dBZ) for reflectivity

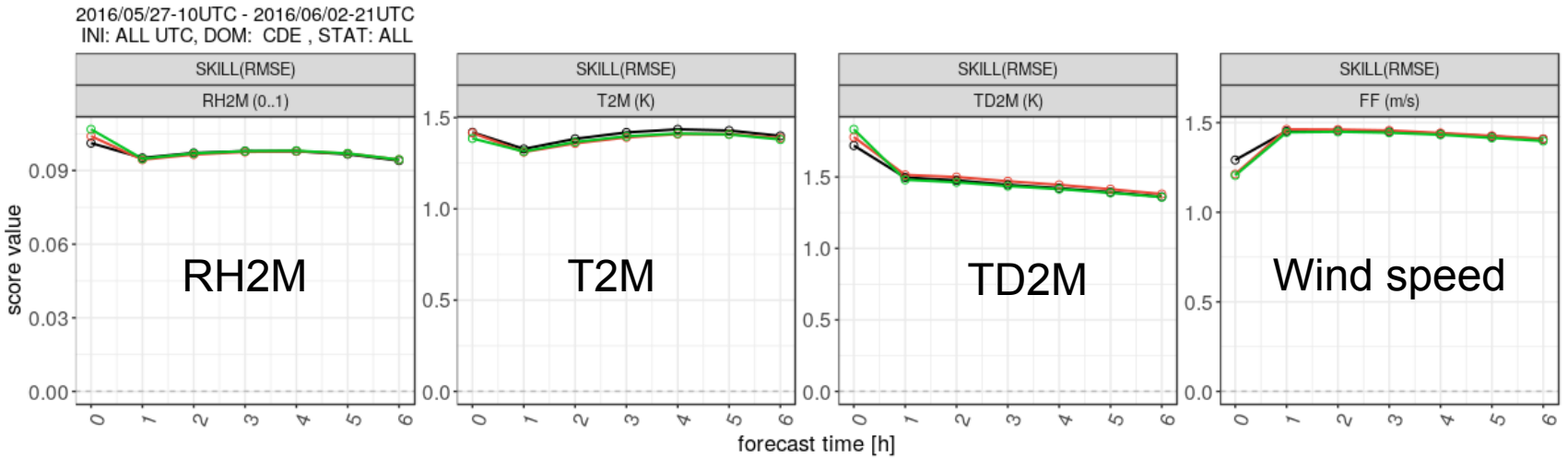
# Study I: Comparison of AIG and RTPP (spread skill ratio & RMSE)

**E\_RP0.75** : RTPP ( $\alpha_p = 0.75$ ) only; **E\_AIG0.10**: AIG ( $\alpha_a = 0.1$ ) only

**E\_AIG0.10RP0.75**: AIG ( $\alpha_a = 0.1$ ) + RTPP ( $\alpha_p = 0.75$ )



# Study I: Comparison of AIG and RTPP (RMSE of ensemble forecast)

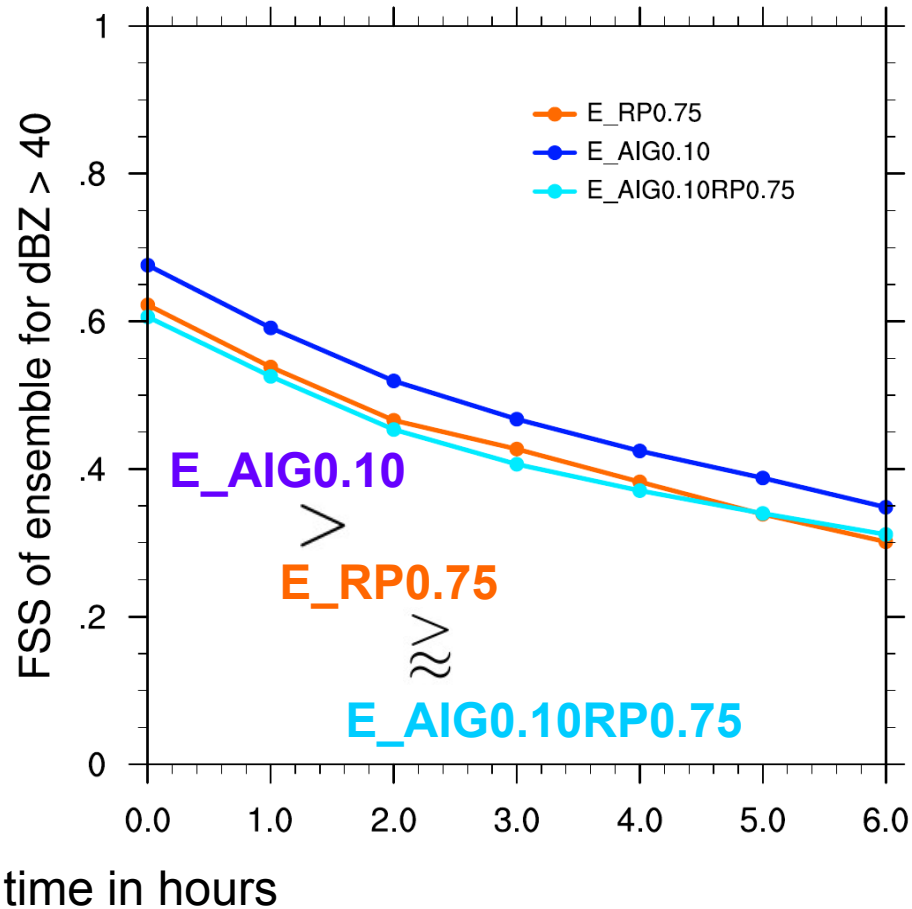
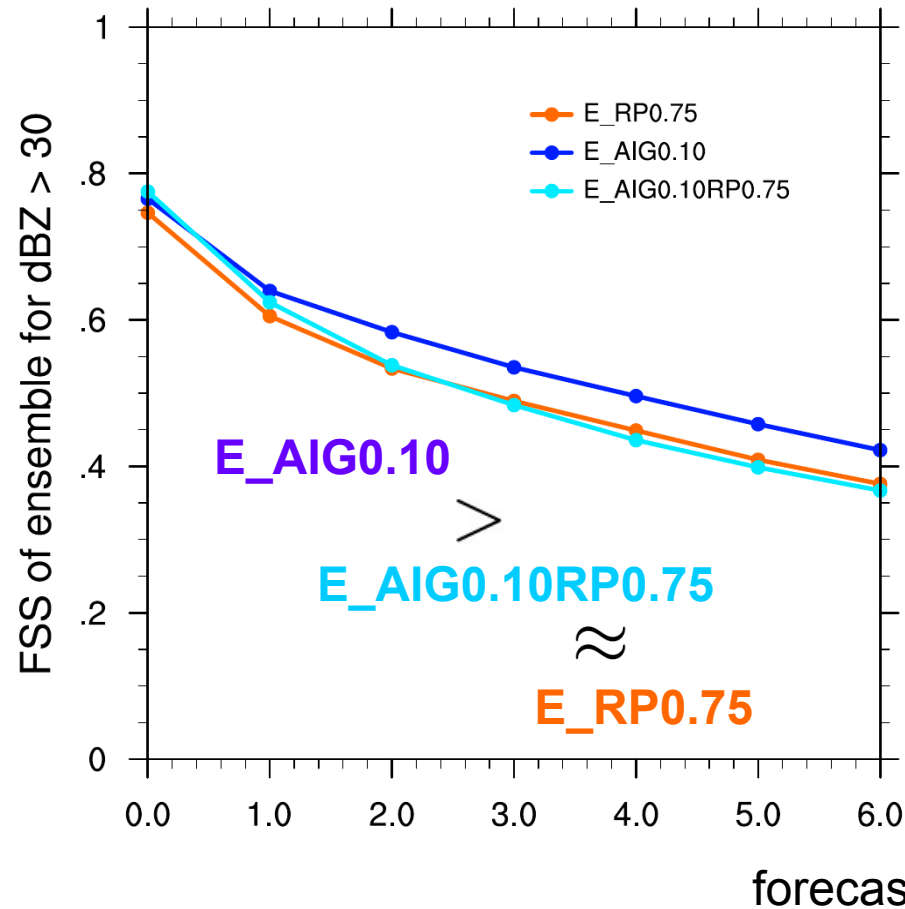


Verification of 6-h ensemble forecast against SYNOP

$$E_{RP0.75} \approx E_{AIG0.10} \approx E_{AIG0.10RP0.75}$$

# Study I: Comparison of AIG and RTPP (Fraction skill score (FSS) of reflectivity in ensemble forecast)

FSS with scale of 30 km for different thresholds 30 and 40 dBZ: the higher, the better





# Study I: Comparison of AIG and RTPP (reflectivity composite in initial time & forecast)

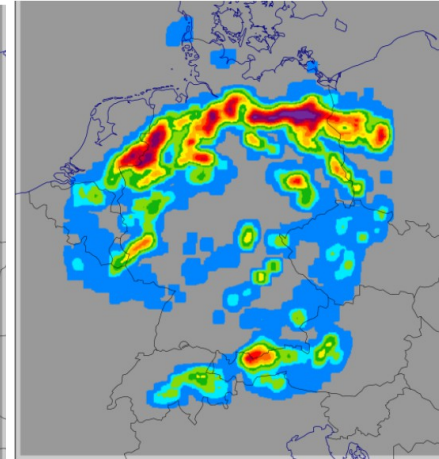
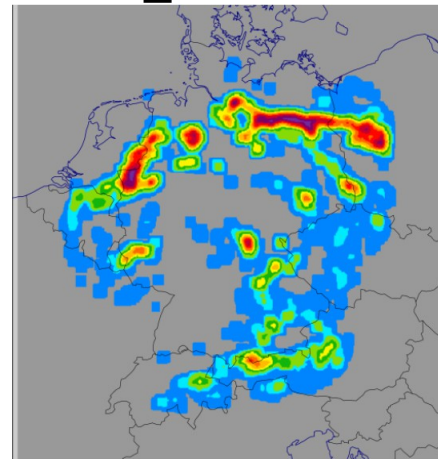
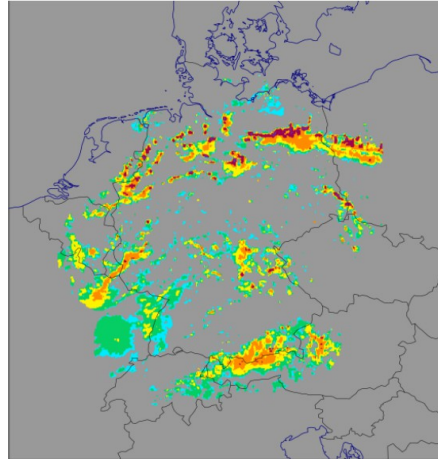
14:00 30 May, 2016

Obs

E\_AIG0.10

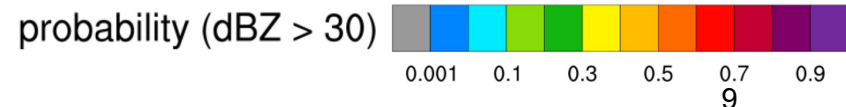
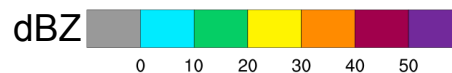
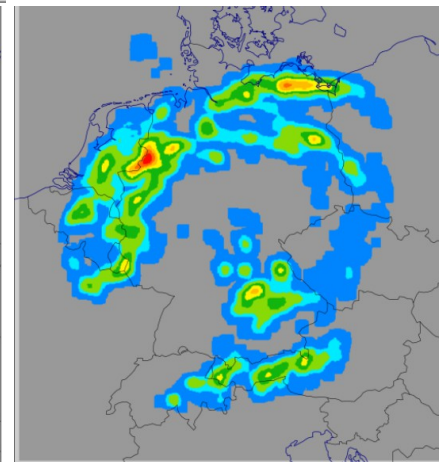
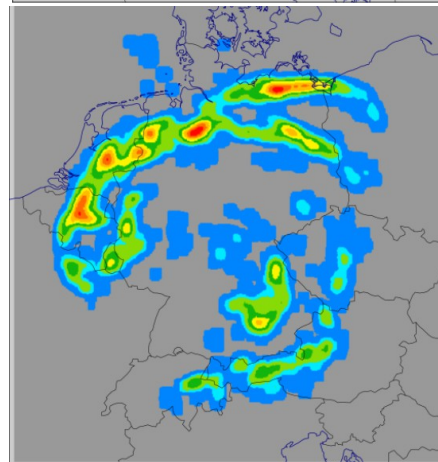
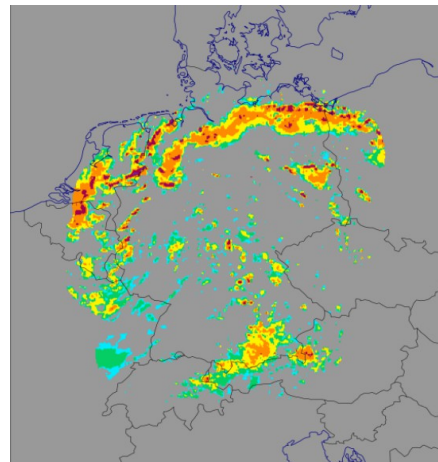
E\_RP0.75

1. Column:  
Reflectivity  
composite

Initial  
time

2.&3. Columns:  
How much  
percent of  
ensemble  
members  
exceed  
30 dBZ

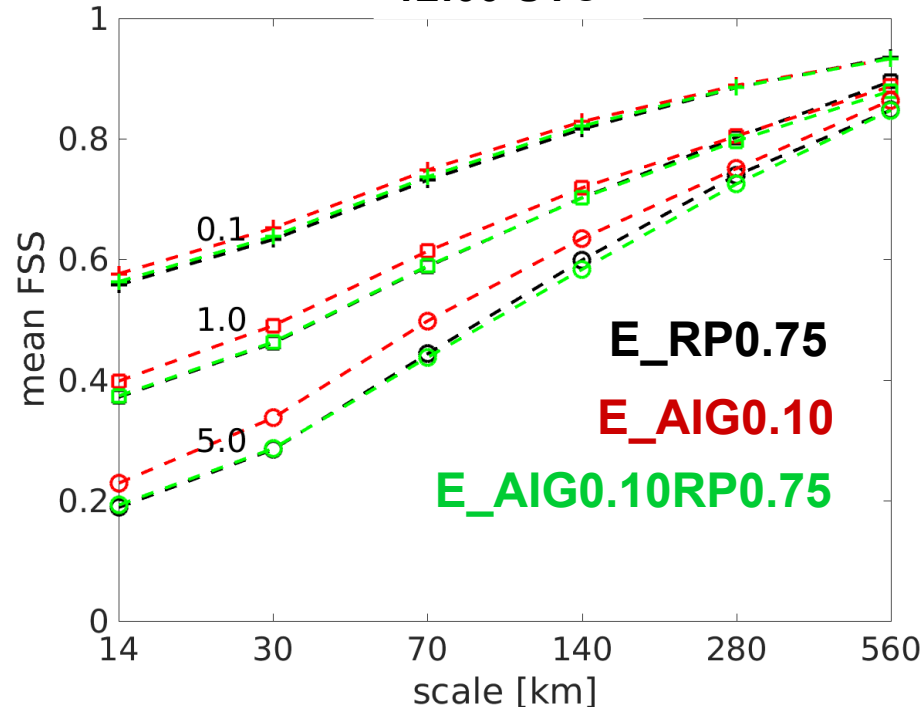
3 h



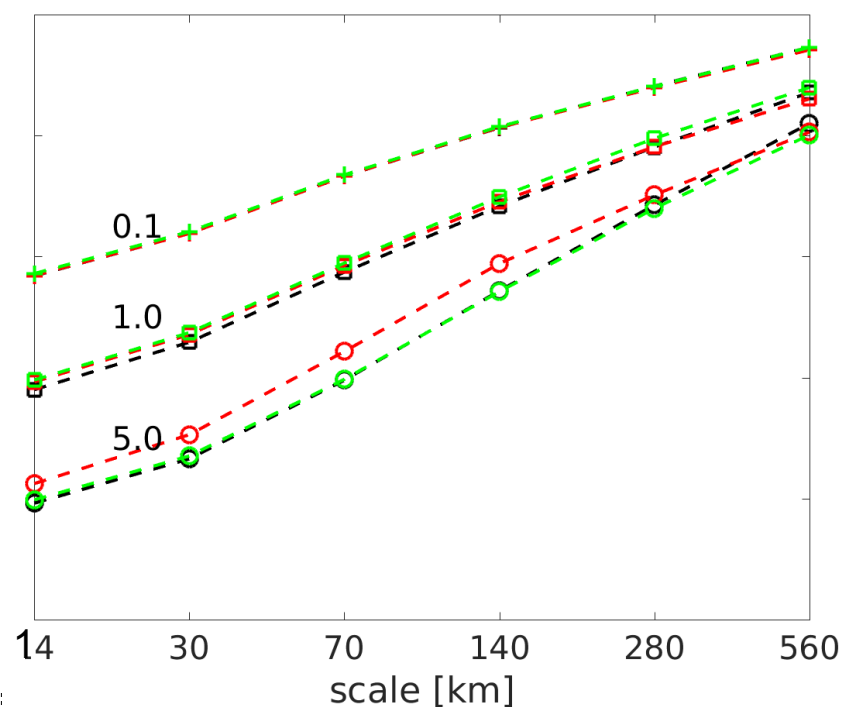
# Study I: Comparison of AIG and RTPP (Fraction skill score of precipitation forecast)

FSS for different precip. rate thresholds 0.1, 1.0 & 5.0 mm/h and scales 14,..., 560 km

12:00 UTC



16:00 UTC



0.1 mm/h: **E\_AIG0.10**  $\gg$  **E\_AIG0.10RP0.75**  $\approx$  **E\_RP0.75**

1.0 mm/h: **E\_AIG0.10**  $>$  **E\_AIG0.10RP0.75**  $\approx$  **E\_RP0.75**

5.0 mm/h: **E\_AIG0.10**  $>$  **E\_AIG0.10RP0.75**  $\approx$  **E\_RP0.75**

0.1 mm/h: **E\_AIG0.10**  $\approx$  **E\_AIG0.10RP0.75**  $\approx$  **E\_RP0.75**

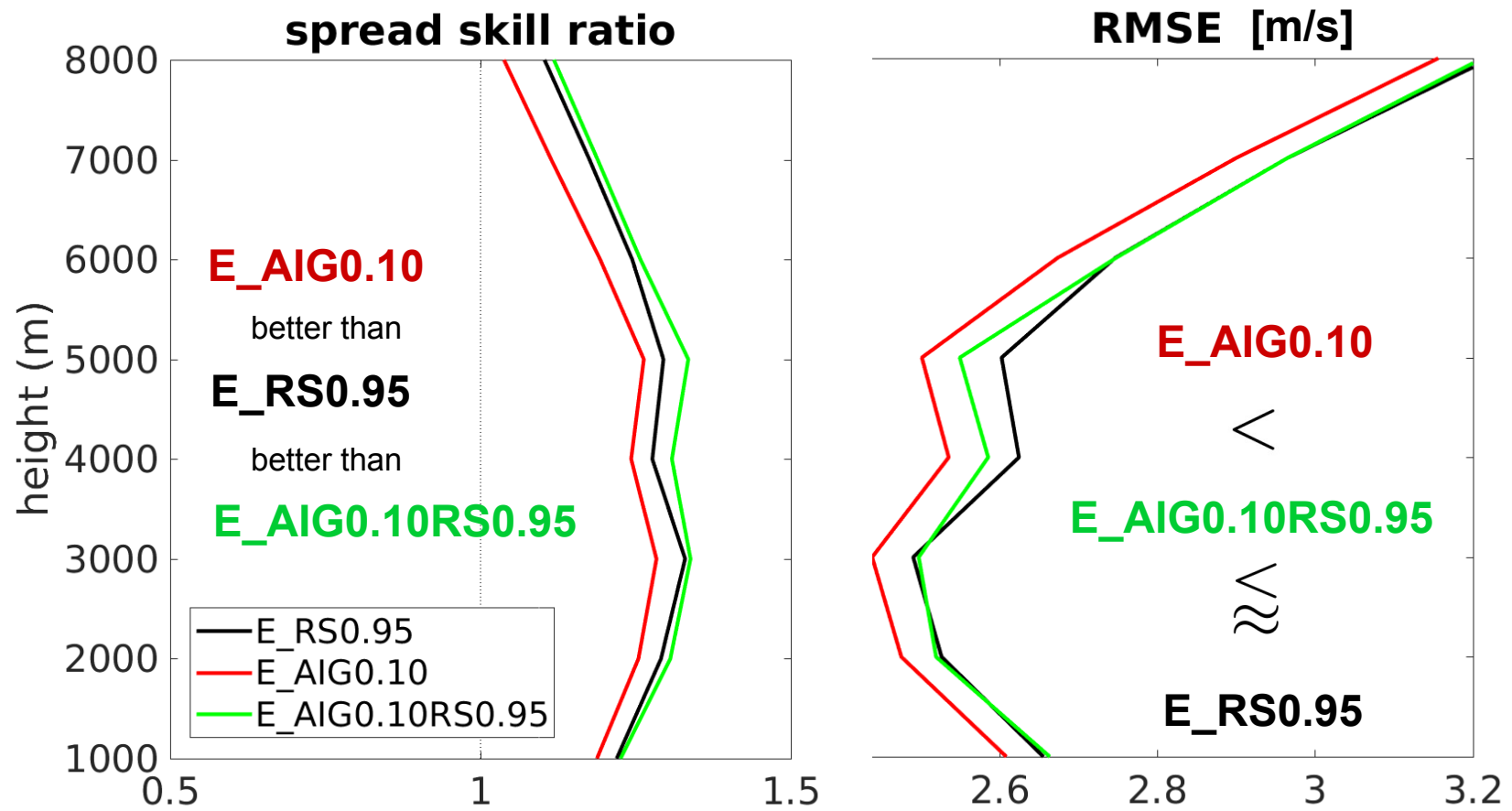
1.0 mm/h: **E\_AIG0.10**  $\approx$  **E\_AIG0.10RP0.75**  $\approx$  **E\_RP0.75**

5.0 mm/h: **E\_AIG0.10**  $>$  **E\_AIG0.10RP0.75**  $\approx$  **E\_RP0.75**

# Study II: Comparison of AIG and RTPS (spread skill ratio & RMSE)

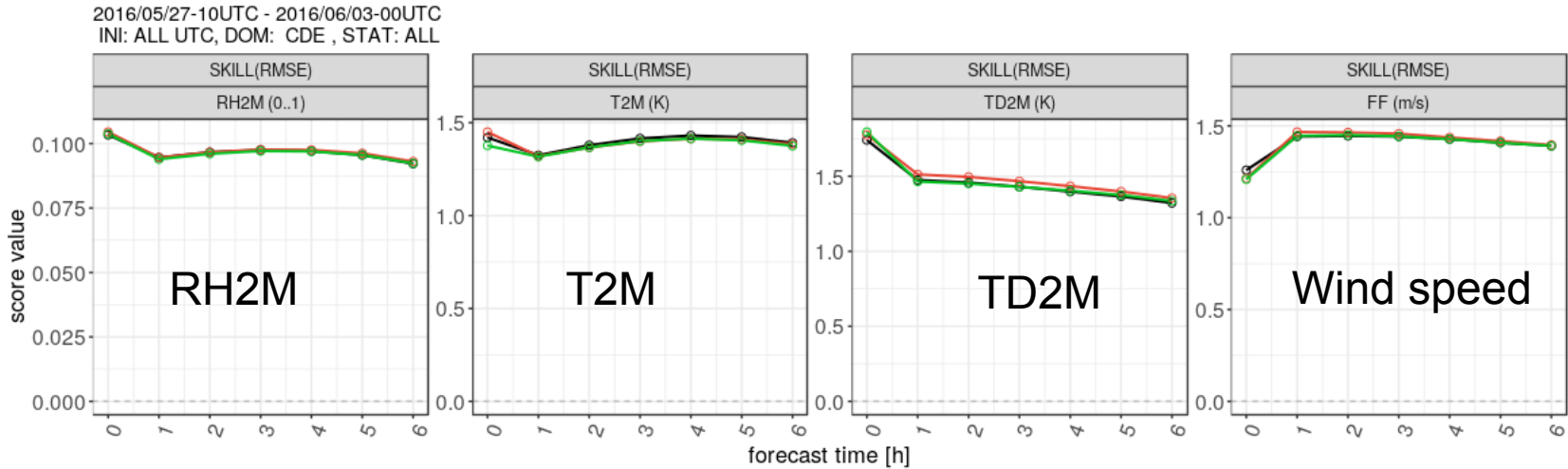
**E\_RS0.95** : RTPS ( $\alpha_s = 0.95$ ) only; **E\_AIG0.10**: AIG ( $\alpha_a = 0.1$ ) only

**E\_AIG0.10RS0.95**: AIG ( $\alpha_s = 0.1$ ) + RTPS ( $\alpha_a = 0.95$ )



Verification of first guess ensemble against **Radial Wind** within assim. cycles

# Study II: Comparison of AIG and RTPS (RMSE of ensemble forecast)

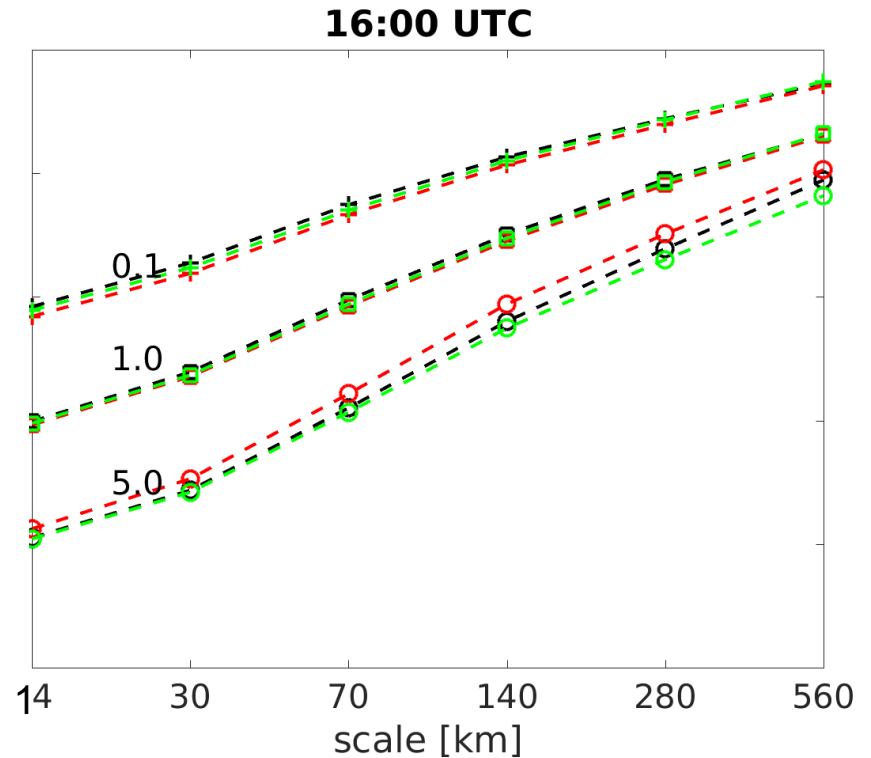
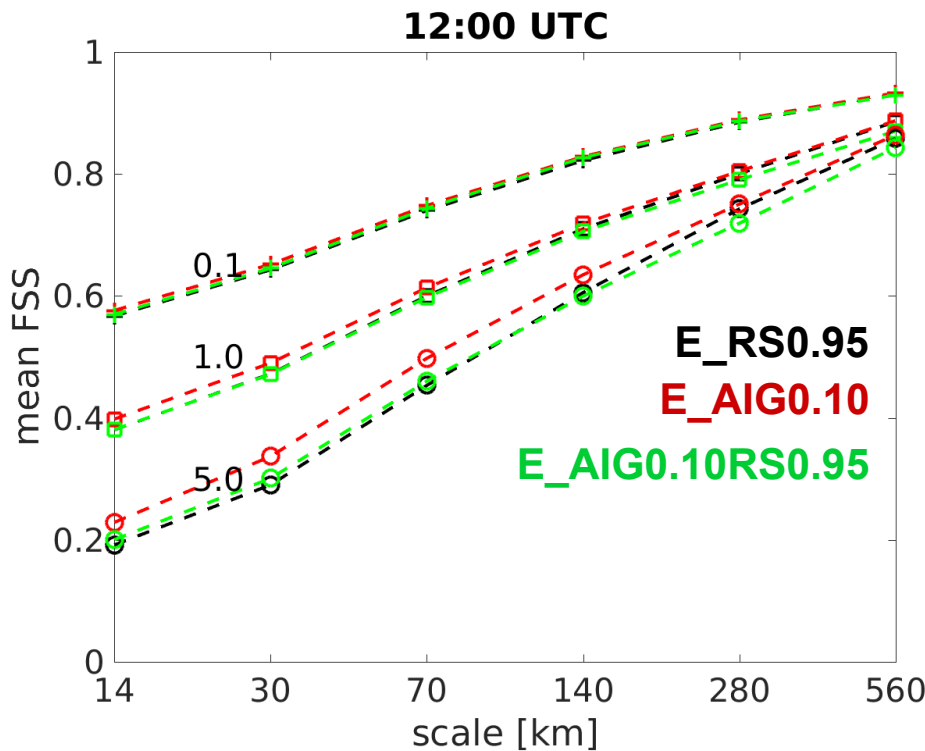


Verification of 6-h ensemble forecast against SYNOP

$$E_{RS0.95} \approx E_{AIG0.10} \approx E_{AIG0.10RS0.95}$$

# Study II: Comparison of AIG and RTPS (Fraction skill score of precipitation forecast)

FSS for different precip. rate thresholds 0.1, 1.0 & 5.0 mm/h and scales 14,..., 560 km



0.1 mm/h: **E\_AIG0.10**  $\approx$  **E\_AIG0.10RS0.95**  $\approx$  **E\_RS0.95**

**E\_AIG0.10**  $\approx$  **E\_AIG0.10RS0.95**  $\approx$  **E\_RS0.95**

1.0 mm/h: **E\_AIG0.10**  $\approx$  **E\_AIG0.10RS0.95**  $\approx$  **E\_RS0.95**

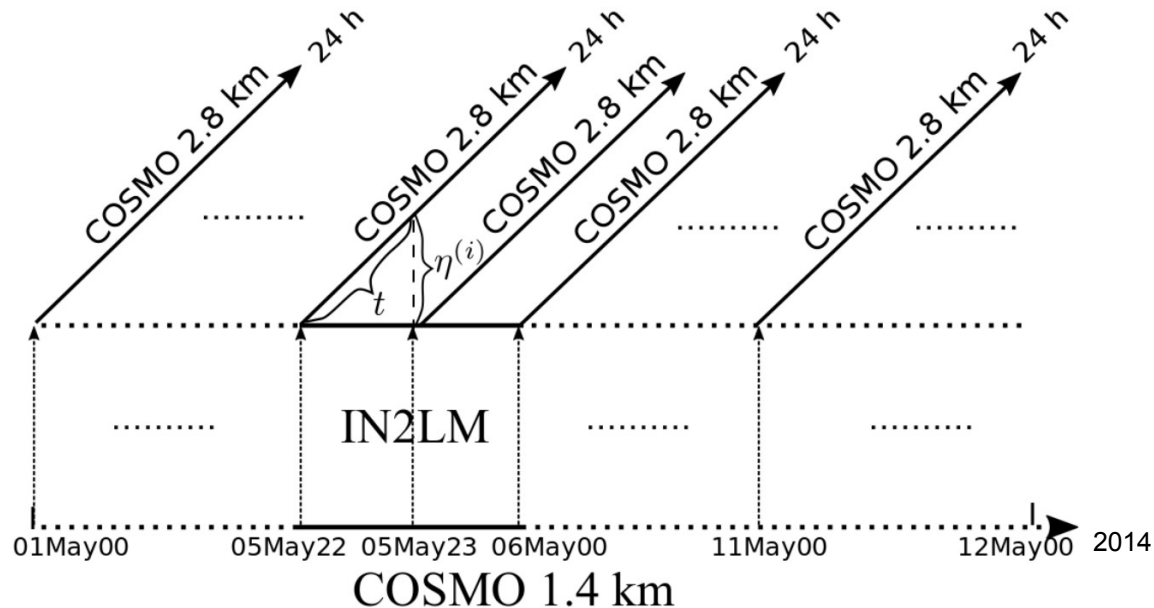
**E\_AIG0.10**  $\approx$  **E\_AIG0.10RS0.95**  $\approx$  **E\_RS0.95**

5.0 mm/h: **E\_AIG0.10**  $>$  **E\_AIG0.10RS0.95**  $\approx$  **E\_RS0.95**

**E\_AIG0.10**  $>$  **E\_AIG0.10RS0.95**  $\approx$  **E\_RS0.95**

# Introduction of additive inflation based on model truncation error

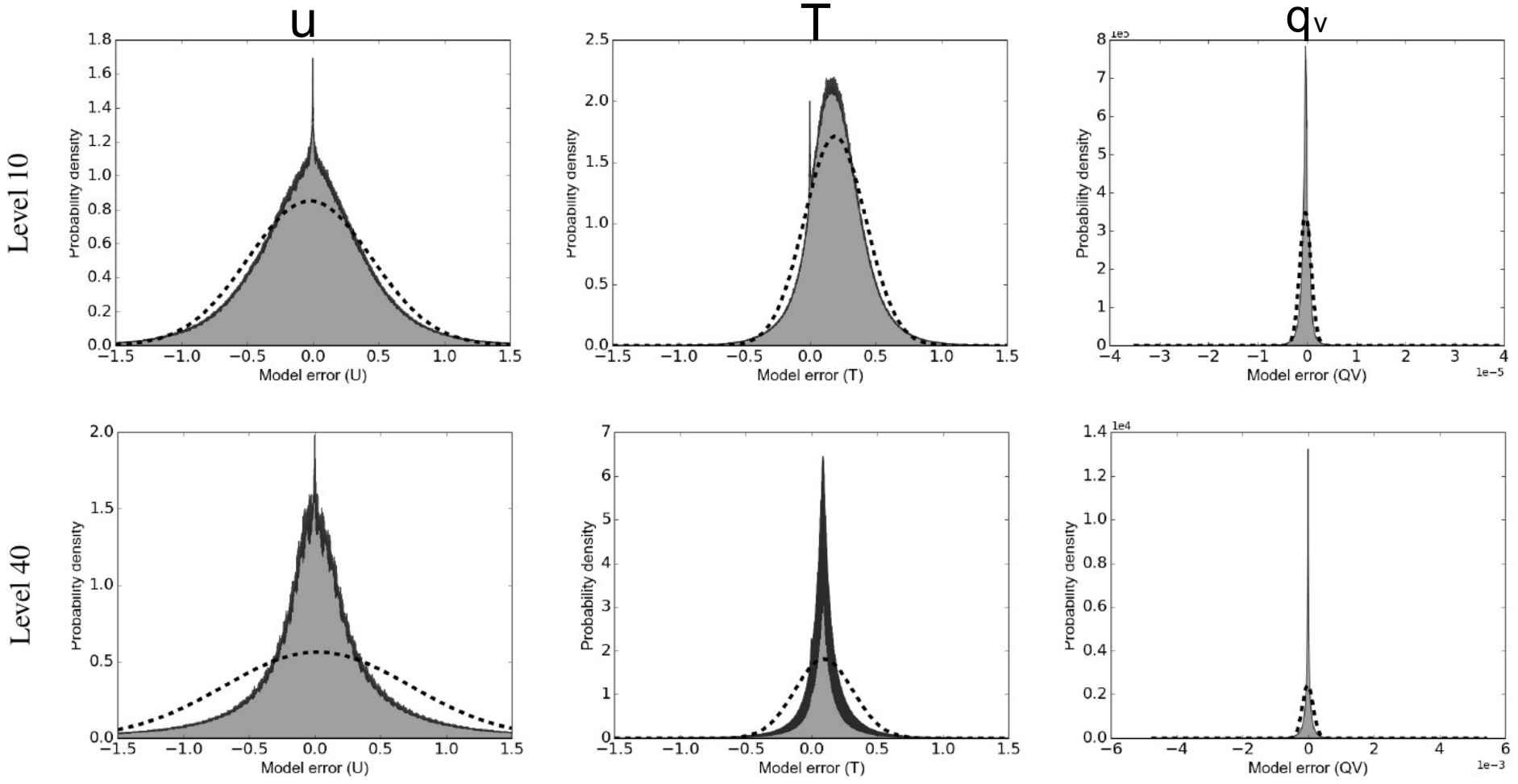
- Model truncation error is one of important sources of model error
- The refinement of the horizontal resolution improves the convective-scale precip. forecasts (e.g., Clark et al. (2016))
- Creation of sample archive for model truncation error



- Approach: choose  $t = 1$  hour,  $\mathbf{x}^{a(i)} \leftarrow \mathbf{x}^{a(i)} + \alpha_b \eta^{(i)}$

$\eta^{(i)}$  samples represent **unresolved/small-scale** model error  
We call it “small-scale” additive inflation, denoted by “ **AIS** ”

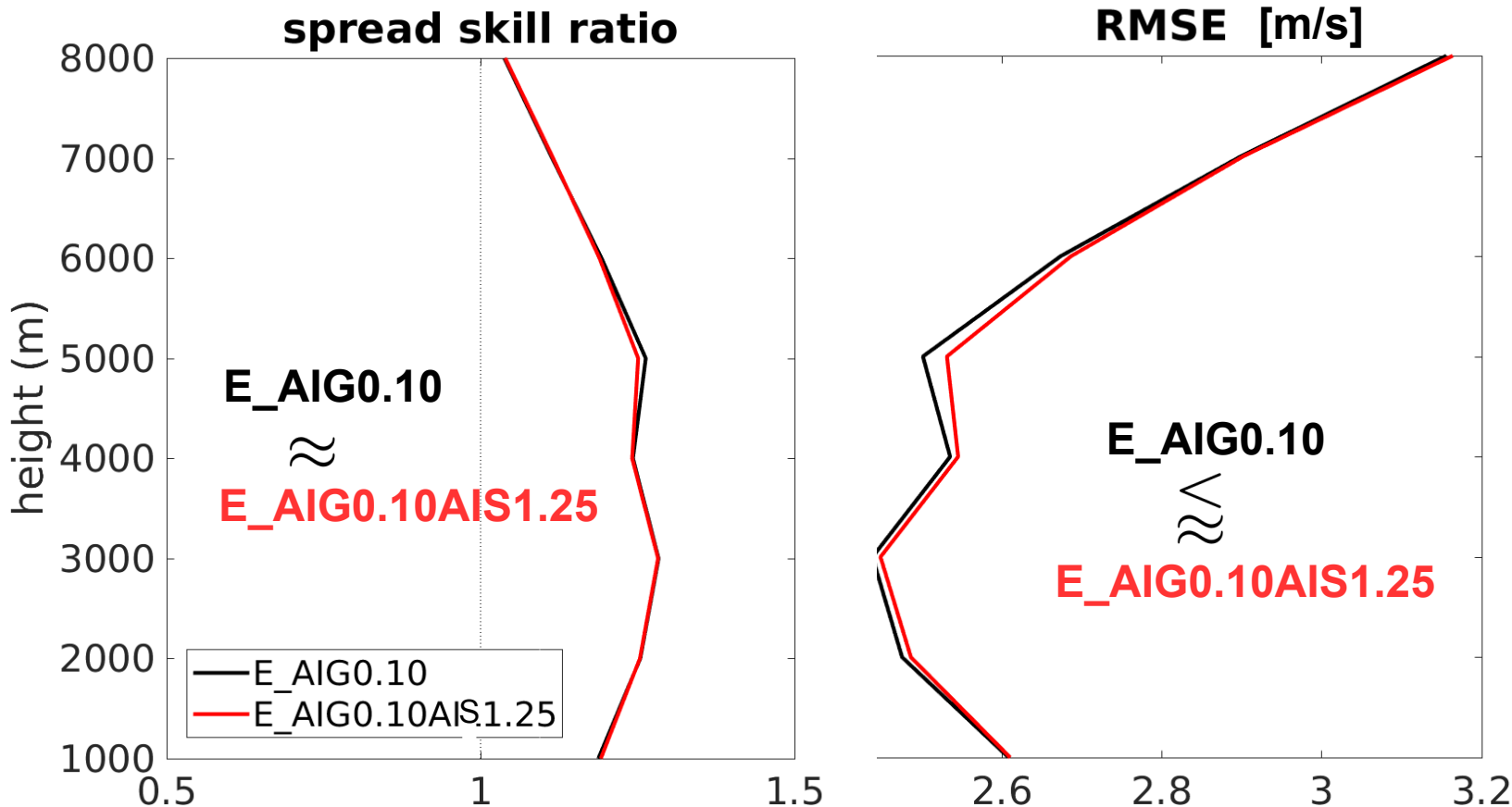
# Introduction of additive inflation based on model truncation error (Histogram of model error samples)



# Study III: Comparison of AIG and AIG+AIS(spread skill ratio & RMSE)

**E\_AIG0.10**: AIG ( $\alpha_a = 0.1$ ) only

**E\_AIG0.10AIS1.25**: AIG ( $\alpha_a = 0.1$ ) +  
AIL ( $\alpha_b = 1.25$ ) with  $u, v, T, q_v$  perturbed

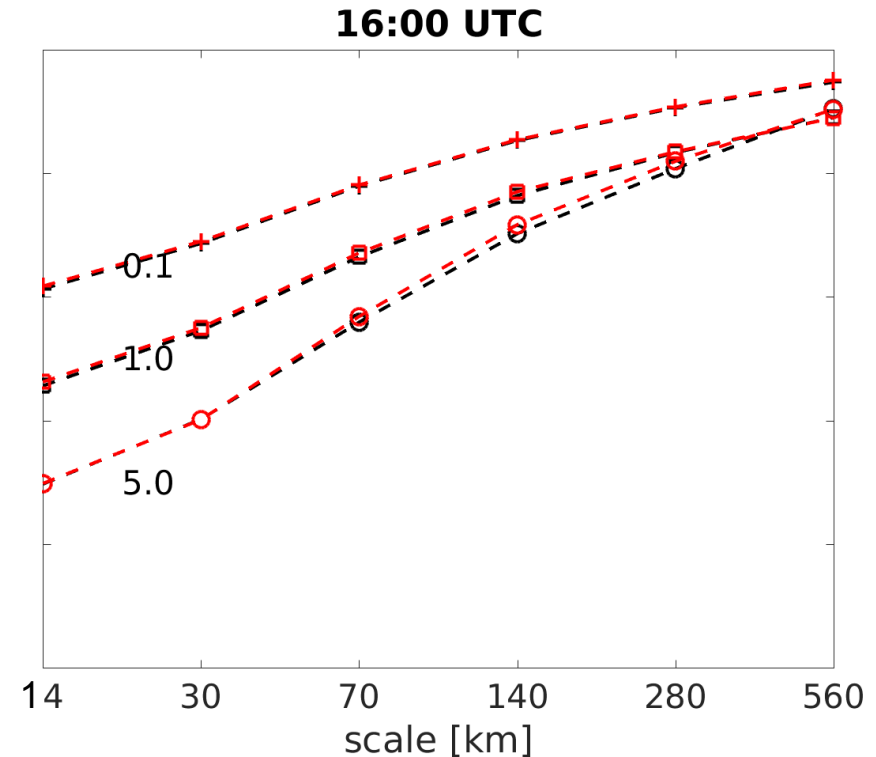
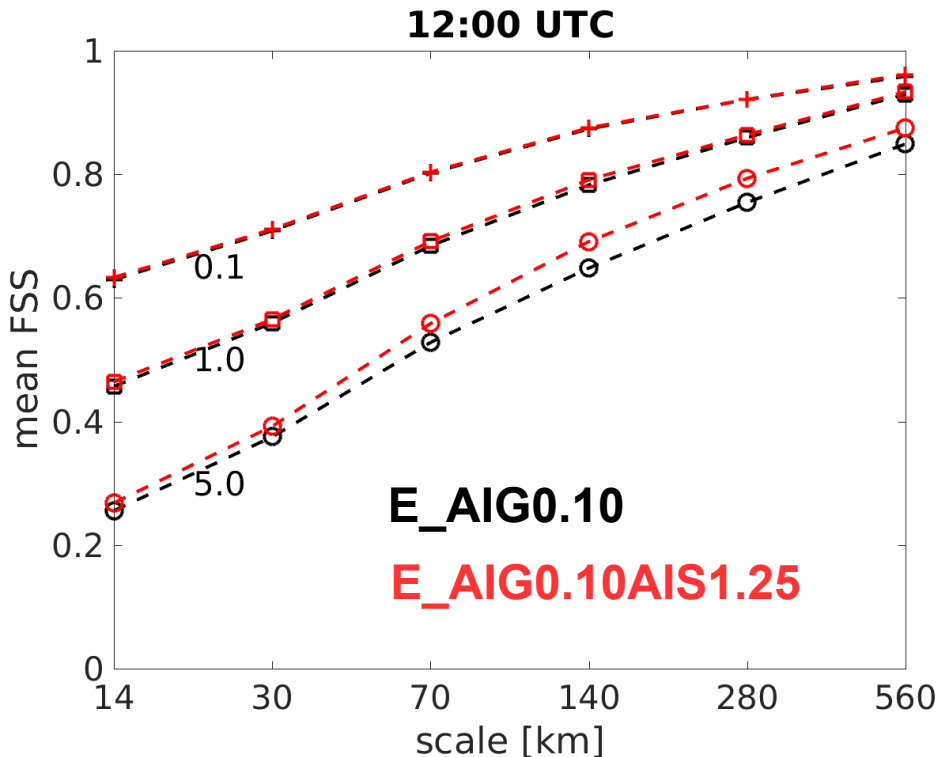


Verification of first guess ensemble against **Radial Wind** within assim. cycles



# Study III: Comparison of AIG and AIG+AIS (Fraction skill score of precipitation forecast)

FSS for different precip. rate thresholds 0.1, 1.0 & 5.0 mm/h and scales 14,..., 560 km



0.1 mm/h: **E\_AIG0.10AIS1.25** ≈ E\_AIG0.10

**E\_AIG0.10AIS1.25** ≈ E\_AIG0.10

1.0 mm/h: **E\_AIG0.10AIS1.25** ≈ E\_AIG0.10

**E\_AIG0.10AIS1.25** ≈ E\_AIG0.10

5.0 mm/h: **E\_AIG0.10AIS1.25** > E\_AIG0.10

**E\_AIG0.10AIS1.25** >> E\_AIG0.10

# Conclusion and Outlook

## Conclusion:

1. Large-scale additive inflation alone outperforms RTPP, RTPS and combination both in cycling and short-term precip. forecast for convective-scale data assimilation
2. Small-scale additive inflation based on model truncation error further improves large-scale additive inflation for short-term precip. forecast

## Outlook:

1. To tune small-scale additive inflation
2. To compare small-scale additive inflation with warm bubbles and stochastic boundary layer perturbations
3. Papers in preparation:

Y. Zeng, T. Janjic, A. de Lozar, U. Blahak, M. Sommer, H. Reich, A. Seifert, 2018: Representation of model error for data assimilation on convective scale. Part I: Additive noise based on model truncation errors.

Y. Zeng, T. Janjic, A. de Lozar, U. Blahak, A. Seifert, S. Rasp, G. C. Craig, 2018: Representation of model error for data assimilation on convective scale. Part II: Comparison of additive noise and differently specified boundary layer uncertainties.

# Reference

Bick, T., C. Simmer, S. Trömel, K. Wapler, K. Stephan, U. Blahak, Y. Zeng, and R. Potthast, 2016: Assimilation of 3d-radar reflectivities with an ensemble kalman filter on the convective scale. 142, 1490–1504.

Clark, P., N. Roberts, H. Lean, S. P. Ballard, and C. Charlton-Perez, 2016: Convection-permitting models: a step-change in rainfall forecasting. *Meteorological Applications*, 23 (2), 165–181.

Schraff, C., H. Reich, A. Rhodin, A. Schomburg, K. Stephan, A. Perianez, and R. Potthast, 2016: Kilometre-scale ensemble data assimilation for the Cosmo model (KENDA). *Quart. J. Roy. Meteor. Soc.*, 142, 1453–1472.

Whitaker, J. S. and T. M. Hamill, 2012: Evaluating methods to account for system errors in ensemble data assimilation. *Mon. Wea. Rev.*, 140(9), 3078–3089.

Zeng, Y., U. Blahak and D. Jerger, 2016: An efficient modular volume-scanning radar forward operator for NWP models: description and coupling to the COSMO model. *Quart. J. Roy. Meteor. Soc.*, 142, 3234–3256.

Zhang, F., C. Snyder, and J. Sun, 2004: Impacts of initial estimate and observation availability on convective-scale data assimilation with an ensemble Kalman filter. *Mon. Wea. Rev.*, 132(5), 1238–1253.

## Thank you for your attention