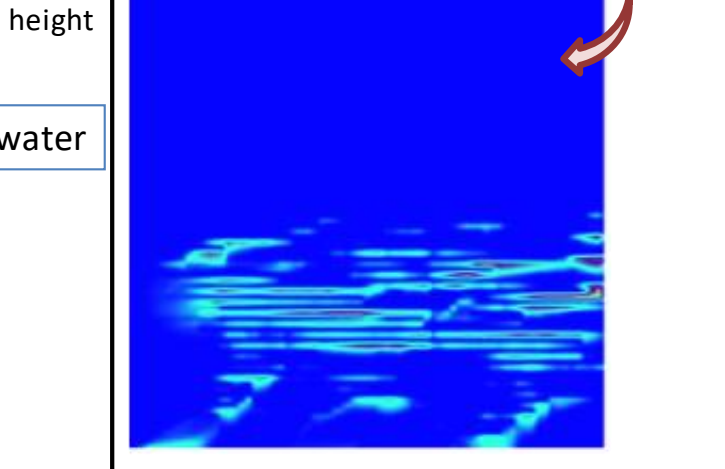
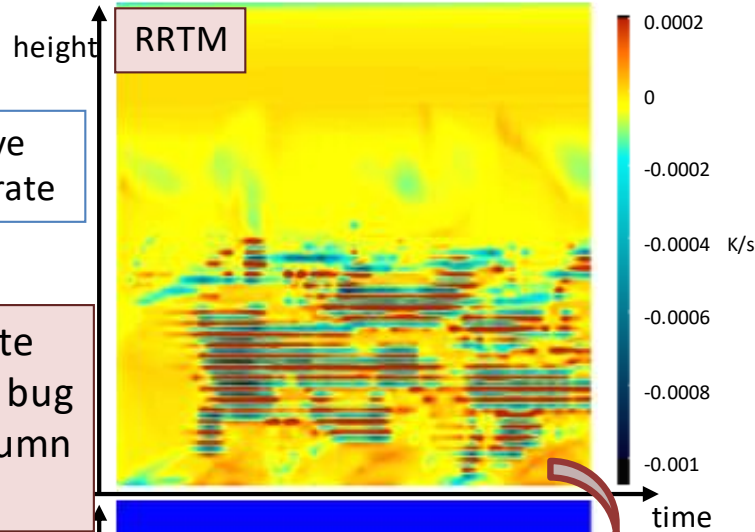


New radiation scheme ecRad in ICON + uncertainties in radiation modelling

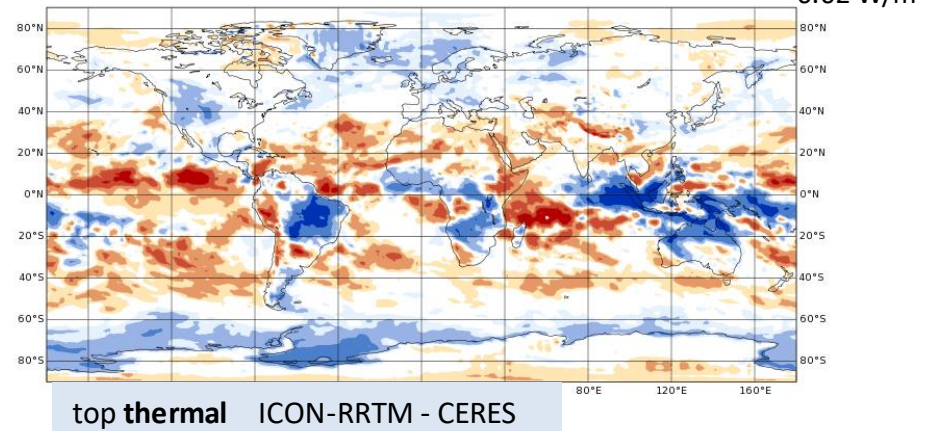
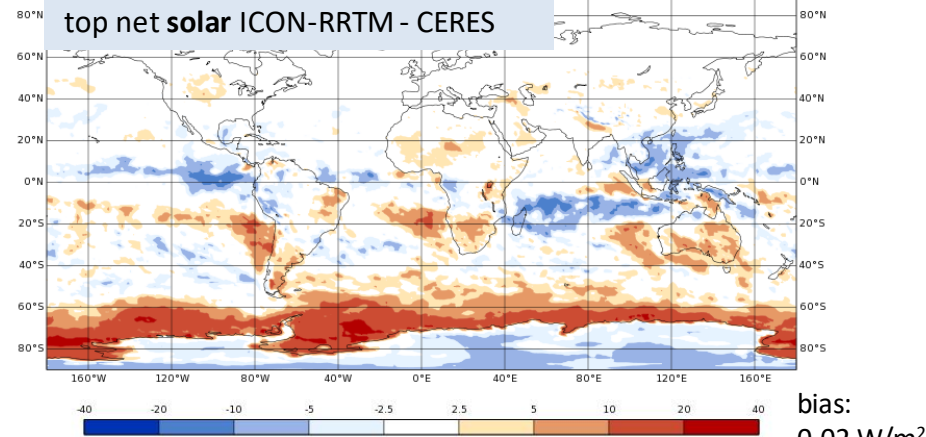
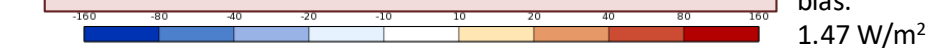
*Sophia Schäfer¹, Martin Köhler¹, Robin Hogan^{2,3}, Carolin Klinger⁴,
Daniel Rieger¹, Alberto de Lozar¹*

¹Deutscher Wetterdienst, ²ECMWF, ³University of Reading, ⁴Ludwig-Maximilians-Universität München

Issues in current RRTM radiation scheme



Error in global ICON 24h forecasts (Jan. 2018)



- **Gas optics:** RRTMG (Iacono et al. 2008)

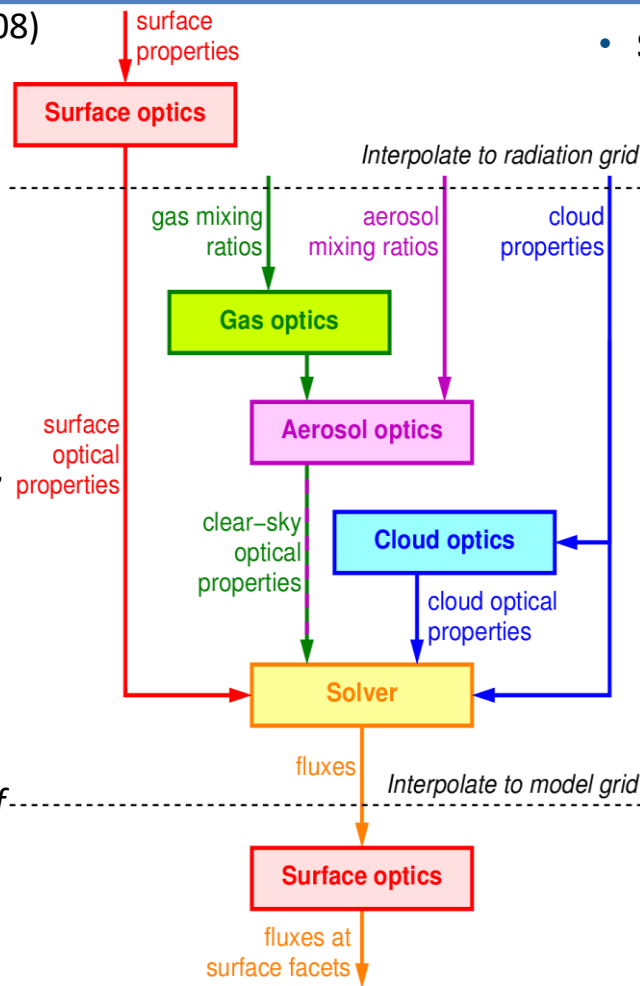
- Plan to develop new scheme with fewer spectral intervals

- **Aerosol optics:** variable species number and properties (set at run-time)

- **Cloud optics:**

- **liquid:** SOCRATES (MetOffice), Slingo (1989)
- **ice:** Fu 1996, 1998 (default), Yi et al. 2013 or Baran et al. 2014

- *Surface (under development)*
Rigorous and consistent treatment of urban and forest canopies

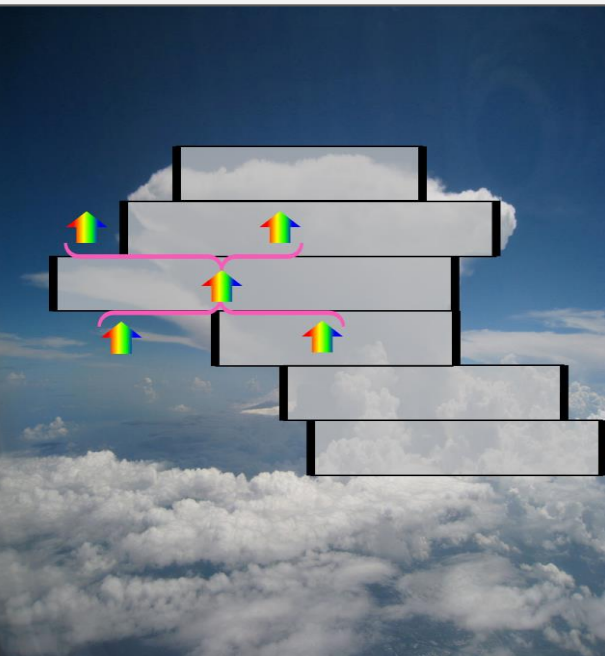


- **Solvers** for radiative transfer equations:

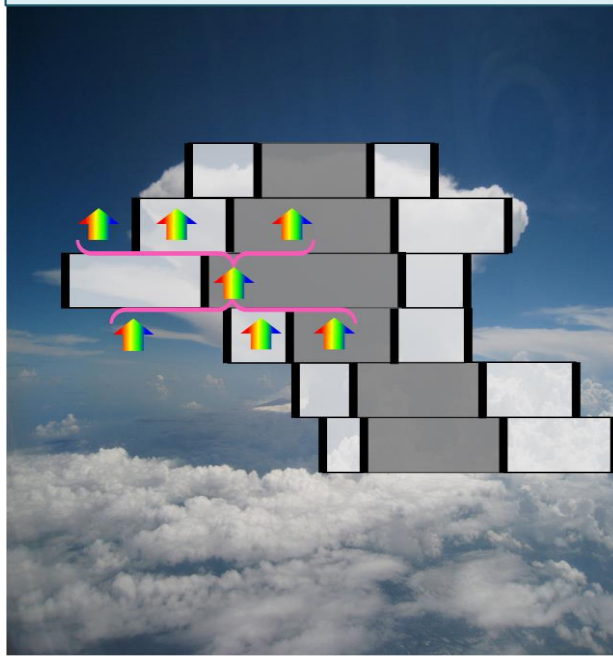
- **McICA** (Pincus et al. 2005), **Tripleclouds** (Shonk & Hogan, 2008) or **SPARTACUS** (Schäfer et al. 2016, Hogan et al. 2016)
- SPARTACUS makes ecRad the only global radiation scheme that can do **3D** radiative effects
- Longwave scattering optional
- Can configure **cloud overlap**
- **Cloud inhomogeneity:** can configure width and shape of PDF

All solvers for global models **simplify** by treating **only vertical** dimension explicitly.

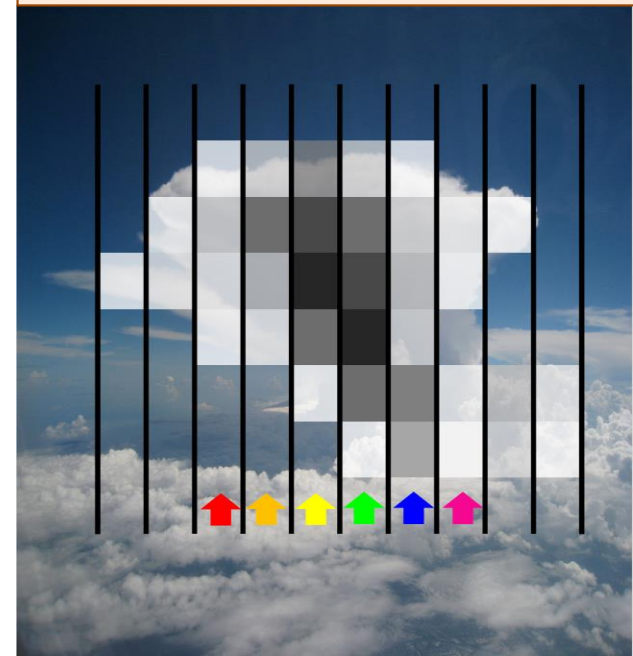
Two-stream solver: solve in **cloudy / clear regions**, partition at layer boundaries according to **overlap** (e.g. RRTM in ICON)



Tripleclouds/SPARTACUS: similar; 3 regions: **clear, thin cloud, thick cloud** → cloud inhomogeneity

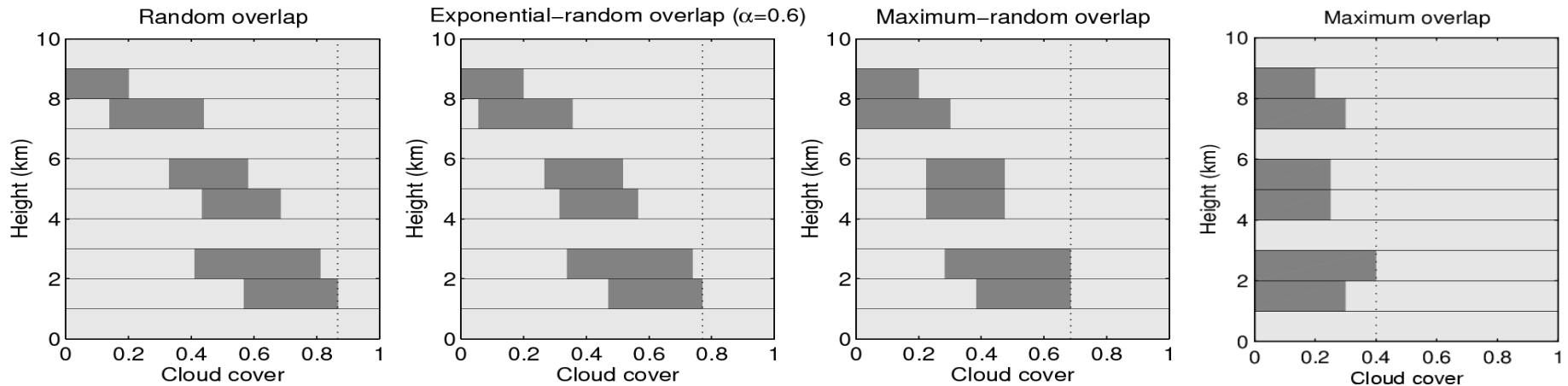


McICA: draw **random clouds** in **sub-columns** according to overlap + inhomogeneity; **distribute spectral intervals** in 1 sub-column each → **fast, random noise**



Cloud vertical overlap

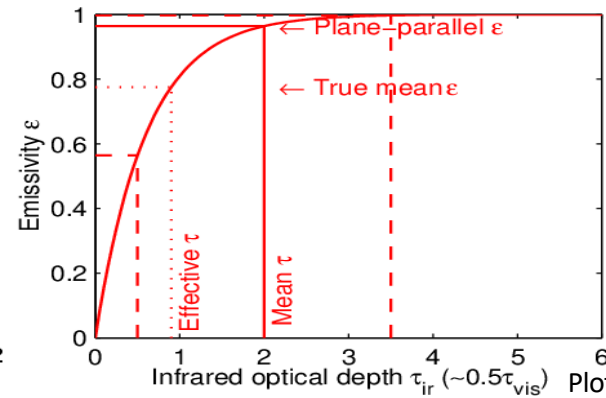
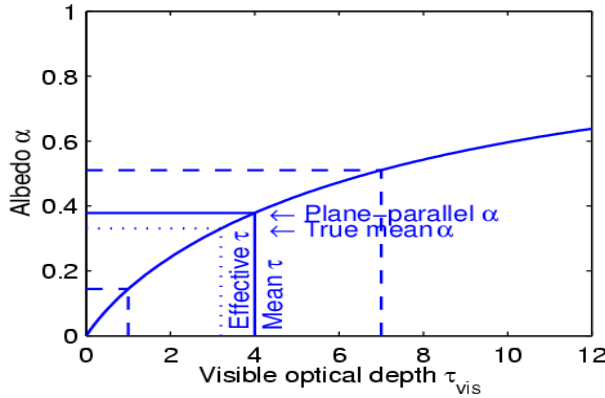
- **ICON RRTM**: different in shortwave, longwave and in total cloud cover (and bugs)
- In **ecRad**: exponential-random, maximum-random, exponential-exponential



- Based on observations (Hogan & Illingworth 2000): **exponential-random overlap**, decorrelation length ca. 1.6 km; simulation studies have 100-600 m (Neggers et al. 2011, Corbetta et al., 2015)
- **MclCA solver**: draws random number in each layer to decide cloudy/clear, numbers for neighbouring layers correlated according to overlap rules

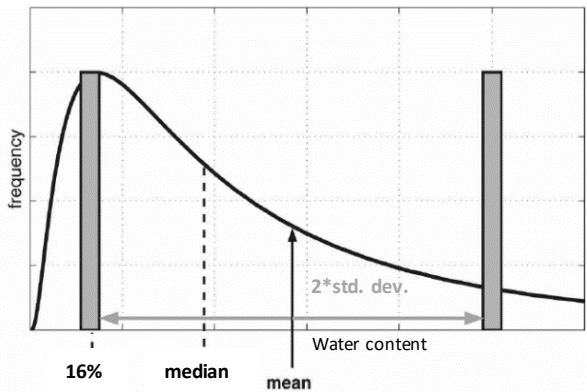
Cloud inhomogeneity

- Reflectivity and longwave emissivity are **non-linear functions** of optical depth / cloud water content



- **ICON RRTM** reduces optical depth by 0.77 (liquid) or 0.8 (ice)
- **COSMO** reduces by 0.5

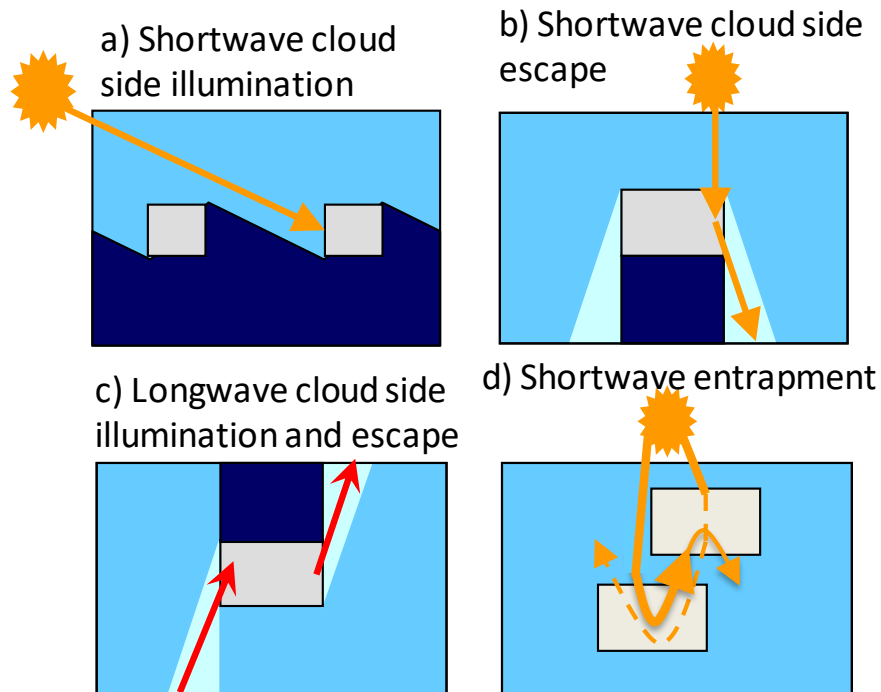
- **ecRad inhomogeneity parameters:** type of cloud water distribution (gamma / lognormal PDF), **fractional standard deviation** = $\frac{\text{standard deviation}}{\text{mean}}$



- **Tripleclouds:** two cloudy regions (equal size, preserve standard deviation of cloud water PDF)
- **MclCA:** Draw random number $\in [0,1]$ for each cloudy layer, correlated according to vertical inhomogeneity correlation; then scale with cloud water PDF value at this percentile

Adapted from Shonk & Hogan (2008)

3D effects: Physical mechanisms

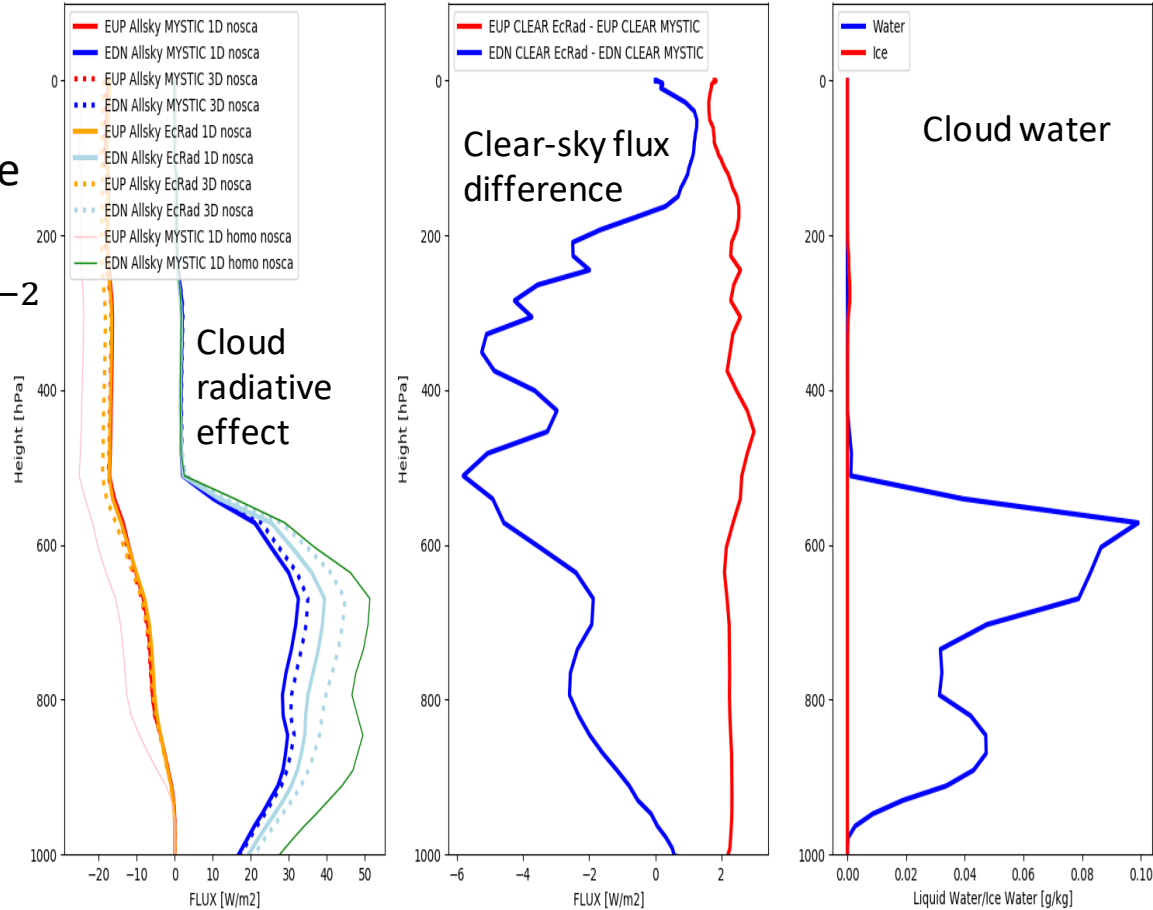


- **Shortwave cloud side illumination** increases cloud reflectivity, **cloud side escape** decreases cloud reflectivity
- **Longwave cloud side illumination and escape** increase cloud warming effect
- **Shortwave entrapment** decreases cloud reflectivity
- Similar at complex surfaces (trees, mountains, buildings)
- **Usually neglected, SPARTACUS** solver in ecRad can treat them for sub-grid obstacles

- **Magnitude of sub-grid cloud effects:** cloud inhomogeneity $\leq 30 \text{ W/m}^2$, 3D effects $\leq 5 \text{ W/m}^2$

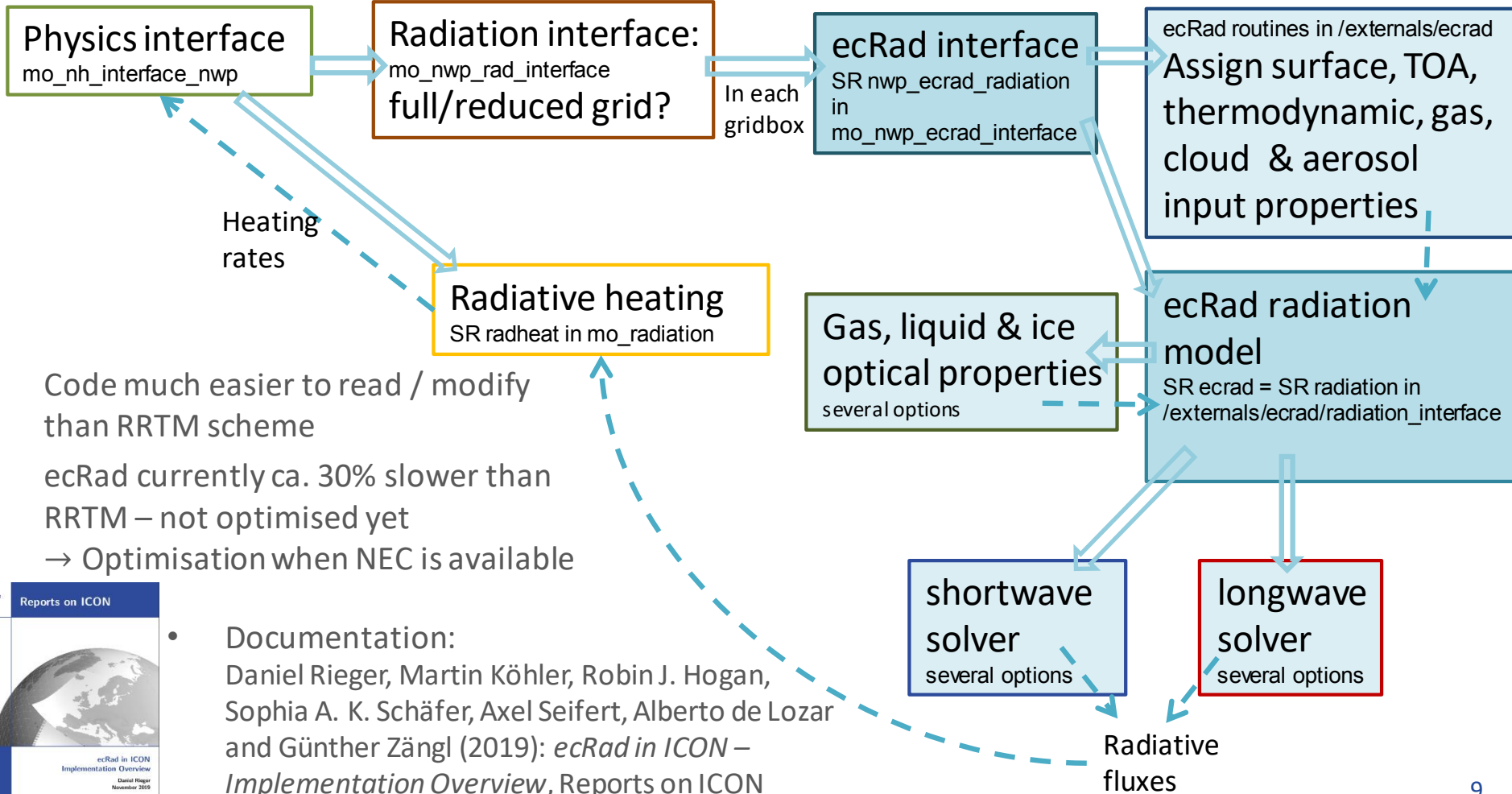
ecRad longwave evaluation (against Monte Carlo scheme)

- **Shortwave:** ecRad compares well with 3D Monte Carlo radiation (Hogan et al., 2019)
- **Longwave:** ecRad and Monte Carlo fluxes **agree well** in simple water or ice clouds, some clear-sky difference due to different gas models
- **Large uncertainty** of up to 30 Wm^{-2} due to **inhomogeneity**
- **3D effects** of up to 5 Wm^{-2}
- **ecRad captures effects;** somewhat **underestimates inhomogeneity,** **overestimates 3D effects**
- Inhomogeneity between water and ice can be important, not yet represented



Monte Carlo calculations and plots by C. Klinger

ecRad in ICON (implemented by Daniel Rieger)



- Code much easier to read / modify than RRTM scheme
- ecRad currently ca. 30% slower than RRTM – not optimised yet
→ Optimisation when NEC is available

- Documentation: Daniel Rieger, Martin Köhler, Robin J. Hogan, Sophia A. K. Schäfer, Axel Seifert, Alberto de Lozar and Günther Zängl (2019): *ecRad in ICON – Implementation Overview*, Reports on ICON



Namelist parameters for ecRad

To use ecRad, **need to specify** in configure: `./configure --with-ecrad`

+ in ICON namelist:

```
&nwp_phy_nml
inwp_radiation = 4           ! 0: no radiation, 1: RRTM, 2: RG, 3: PSRAD, 4: ecRad
&radiation_nml
ecRad_data_path = '<ICON-directory>/externals/ecrad/data'
```

Can configure model behaviour:

```
&radiation_nml
icld_overlap=2              ! Cloud overlap (in RRTM only changes sw); 1: maximum-random, 2: exponential-
                             random, 3: maximum, 4: random
irad_aero = 0               ! Aerosols; 0: no aerosol, 2: constant, 5: Tanre climatology, 6: Tegen climatology
iliquid_scat = 0           ! Liquid optics scheme: 0: SOCRATES, 1: Slingo (1989)
iice_scat = 0              ! Ice optics scheme: 0: Fu et al. (1996), 1: Baran et al. (2016)
llw_cloud_scat = .true.    ! Do longwave cloud scattering? etc.
```

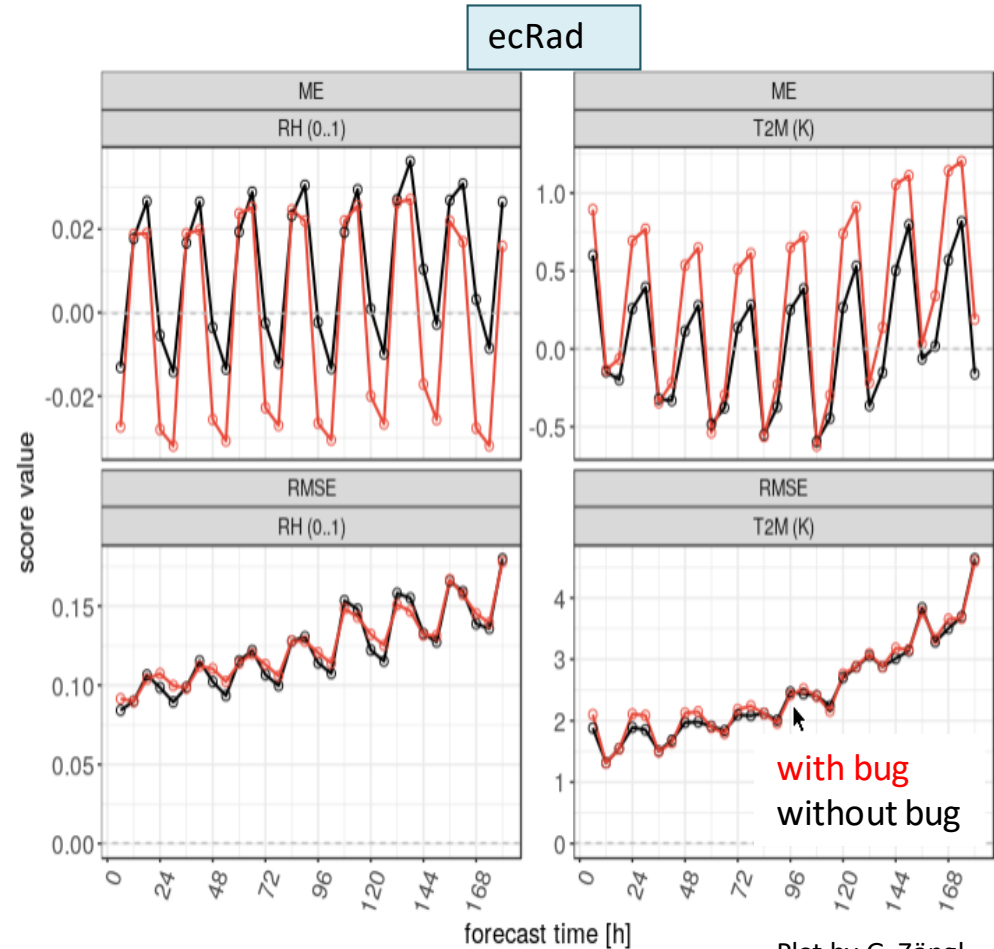
Additional ecRad namelist parameters set in SR setup_ecrad in mo_nwp_ecrad_init

```
ecrad_conf%i_solver_sw     = ISolverMcICA ! Short-wave solver
ecrad_conf%i_solver_lw     = ISolverMcICA ! Long-wave solver
ecrad_conf%do_3d_effect    = .false.     ! Do we include 3D effects?
ecrad_conf%do_lw_aerosol_scattering = .false. ! LW scattering due to aerosol etc.
```

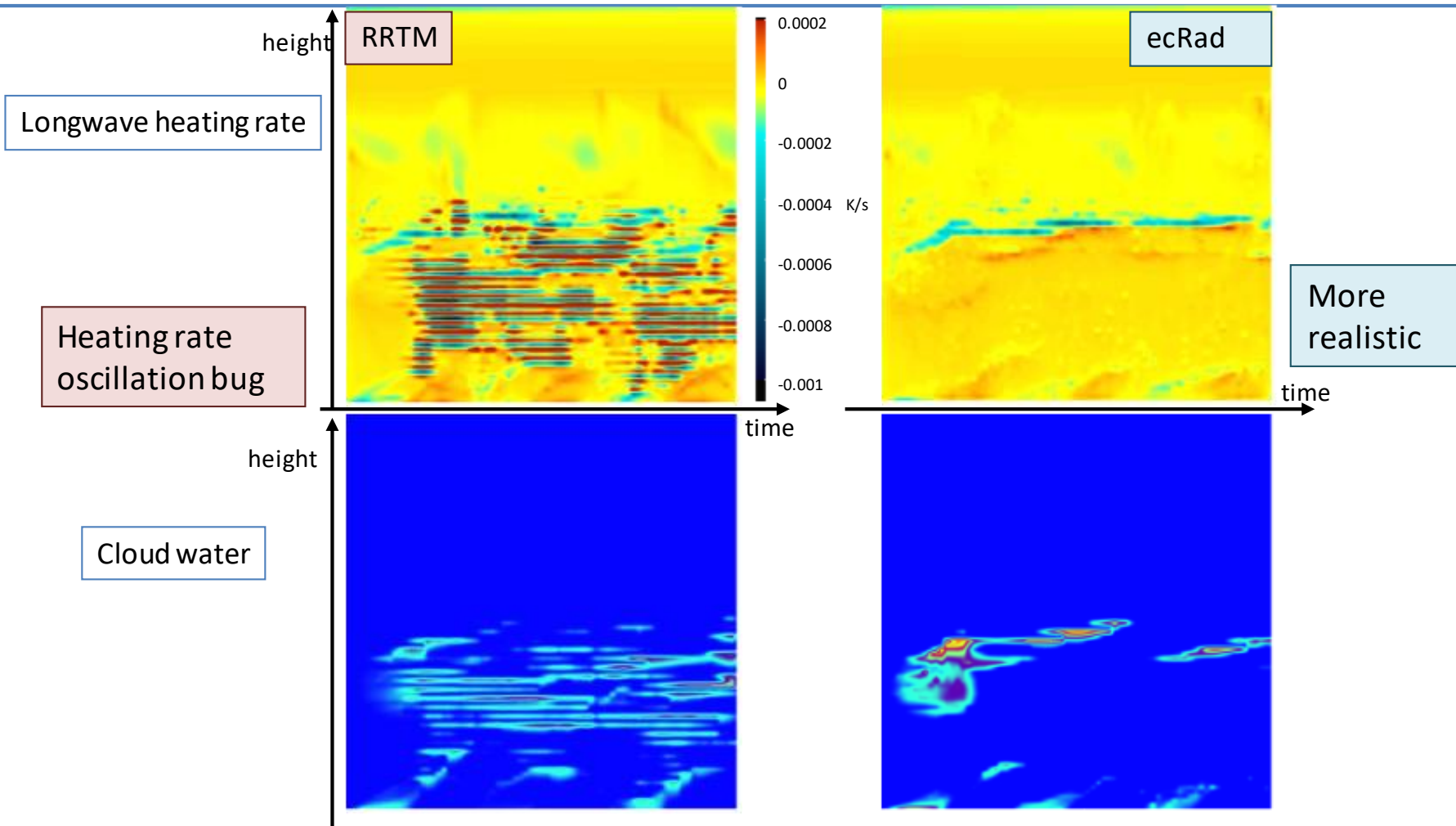
Not all combinations possible. ecRad documentation at <https://confluence.ecmwf.int/display/ECRAD>

Fixed bug in temperature input

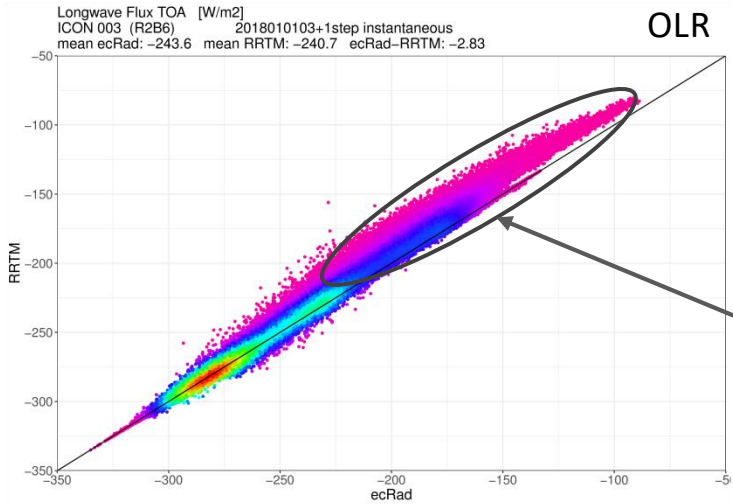
- Air **temperature in lowest level** had been set to ground temperature
→ heating rate errors
→ errors in temperature, humidity, surface fluxes,...
- Bug also in RRTM (but results less sensitive)
- **Now:** Heating rate, temperature, ... **results improve** with both radiation schemes, **ecRad better** than RRTM



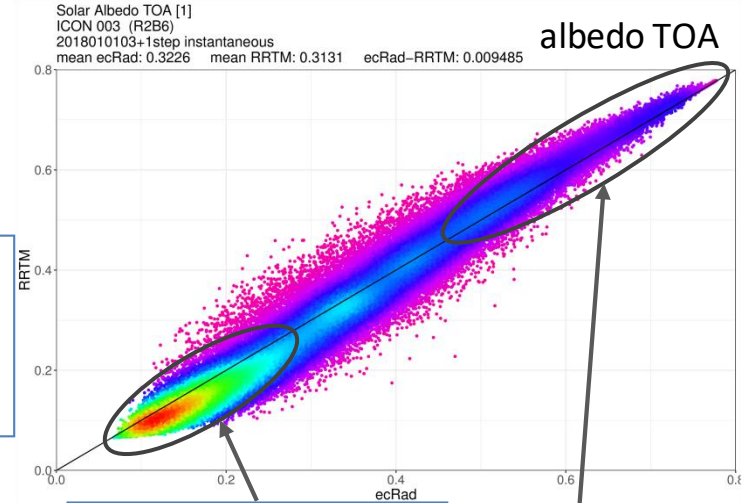
ecRad versus RRTM : ICON single column model



ecRad versus RRTM: January 2018 global one step

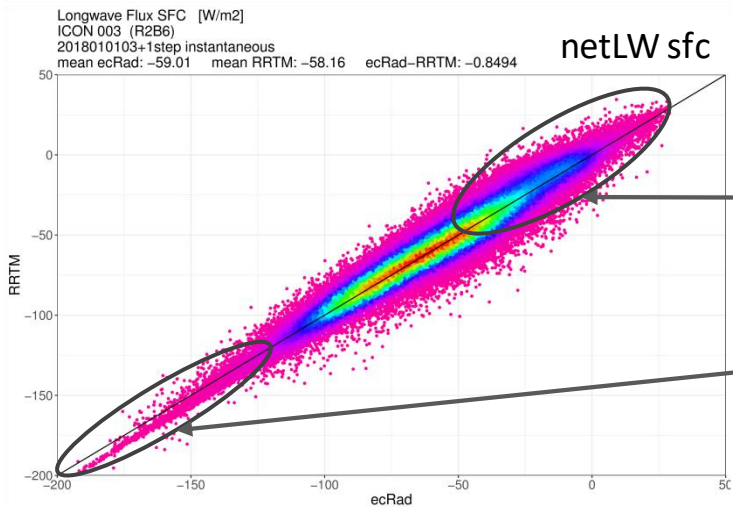


cold high clouds:
optically thinner in
LW
(offset = $\sim 15\text{Wm}^{-2}$)



thin clouds or surface:
more reflective ($\sim 2\%$)

thick clouds:
less reflective ($\sim 2\%$)



low clouds:
optically thinner ($\sim 5\text{Wm}^{-2}$) in LW

clear sky:
less surface cooling ($\sim 5\text{Wm}^{-2}$)
optically thicker dry atmosphere in
LW

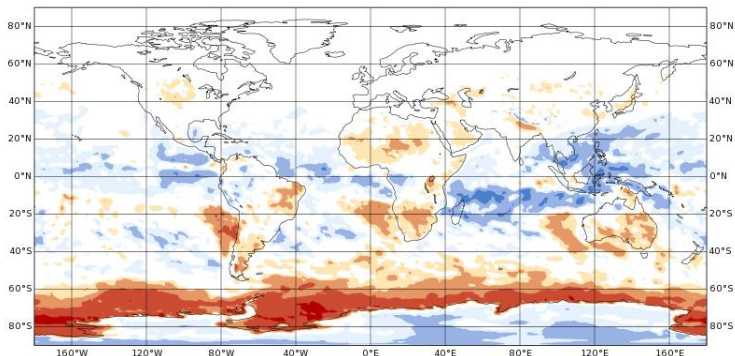
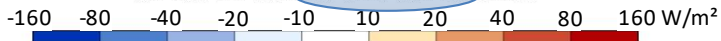
Clear-sky results agree
much closer
→ difference mainly in
cloud effects

Differences: ecRad - RRTM, 24h forecasts, January 2018

TOA solar vs. CERES

ACCSOB_T del2_375 - CERES_2018010100 to 2018010200
Min: -78.34 Max: 141.6 Mean: -0.5227 RMS: 17.51 Mem: 31

bias: -0.52 W/m²

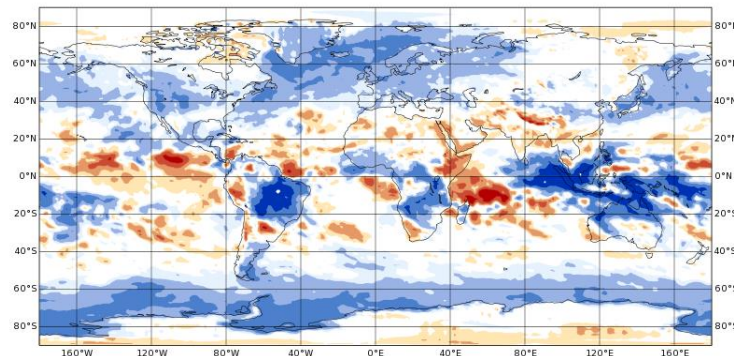
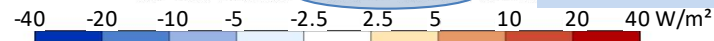


ecRad

TOA thermal vs. CERES

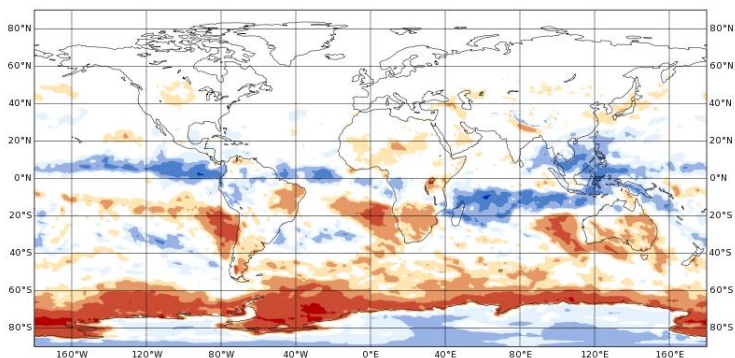
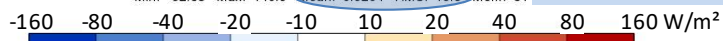
ACCTHB_T del2_375 - CERES_2018010100 to 2018010200
Min: -44.68 Max: 34.52 Mean: -2.969 RMS: 7.822 Mem: 31

bias: -2.99 W/m²



ACCSOB_T del2_374 - CERES_2018010100 to 2018010200
Min: -92.35 Max: 140.9 Mean: 0.9204 RMS: 19.3 Mem: 31

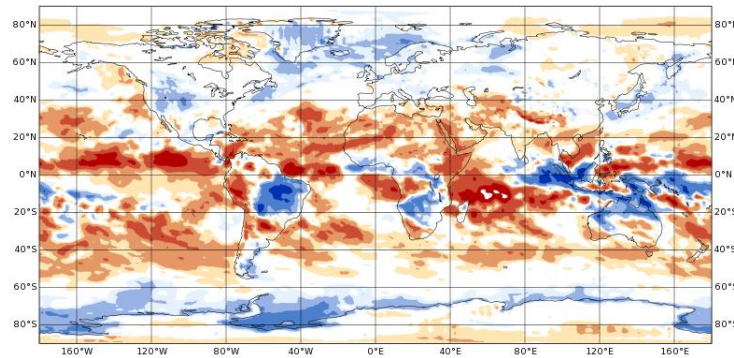
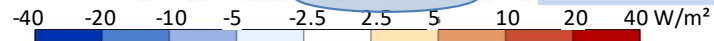
bias: 0.92 W/m²



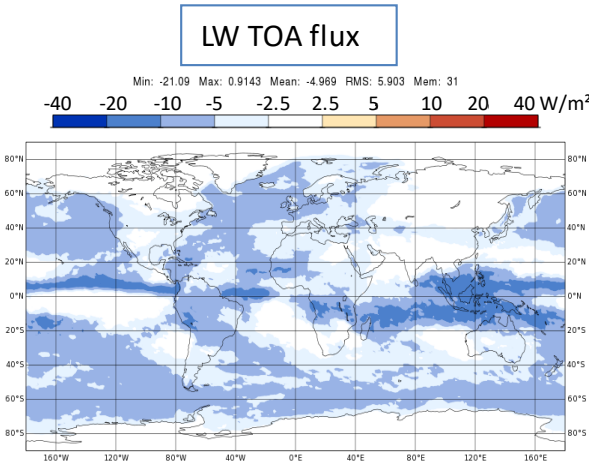
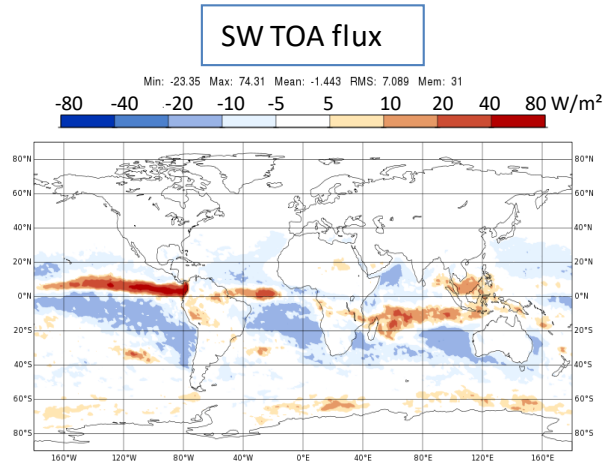
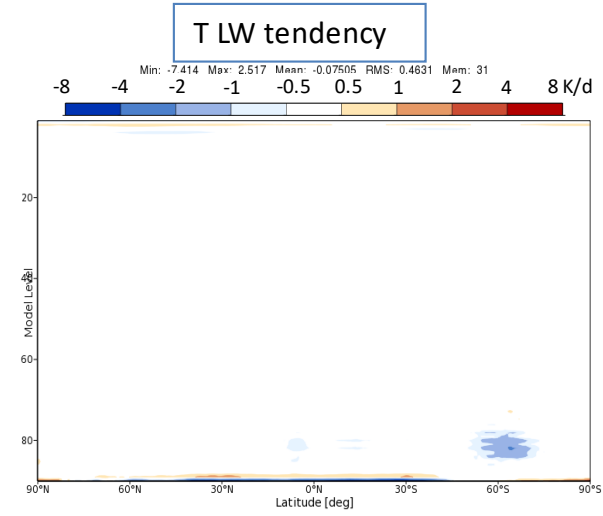
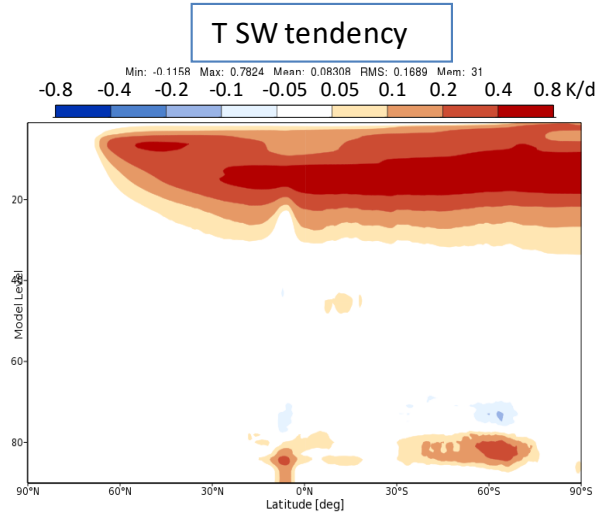
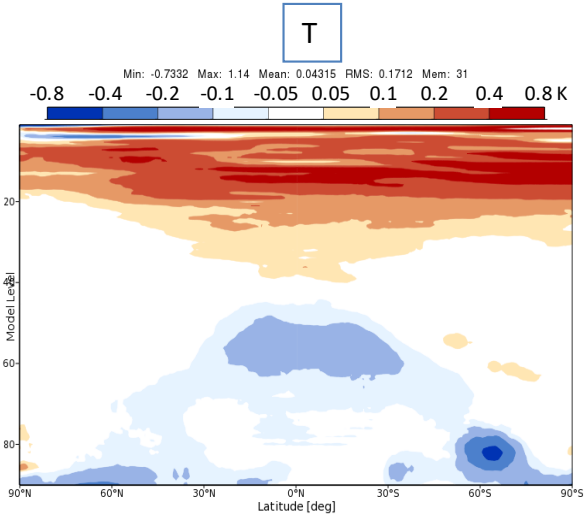
RRTM

ACCTHB_T del2_374 - CERES_2018010100 to 2018010200
Min: -38.39 Max: 50.31 Mean: 1.999 RMS: 7.625 Mem: 31

bias: 1.99 W/m²

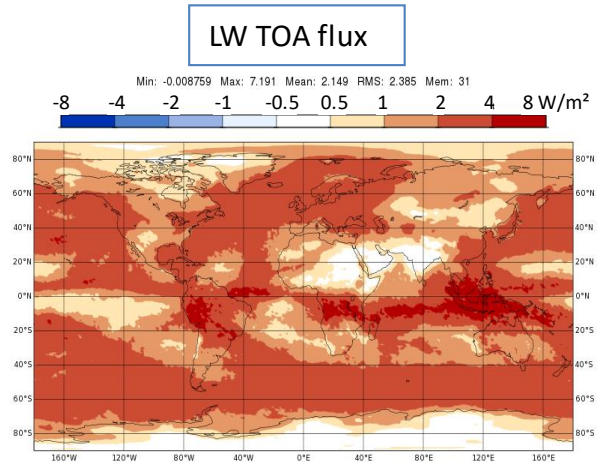
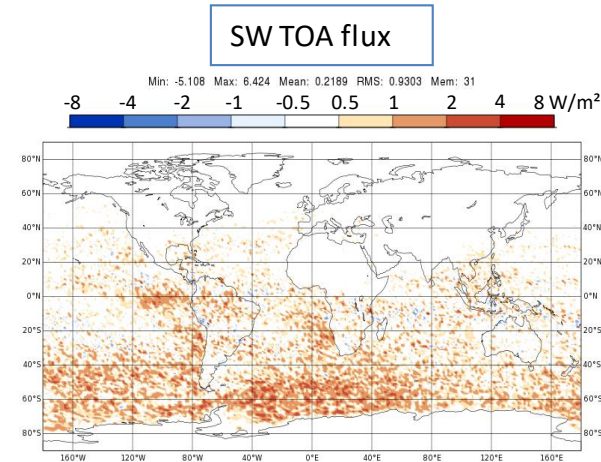
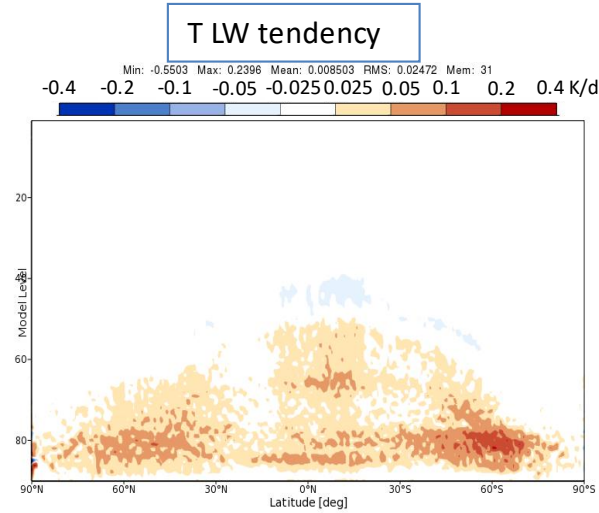
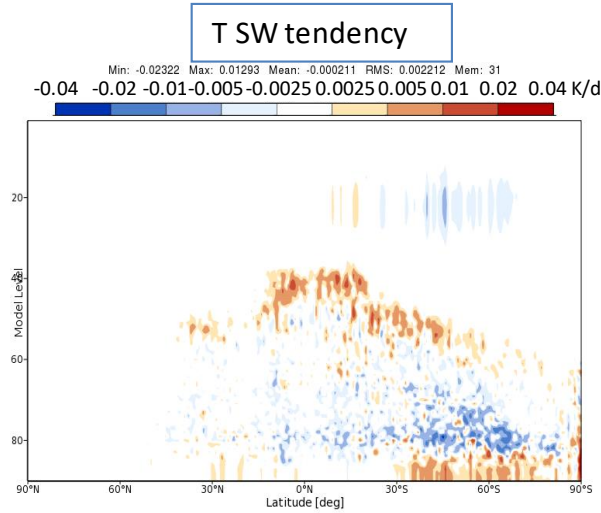
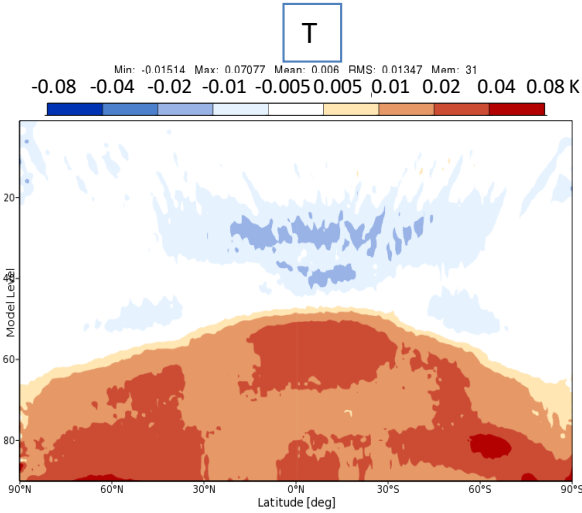


Differences: ecRad - RRTM



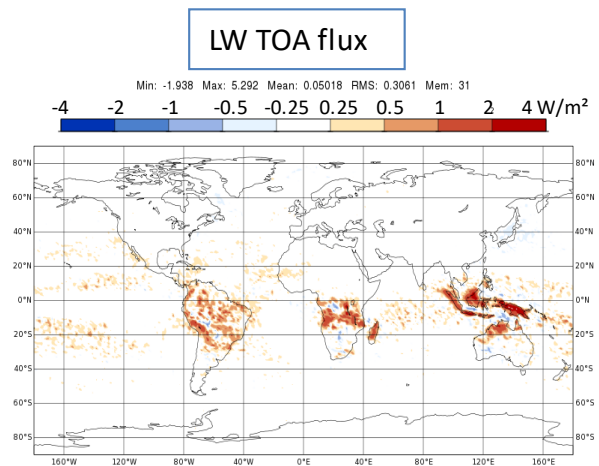
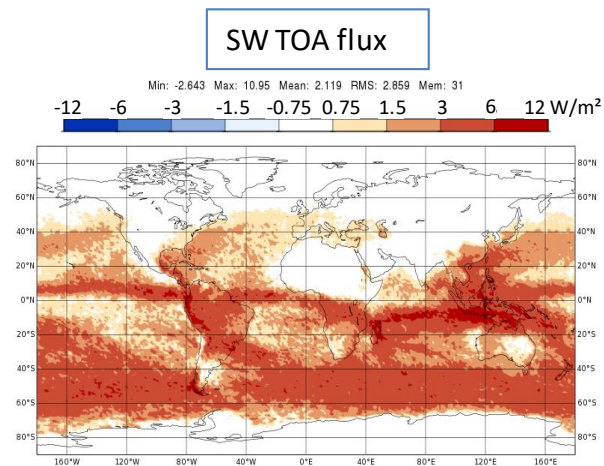
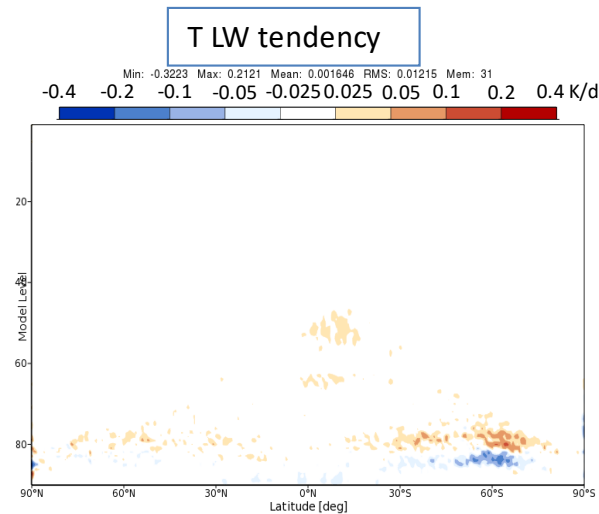
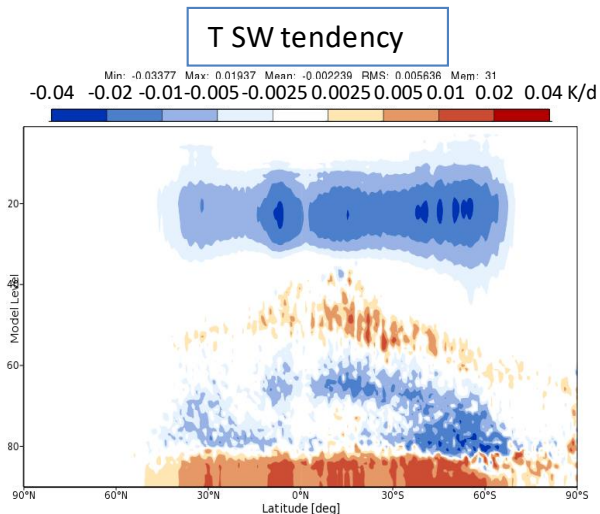
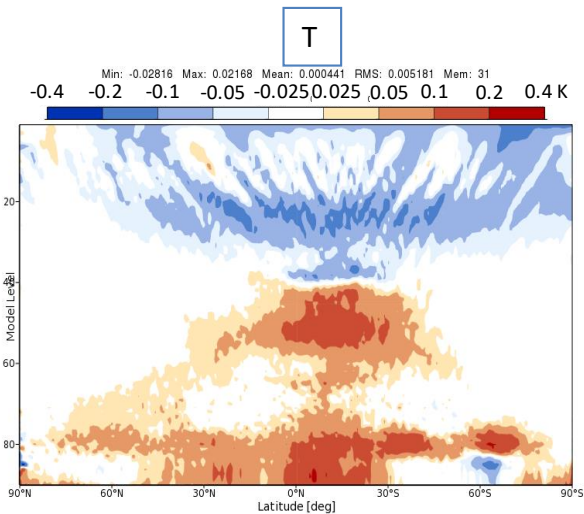
Have to check
stratospheric
heating – gas
optics difference

ecRad LW cloud scattering ON - OFF



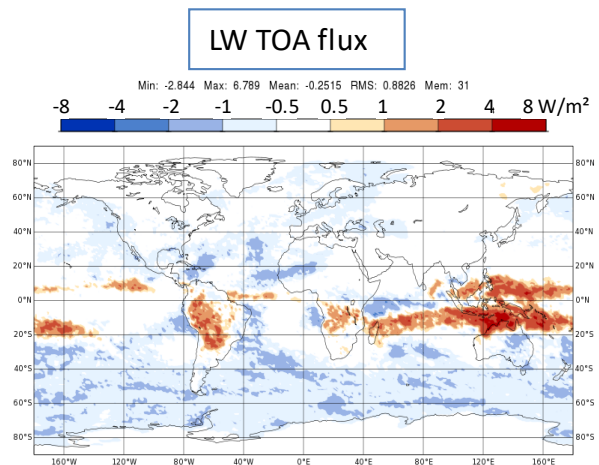
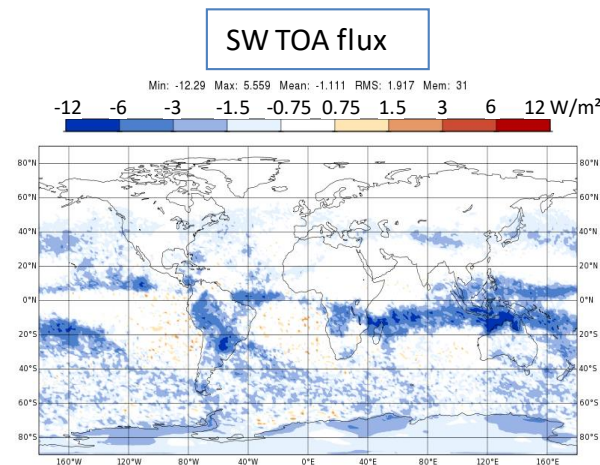
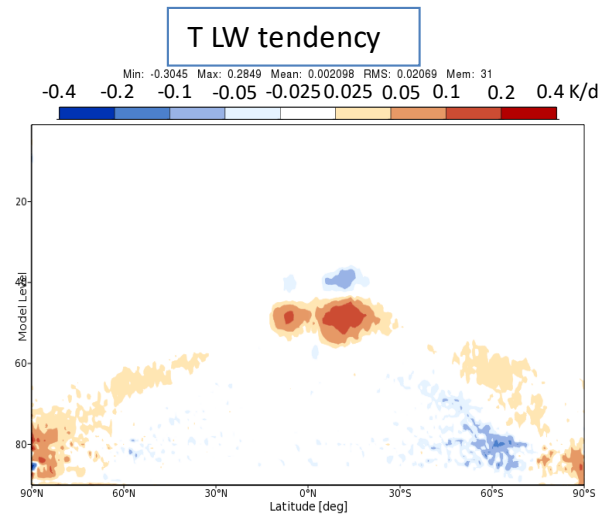
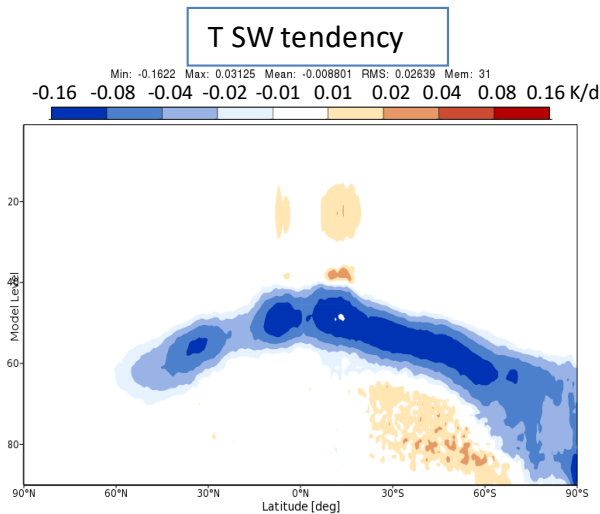
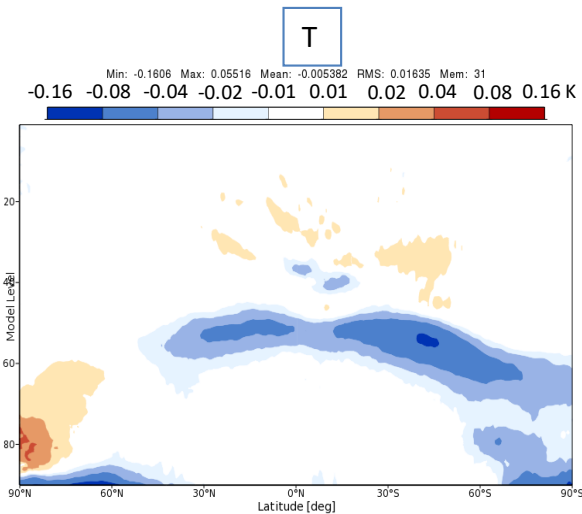
Longwave
scattering slightly
warms atmosphere

ecRad solver: TripleClouds - McICA



Global uncertainty
due to solver
assumptions

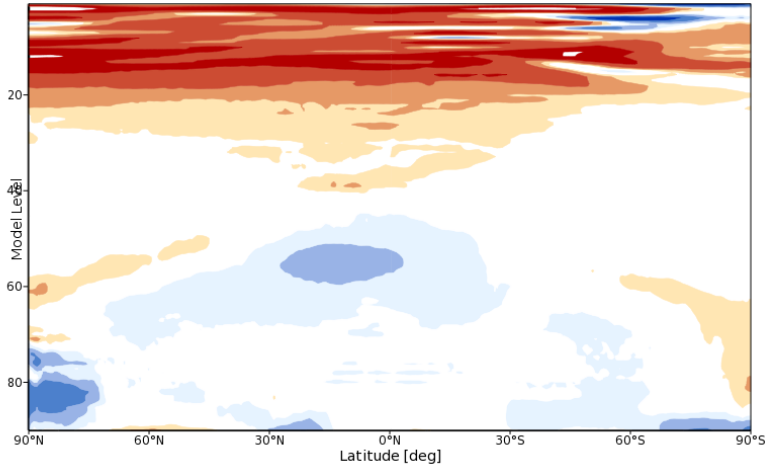
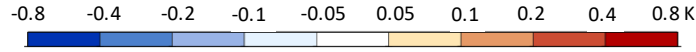
ice optics: Baran - Fu



Considerable uncertainty in ice optics assumptions

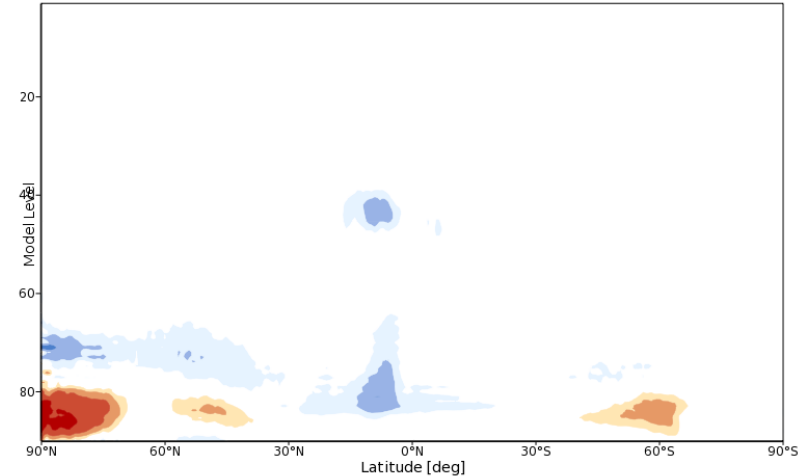
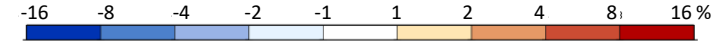
T

T ml dei2_382 - dei2_381 2018070100 to 2018070200
Min: -0.8708 Max: 1.887 Mean: 0.04883 RMS: 0.1782 Mem: 31



Cloud cover

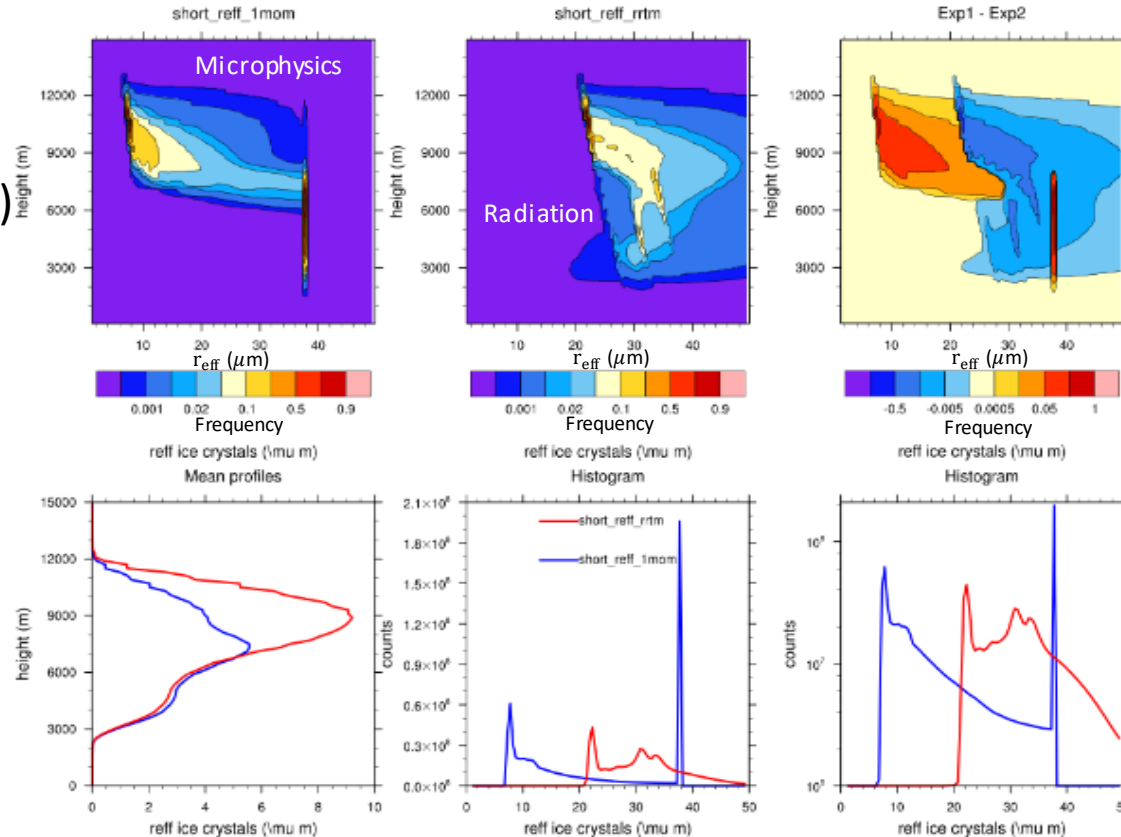
CC ml dei2_382 - dei2_381 2018070100 to 2018070200
Min: -6.899 Max: 14.44 Mean: -0.1547 RMS: 0.8798 Mem: 31



- Clouds generally optically thinner in ecRad → cooler at cloud base, warmer at cloud top
→ More cloud at base, less cloud at top
- Also less tropical convection

Input: cloud particle effective radius

- Calculated from cloud water, needs knowledge or assumptions on cloud particle size distribution and geometry
- Important for radiation (small particles dominate radiative effect)
- **Currently:** ice effective radius for radiation **inconsistent** with microphysics (liquid water better)
- Alberto de Lozar uses 1-moment- and 2-moment-microphysics assumptions to calculate effective radius for radiation
→ test radiation effect



Summary

- **ecRad implemented in ICON** (D. Rieger)
- In troposphere: **ecRad** and removal of temperature input bug **improve results** over RRTM.
- ecRad represents cloud **inhomogeneity**, SPARTACUS solver can parametrise sub-grid **3D** obstacles
- Modular scheme allows uncertainty estimation

Ongoing and future work:

- Understand cloud feedbacks, estimate uncertainties
- Evaluation of ecRad in ICON against Monte Carlo radiation calculations (with C. Klinger at LMU) and line-by-line calculations (project CKMIP at ECMWF)
- Vector optimisation for NEC@DWD
- Cloud particle effective radius parametrisation consistent with microphysics (A. de Lozar)
- Improved treatment of surface albedo and emissivity (M. Köhler, B. Fay)
- Generalise ecRad to user-defined number of cloud particle species (with R. Hogan at ECMWF)
- Extend ice optics to larger ice particles like snow or graupel (with U. Blahak and colleagues at Israel Meteorological Service)
- 3D effects of resolved (LES) clouds (with B. Mayer, C. Klinger and F. Jakub at LMU)
- Evaluation for all applications

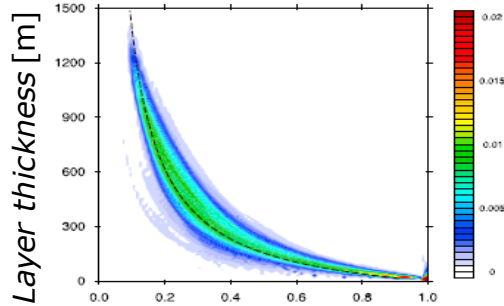
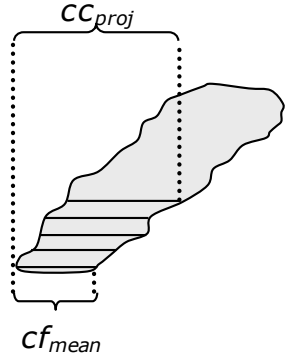
Thank you for your attention!

Contact: sophia.schaefer@dwd.de

vertical SGS cloud overlap (M.Köhler)

Neggers, Heus, Siebesma, 2011

Decorrelation length scale:
220m



$$c_{f_{mean}} / CC_{proj}$$

LES simulation of BOMEX cumulus ($dz=10m$)

Corbetta, Orlandi, Heus, Neggers, Crewell, 2015

100-600m

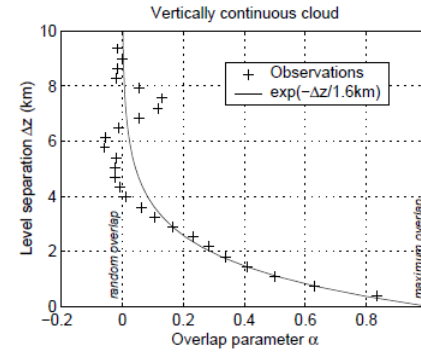
Table 1. β Parameter (m^{-1}) (Fitted Using Daily Mean R Values) and Decorrelation Length (m) on Different Days in 2013 Featuring Boundary Layer Clouds Calculated From Observations and LES Simulations for 3 and 15 min Time Resolutions

Day	3 min				15 min			
	$\beta \times 10^3$		Decorrelation Length		$\beta \times 10^3$		Decorrelation Length	
	Observations	LES	Observations	LES	Observations	LES	Observations	LES
27 Apr	4.9	4.5	590	180	-	-	-	-
19 May	5.8	6.2	157	127	-	-	-	-
5 Jun	4.7	6.5	160	148	5.2	6.3	170	202
10 Jun	4.4	4.9	253	104	4.7	5.7	213	153
20 Aug	5.3	6.6	249	120	5.0	7.0	237	239

Jülich cases LES forced by ECMWF ($dz=40m$)

Hogan, Illingworth, 2000

Decorrelation length scale:
1600m



Chilbolton, radar, $dz=360m$, $dt=1h$

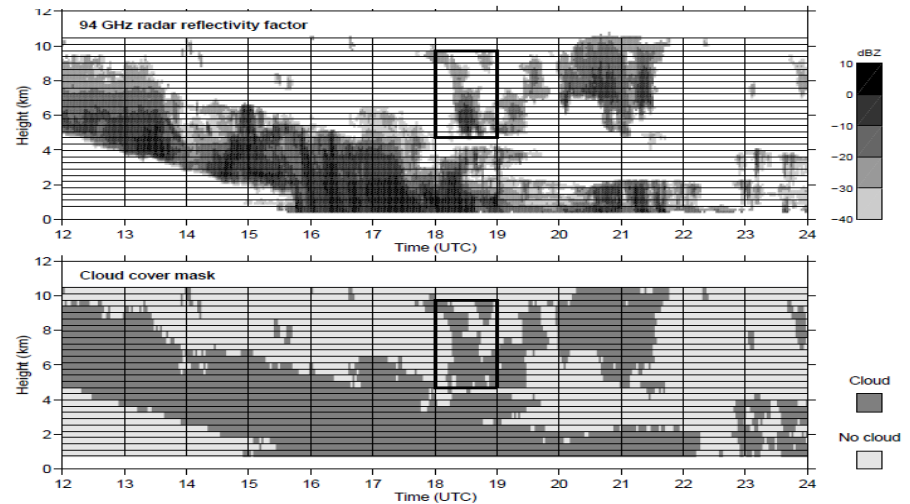
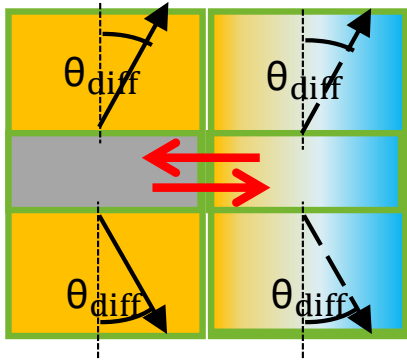


Figure 2. An example of cloud radar data used to derive the cloud cover mask, from 11 December 1998. Intermittent light drizzle was measured at the ground between 17 and 19 UTC. The resolution of the grid is 360 m and 1 hour.

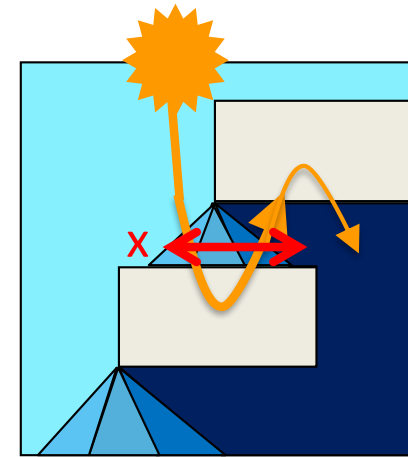
SPARTACUS: Incorporating 3D effects in a rapid radiation scheme

SPARTACUS radiation solver (SPeedy Algorithm for Radiative TrAnsfer through CloUd Sides, Schäfer et al., 2016, Hogan et al, 2016): cloud treatment based on 1D Tripleclouds solver, but includes 3D effects.

- Incorporate **3D cloud side effects** as additional terms in 2-stream calculation representing transfer between clear / cloudy regions \propto cloud side area.



SPARTACUS treatment of cloud side transfer (left) and entrapment (right).

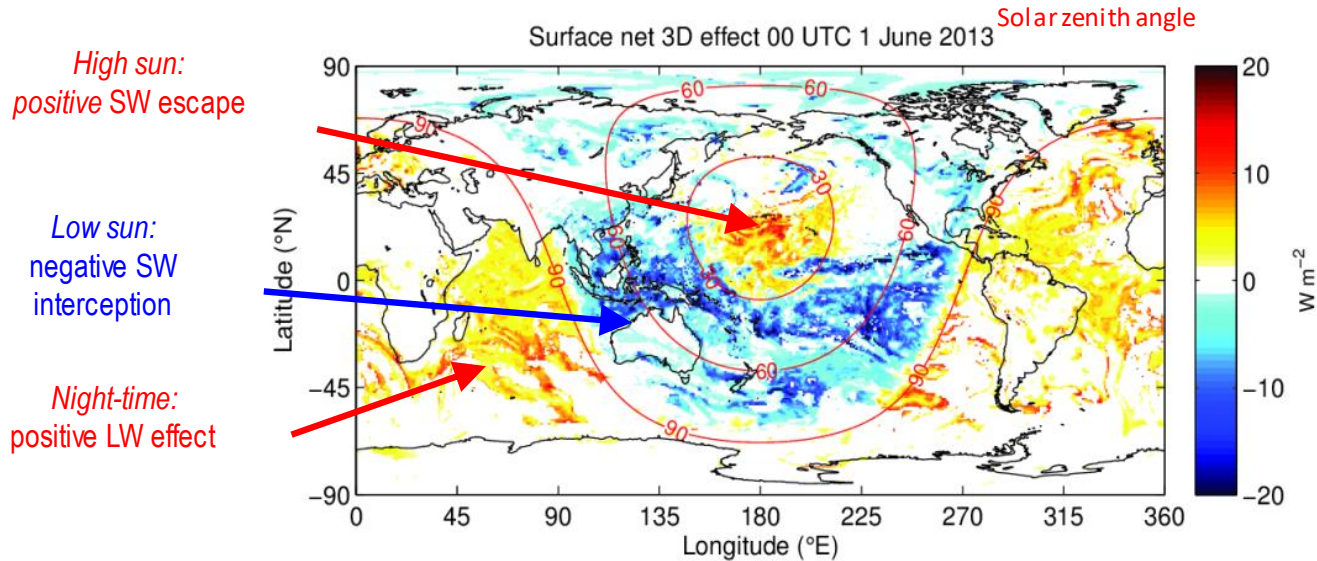


- Entrapment:** Estimate mean horizontal path travelled x .
- SPARTACUS $10^4 - 10^7 \times$ cheaper than 3D Monte Carlo, suitable for global model.

Global 3D cloud side effects

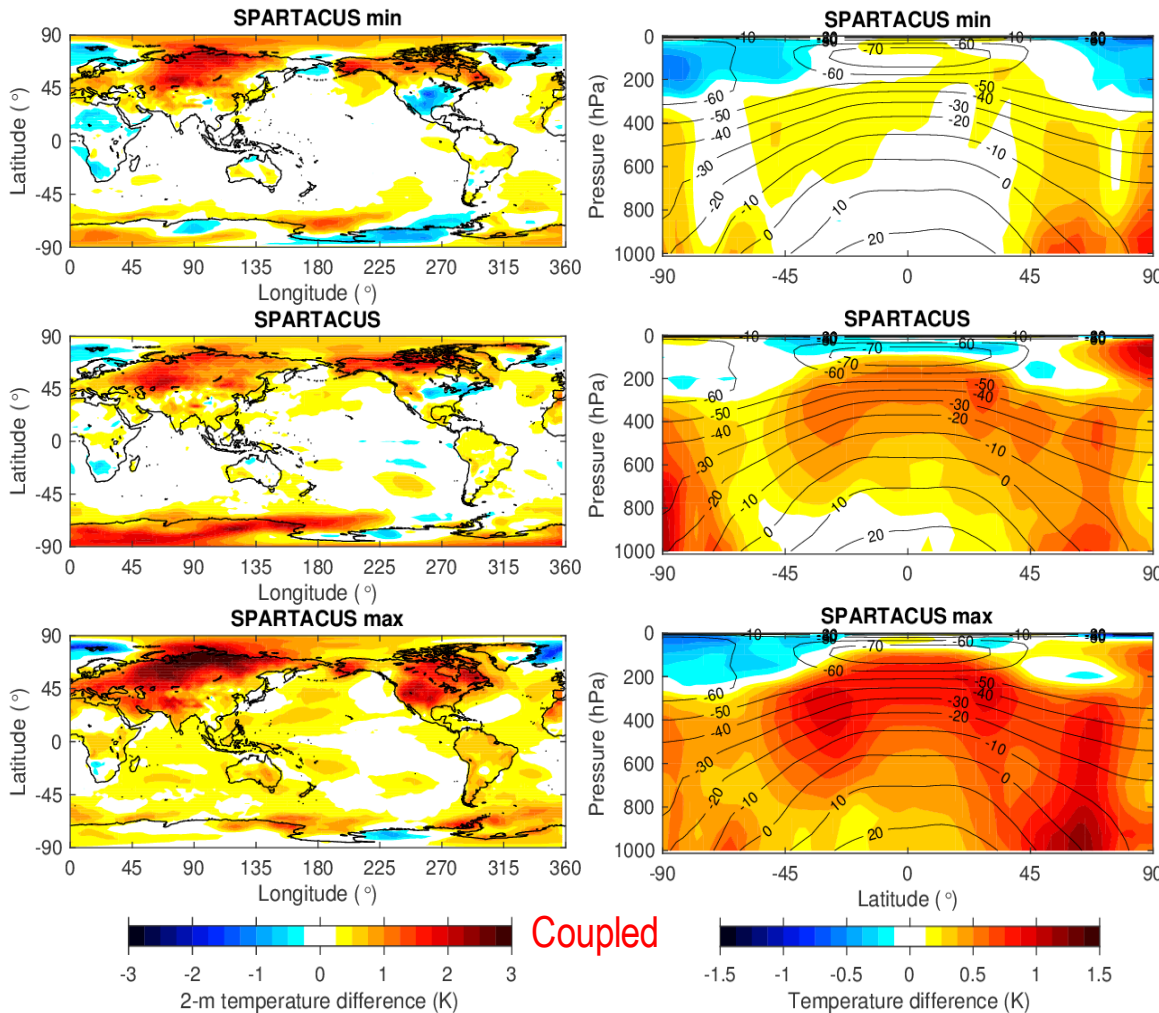
Cloud side effects (instantaneous)

- Sign of total cloud side effect depends on cloud type and solar zenith angle.
- Except for low sun: weaker than entrapment



Cloud side effects on net downward surface flux in ERA-Interim scene

Global total 3D cloud effects



Total 3D effect on climate

- **Global fluxes (net down, surface):**

Longwave $+1.6 \text{ Wm}^{-2}$,
 Shortwave $+0.8 \text{ Wm}^{-2}$,
 Total $+2.4 \text{ Wm}^{-2}$

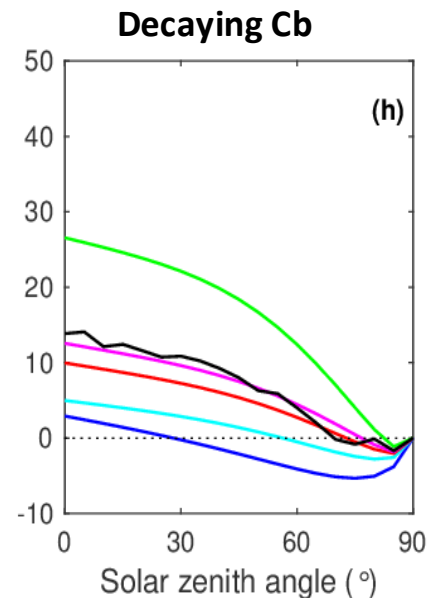
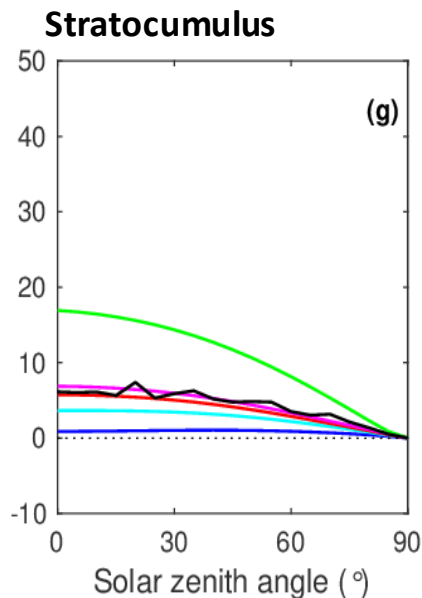
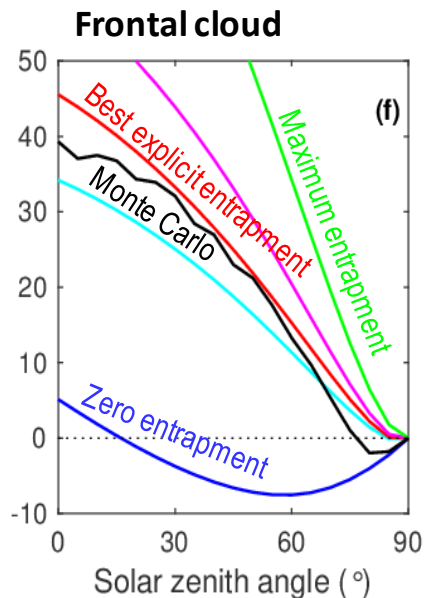
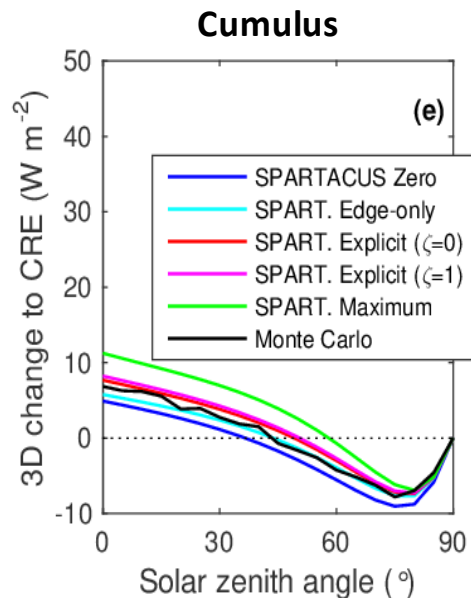
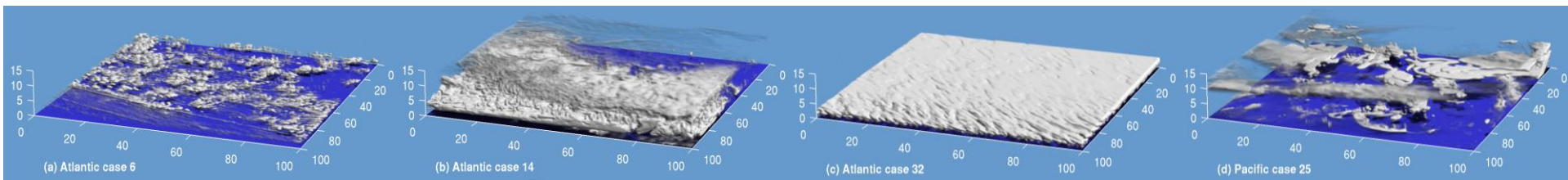
- **Temperature increases** by around 1K.
- Depends on entrainment and cloud geometry (Schäfer et al., in prep.)

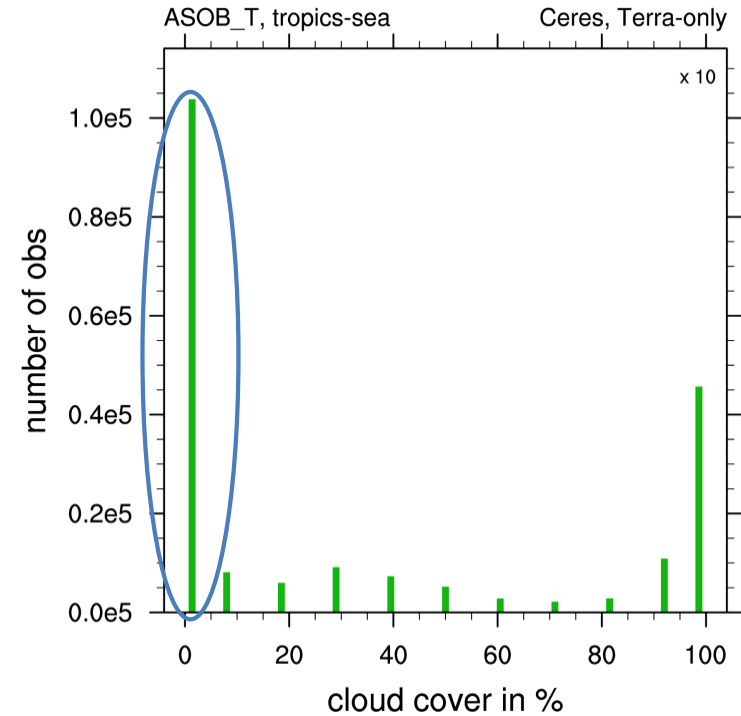
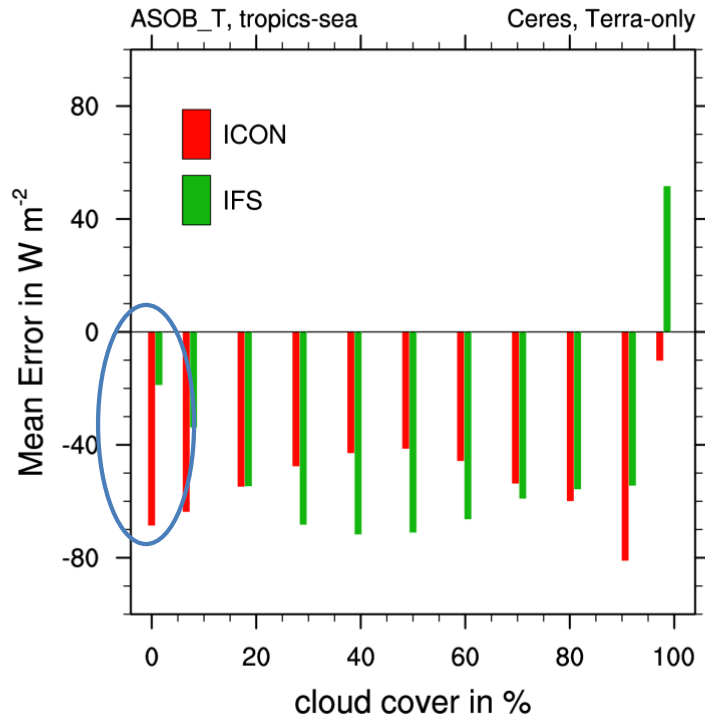
Mean 3D effect on temperature in four 1-year simulations with coupled ocean, with minimum (top) / calculated (middle) / maximum (base) entrainment.

Plots by R. Hogan

Shortwave 3D evaluation (Hogan et al., 2019, JAS)

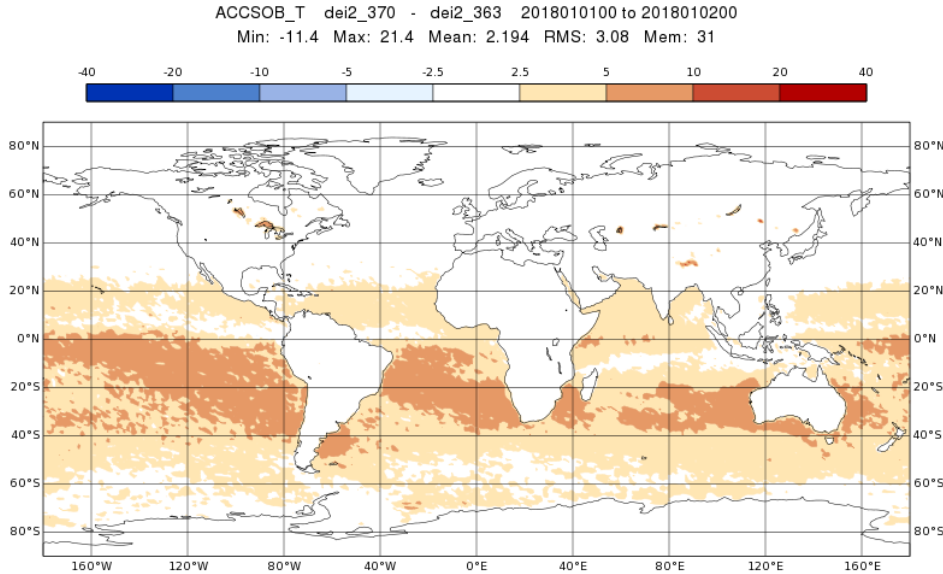
- SPARTACUS with **explicit entrainment** matches **Monte Carlo** well on average on 100 km x 100 km scenes
- Huge difference between **maximum entrainment** and **zero entrainment**





- ICON oper (13km)
- CERES TERRA (10:30am/pm at equator)
- compared colocated for consistent 25km grid

difference IFS-ICON ocean direct albedo



mean difference: 2.194 W/m²

ICON (40km), January 2018, 31 forecasts of 24h

Try other parametrisations (potentially also for waves / whitecaps)

- ICON: [Yang et al. \(2008\)](#)
- IFS: [Taylor et al. \(1996\)](#)

