Treatment of model and observation error in an ensemble data assimilation system

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Data assimilation algorithm combine forecast and observations to produce the best analysis



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Our goal: Best analysis for a prediction with the numerical model that we are using.

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Ensemble Kalman filter methods:

Step 1: Update/Analysis

Combines prior information to obtain an estimate (analysis) \mathbf{w}_k^a of the truth on discrete spatial grid and its error \mathbf{B}_k^a .

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The ensemble Kalman filter requires us to generate a number r of ensemble members $\mathbf{w}_{k}^{a,i}$, $i = 1, ..., N_{ens}$:

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Step 3: Propagation of the ensembles with full nonlinear model

$$\mathbf{w}_{k}^{a,1}, \mathbf{w}_{k}^{a,2}, \dots, \mathbf{w}_{k}^{a,N_{ens}-1}, \mathbf{w}_{k}^{a,N_{ens}} \Rightarrow \mathbf{w}_{k+1}^{b,1}, \mathbf{w}_{k+1}^{b,2}, \dots, \mathbf{w}_{k+1}^{b,N_{ens}-1}, \mathbf{w}_{k+1}^{b,N_{ens}} \Rightarrow \mathbf{w}_{k+1}^{b,1}, \mathbf{B}_{k+1}^{b}$$

The inputs to the Kalman filter are:

- An initial state at time t₀ and the corresponding covariance matrix B₀
- Observations y^o_k and observational error covariance R_k at each analysis time
- Covariance matrix of model error Q_k

We do not need to specify the covariances matrices of background error \mathbf{B}_{k}^{b} . It is generated and propagated by the filter using full nonlinear dynamics of the model.

$$\mathbf{B}_k^b pprox rac{1}{N_{ens}-1} \sum_{i=1}^{N_{ens}} [\mathbf{w}_k^{b,i} - \mathbf{w}_k^b] [\mathbf{w}_k^{b,i} - \mathbf{w}_k^b]^{ op}.$$

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$$\mathbf{B}_k^b \approx \frac{1}{N_{ens}-1} \sum_{i=1}^{N_{ens}} [\mathbf{w}_k^{b,i} - \mathbf{w}_k^b] [\mathbf{w}_k^{b,i} - \mathbf{w}_k^b]^T.$$

However, we do need to specify \mathbf{Q}_k and \mathbf{R}_k for all k as well as \mathbf{B}_0 .

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Model error

One major contributor to the background error is model error.

Appropriate model error statistics for use in data assimilation algorithms are not known.

What we consider model error:

- accuracy of numerical schemes
- unrepresented subgrid scale processes
- inaccurate forcing and boundary conditions
- representation of orography as well as parametrization uncertainty.

Model Error



from time k to time k+1 atmosphere evolves without us knowing perfectly time propagator, F^c .

from time k to time k+1 numerical model, F, propagates w^r

Model error is the difference:

$$\Pi \mathbf{F}^{\mathbf{c}}(w) - F(w^{r}).$$

Model error

- Model error in DA is assumed to be additive.
- It is defined in the model space, i.e. w^{r,true} = w^r + q
- Usually q statistics prescribed (modelled)

- Errors are assumed to be
 - white in time
 - stationary
 - without bias
 - normally distributed.

In ensemble Kalman filter algorithms, full nonlinear numerical model is used to propagate each analysis ensemble member $\mathbf{w}_{k}^{a,i}$.

Instead of only propagating the analysis ensemble to obtain the new forecast ensemble, model error \mathbf{q}_k^i can be added to:

$$\mathbf{w}_{k+1}^{b,i} = M_{k+1,k}(\mathbf{w}_k^{a,i}) + \mathbf{q}_{k+1}^i$$

where \mathbf{q}_{k}^{i} will be a sample randomly drawn using model error covariance matrix \mathbf{Q}_{k} .

This is done at analysis times!

In case nothing is done about it, the observations even for linear system might degrade the analysis (Dee 1995):











 ϵ_k^o consists of measurement error and representativeness error. It can be divided into three parts:

$$\epsilon_k^o = \epsilon_k' + \epsilon_k'' + \epsilon_k^m$$

where

$$\begin{aligned} \epsilon'_k &\equiv & \mathsf{H}^c_k \mathsf{w}(\cdot, t_k) - \mathsf{H}^c_k \mathsf{\Pi} \mathsf{w}(\cdot, t_k) \\ &= & \mathsf{H}^c_k (\mathsf{I} - \mathsf{\Pi}) \mathsf{w}(\cdot, t_k) \end{aligned}$$

 ϵ'_k – will be called *error due to unresolved scales*.

$$\begin{aligned} \epsilon_k'' &\equiv \mathbf{H}_k^c \mathbf{\Pi} \mathbf{w}(\cdot, t_k) - \mathbf{H}_k \mathbf{\Pi} \mathbf{w}(\cdot, t_k) \\ &= [\mathbf{H}_k^c - \mathbf{H}_k] \mathbf{\Pi} \mathbf{w}(\cdot, t_k) \end{aligned}$$

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 ϵ_k'' – will be called *forward interpolation error*.

Representativeness error (Lorenc 1986; Daley 1993; Cohn 1997)

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- it depends on the observation type
- ► Example: for 40 × 40 km Radiosonde/Dropsonde wind observation, observational error < 0.5 m/s and Assigned error: 2 3 m/s</p>

Data assimilation approach for dealing with observation and model error statistics

- ► parametrize the covariances $\mathbf{Q} = \mathbf{VQ}(\alpha)\mathbf{V}$ and estimate missing parameters from observation minus forecast statistics (Dee 1995, Dee and Arlindo 1997) $v_k = w_k^o H(w)$, which are assumed $\mathbf{HP}^{\mathsf{f}}\mathbf{H}^{\mathsf{T}} + \mathbf{R}$
- or from two observing system (Dee and Arlindo 1997) that simultaneously measure the same quantity.
- similar for the observation error statistics (Desrozier et al. 2005)
- or Jung et al. MWR 2012 for assimilation of radar data during forecast perturb shape parameters in the ensemble plus multiplicative and additive inflation (every 5 min *u*, *v*, *θ* perturbed with noise). Noise had std 0.5, 0.5, 0.5 with SM scheme and with DM 0.75,0.75,0.75.

Example: representativeness error





Initial state

Left: The state $w^r(\lambda, \varphi, t)$ which is being estimated. Right: The full state $w(\lambda, \varphi, t)$ from which the observations are taken.

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Cross section at $\lambda = \pi$ and estimate of $w^r(\lambda, \varphi, t)$. Upper: $\mathbf{H}_{2k}^c[\mathbf{H}_{1k}^c W_k^{uu}(\cdot, \cdot)]^T$ approximated by zero. Lower: $\mathbf{H}_{2k}^c[\mathbf{H}_{1k}^c W_k^{uu}(\cdot, \cdot)]^T$ approximated by adaptive method.

Conclusion

- Data assimilation algorithms require us to specify the statistical properties of the observation and model error.
- Both of these errors depend on the state of the atmosphere.
- Since we are searching for the best estimate for the scales that our model can represent,
- the unresolved scales are part of the model error as well as observation error.
- Not include them can have a significant impact on accuracy of the analysis.