

3D diffusion in steep terrain: testing and stability of horizontally explicit, vertically implicit discretizations

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3D diffusion in general coordinate systems

scalar diffusion equation

$$\rho \frac{\partial s}{\partial t} = -\nabla_j H^j = -\frac{\partial}{\partial x^j} H^j - \Gamma^j_{jk} H^k$$

vectorial diffusion equation

$$\rho \frac{\partial v^i}{\partial t} = -\nabla_j T^{ij} = -\frac{\partial}{\partial x^j} T^{ij} - \Gamma^i_{jk} T^{kj} - \Gamma^j_{jk} T^{ik}$$

with gradient expressions for scalar diffusion flux vector

$$H^i = -\rho K_s \; g^{ij} \nabla_j s$$

and momentum flux (stress) tensor

$$T^{ij} = -\rho K_m \left(g^{il} \nabla_l v^j + g^{jl} \nabla_l v^i \right) - \rho K_d g^{ij} \nabla_l v^i$$

These formulae cannot be used directly since many meteorological models use spherical (i.e. non-terrain-following) and normalized base vectors \rightarrow





Diffusion in spherical + terrain-following coordinates λ , ϕ , $\zeta(\lambda, \phi, r=R+z)$

scalar flux divergence: terrain following coordinates vertical ∂H^{*1} $\partial \zeta \ \partial H^{*1}$ $1 \partial H^{*2}$ $1 \partial \zeta \partial H^{*2}$ $\partial \zeta \, \partial H^{*3}$ ∂s 1 1 $r\cos\phi \,\partial\lambda$ $\partial \lambda$ $\partial \phi$ $r \partial \phi$ ∂z $\partial \zeta$ $r\cos\phi$ r ∂C $\tan \phi H^{*2}$ $\frac{2}{-}H^{*3}$ horizontal (cartesian) earth curvature

diffusion flux vector of scalar s (physical components):

$$\begin{aligned} H^{*1} &= -\rho K_s \frac{1}{r \cos \phi} \left(\frac{\partial s}{\partial \lambda} + \frac{\partial \zeta}{\partial \lambda} \frac{\partial s}{\partial \zeta} \right) \\ H^{*2} &= -\rho K_s \frac{1}{r} \left(\frac{\partial s}{\partial \phi} + \frac{\partial \zeta}{\partial \phi} \frac{\partial s}{\partial \zeta} \right) \\ H^{*3} &= -\rho K_s \frac{\partial \zeta}{\partial z} \frac{\partial s}{\partial \zeta} \end{aligned}$$

analogous: ,vectorial' diffusion of *u*, *v*, *w*

Baldauf (2005), COSMO-Newsl. Nr. 5



Basic discretization strategy for 3D diffusion:

(here: the diffusion coefficients themselves are assumed as given)

- spatially: 2nd order centered finite differences (here: staggered C-grid)
- temporally: horizontally explicit vertically implicit (HE-VI) ٠ analogous to the treatment of the Euler solver, only tridiagonal solvers are 'allowed'

HE-VI

- + good efficiency
- + good scalability \leftarrow only horizontal nearest neighb. data exchange
- metric terms in steep terrain can become unstable In fact, 3D diffusion in older versions of COSMO suffer from this problem

In the following:

- how can numerical stability be improved
- idealized tests for validation





Numerical stability





Linear von-Neumann stability analysis of 2D (i.e. x-z-diffusion) in tilted terrain

consider diffusion with constant diffusion coeff. K and constant steepness



$$\frac{\Delta h}{\Delta z} = m \cdot M_z$$

dimensionless variables: diffusion-Courant number $C_{diff} := K \frac{\Delta t}{\Delta r^2}$ steepness $m := \left. \frac{\partial z}{\partial x} \right|_{\mathcal{L}} = \frac{\Delta h}{\Delta x}$ grid anisotropy $M_z :=$







purely explicit

only 1D vertical diffusion implicit, other terms explicit, no off-centering (0.5)







only 1D vertical diffusion implicit, other terms explicit, off-centering 0.75

=old COSMO-version (until 5.2)







now: all terms with only vertical derivatives are treated implicitly



off-centering 0.6







now: all terms with only vertical derivatives are treated implicitly







Numerical stability for vector diffusion







purely explicit

only 1D vertical diffusion implicit, other terms explicit, off-centering=0.7 (=old COSMO)



(60, 50, 50, 50, 0, 0.7, 0, 0, 0, 0, 0.7, 1e-05)





now: all terms with only vertical derivatives treated implicitly









now: all terms with only vertical derivatives treated implicitly







now: all terms with only vertical derivatives treated implicitly, off-centering 0.6







Testing



Test of scalar diffusion: 3-dim. isotropic gaussian tracer distribution

3D diffusion equation with K=const.:

$$\frac{\partial \phi}{\partial t} = K \Delta \phi$$

analytic Gaussian solution for K=const.:

$$\phi(r,t) = \frac{\Phi_0}{\sqrt{4\pi K(t+t_0)^3}} \exp\left(-\frac{r^2}{4K(t+t_0)}\right),$$

$$r := \sqrt{x^2 + y^2 + z^2}$$

Baldauf (2005) COSMO-Newsl. no. 6







A proposal for an analogous test case for 3D vector diffusion:

isotropic, purely radial vector field:

$$\mathbf{v}(\mathbf{r},t) = v_r(r,t)\hat{\mathbf{e}}_r$$

with

$$v_r(r,t) = const. \frac{r}{\sqrt{K}(t+t_0)^{5/2}} \cdot e^{-\frac{r^2}{8K(t+t_0)}}$$







v,analy: min=5.64219e-07 max=48.4769 v,simul: min=5.64219e-07 max=48.4769

5.1r39_50_v_R1000km_h1000m_3dneu_3dturbT_3dmetrTi0.75_ImetrF

Vector diffusion test



v,analy: min=0.0029977 max=12.1255 v,simul: min=0.000584523 max=12.087

5.1r39_50_v_R1000km_h1000m_3dneu_3dturbT_3dmetrTi0.75_ImetrF

Vector diffusion test



for comparison: <u>without</u> metric terms

Vector diffusion test



v,analy: min=0.0338866 max=5.38961 v,simul: min=0.00519782 max=5.22789

5.1r39_50_v_R1000km_h1000m_3dneu_3dturbT_3dmetrFi0.75_ImetrF



Until now, only the pure 3D diffusion was considered.

Physics-dynamics coupling in COSMO:

physics tendencies are added to the slow dynamic processes and treated in the 3-stage Runge-Kutta time split scheme (Wicker, Skamarock, 2002, MWR)

Real test case

12 May 2015, 06 UTC run at this day COSMO-DE (2.8 km, L50) missed several convective events which produced heavy rain and intensive gusts.

Shown are results of COSMO-D2 (2.2 km, L65) at 20 UTC (i.e. after 14h forecast time)





Real case: ,12 May 2015, 06 UTC run', COSMO-D2, 1h precipitation sum

with 3D diffusion

Start time: 12.05.2015 06:00 UTC Forecast time: 12.05.2015 20:00 UTC Total precipitation [mm/1h] (shaded) C-DE 2.2km L65 5.2addMB_3dturbmetr

Geopot. at 700 hPa [gpdm] (dist. isol. 1.0 gp



difference to 1D diffusion

 Start time:
 12.05.2015
 06:00
 UTC

 Forecast time:
 12.05.2015
 20:00
 UTC

 Total precipitation [mm/1h] (shaded)

C-DE 2.2km L65 5.2addMB_3dturbmetr - C-DE 2.2km L65 5.2addMB







Real case: ,12 May 2015, 06 UTC run', COSMO-D2, 10m wind

with 3D diffusion

C—DE 2.2km L65 5.2addMB_3dturbmetr

Start time: 12.05.2015 06:00 UTC Forecast time: 12.05.2015 20:00 UTC ↓ in 10 m [m/s] (shaded)



difference to 1D diffusion

 C-DE 2.2km L65 5.2addMB_3dturbmetr - C-DE 2.2km L65 5.2addMB







Summary

- Stability analysis indicates that 3D diffusion in terrain following coordinates and HE-VI approach may be stable in *arbitrary steep* terrain if
 - use <u>all</u> terms with only z-deriv. in the tridiagonal solver
 - some off-centering (~0.7) is necessary and leads to

 $K\Delta t \left(\frac{1}{\Delta x^2} + \frac{1}{\Delta y^2}\right) \le \begin{array}{l} 0.45 \quad (\text{scalar case}) \\ 0.146 \quad (\text{vector case}) \end{array}$

- in the v-equations additional forward-backward even slightly increases stability (not implemented/planned yet)
- Testing by idealised tests with known analytic solution successfully carried out both for scalar diffusion (*Baldauf, 2005*) and by a new vector diffusion test
- New implementation runs stable in real case simulations
- Available in COSMO version 5.3
- Will be soon available in ICON (work by Slavko Brdar)
- Publication Baldauf, Brdar (accepted by QJRMS after minor corr.)











R=1000km ('real' planet)

5.1r39_50_s_R1000km_h1000m_3dneu_3dturbT_3dmetrTi0.75_ImetrT



T,analy: min=273 max=274.767 T,simul: min=273 max=274.753

5.1r39_50_s_R1000km_h1000m_3dneu_3dturbT_3dmetrTi0.75_ImetrT



M. Baldauf (DWD)



T,simul: min=273 max=273.934

5.1r39_50_s_R1000km_h1000m_3dneu_3dturbT_3dmetrFi0.75_ImetrT

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T, t=2000, iy=120



T,analy: min=273 max=273.962 T,simul: min=273 max=275.882

5.1r39_50_s_R1000km_h1000m_3dneu_3dturbF_3dmetrFi0.75_ImetrT

only 1D (vertical) diffusion